First Paleozoic *Zoophycos* trace fossils from the Sudetes (the Bardo Unit)

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INTRODUCTION

The trace fossil *Zoophycos* is a complex burrow system consisting of a large number of three-dimensional helically coiled spreiten structures, which are very variable in morphology and are present in marine sediments ranging in age from Cambrian to recent (see Knaust, 2009). Uchman (1995) proposed the term “Zoophycos group” to include all the traces that share certain common morphological characteristics. Although *Zoophycos* structures are common and widespread in ancient and modern sediments, the producing animal has not yet been discovered and the behavioural explanation is still debated (e.g., Bromley, 1991). This ichnogenus was described from diverse sediments representing various palaeoenvironments from infralittoral to abyssal (Frey et al., 1990) and even in a glaciomarine environment (Gong et al., 2008). It is well-documented in the ichnological literature that the bathymetric range of this trace has increased with time (see Bottjer et al., 1988; Kotake, 1997).

This ichnogenus was known in Poland only from the Paleogene flysch deposits of the Polish Carpathians (Wetzel and Uchman, 1998; Uchman, 1998) Jurassic carbonates of Pieniny (Tyszka, 1994) and from the Tourmaisian carbonate deposits of the Kraków Block (Hoffmann et al., 2009). This article reports the first certain discovery of *Zoophycos* in the Paleozoic rocks of the Polish Sudetes.

GEOLOGICAL SETTING

The Bardo Unit is a separate complex structure within the central Polish Sudetes (Fig. 1). The deposits containing *Zoophycos* occur within the Paprotnia Beds, which occur only in the western part of the unit and probably underlie flysch strata. The Paprotnia Beds are an informal unit belonging to an autochthonous/parautochthonous platform to foreland succession of the Bardo Unit (see Haydukiewicz and Muszer, 2002).

The exposed section, 13.7 m in thickness, is composed of claystones and mudstones with thin dark-grey micritic limestone beds and several bentonite layers, lenses and nodules of dark-grey organodetrital limestones and greywackes with intercalations of mudstones (Fig. 2). These deposits pass gradually into the polymictic Wilczka Conglomerates.

The Paprotnia Beds contain a very rich palaeontological record, that has been much studied (see Haydukiewicz and Muszer, 2002), which indicates that the beds belong to the ammonoid *G. crenistria* Zone, which corresponds with the Asian regional substage of the upper Visean (V3b). Both the lithological and palaeontological features of the Paprotnia Beds reflect gradual environmental changes from offshore to onshore conditions (Haydukiewicz and Muszer, 2002).
**SYSTEMATIC ICHNOCYLOGY**

Spreiten Burrows

*Ichnogenus Zoophycos* Massalongo, 1855

*Zoophycos isp.*

(Fig. 3)

**Material and occurrence.** – Several dozen incomplete specimens; the Paprotnia Beds (lower part of the section), the Bardo Unit (Sudetes).

**Description.** – Large spreiten burrow structures showing either simple flat forms, or unilobate helicoidal forms parallel to bedding. Most collected specimens represent parts of larger structures whose architecture is unknown. Some helicoidal forms show downward growth. Measurement of many trace fossils is difficult because of cross-cutting. The interpreted total length of the forms ranges from 70 to 160 mm. Some slabs are encircled by cylindrical fragmentarily preserved marginal tubes 1–5 mm wide. The minimal height of the trace ranges approximately from 14 to 35 mm. The lobes are filled with prominent primary lamellae, the distance between two subsequent lamellae ranging from 1 to 7 mm. The secondary lamellae occur often only in the large forms; the distance of neighbouring ones is 0.7–1 mm. The angle between primary lamellae and secondary lamellae varies from 10 to nearly 30°. The lamina is spirally coiled around a vertical axis which is approximately perpendicular to bedding. The axial tunnel is oval in outline and its diameter reaches from 12 to 30 mm. The angle between the bedding plane and the upper part of the lamina varies between 20–55°.

**Discussion.** – The *Zoophycos* described above clearly corresponds to the constructional model of Gaillard and Olivero (1993) and is very similar to Tournaisian forms from Belgium (Gaillard et al., 1999), but in the Tournaisian material does not show secondary lamellae. The Paprotnia forms and the lower Carboniferous Japanese specimens (see Kotake, 1997) are similar in having an acute angle between major and minor lamella. Our forms have also a very similar distance between two subsequent primary lamellae and the axial tunnel to the forms from the middle Pennsylvanian of Nova Scotia (McIlroy and Falcon-Lang, 2006), but their length are much smaller.

**CONCLUSIONS**

The characteristic features and organization of *Zoophycos* from the Paprotnia Beds are very close to other Carboniferous forms (Kotake, 1997; Gaillard et al., 1999; McIlroy and Falcon-Lang, 2006). Concentrations of *Zoophycos* occur in mudstones, which are underlain by organodetrital limestones. These sediments indicate distinct shallowing. The environment was located between the storm wave and fair weather wave bases, in oxygenated water. The presence of *Zoophycos* in such deposits supports the previously published opinions that the trace-making animal preferred shallow-water environments during the early Carboniferous.
We also note that ancient and recent Zoophycos-group ichnofossils, similar to Zoophycos from the Paprotnia Beds, often occur in sediments containing pebble lags, tuffs or volcanic ash (e.g., Kotake, 1997; Gong et al., 2008; Hoffmann et al., 2009). Recognition of Zoophycos trace fossils in the Carboniferous of the Sudetes will encourage further ichnological study of these deposits.

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REFERENCES


