

# Palynostratigraphy of the Culm deposits of the Moravian–Silesian zone (Poland) at Toszek Castle Hill

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The Early Carboniferous age of the flysch succession in the Moravian–Silesian zone has been based principally on lithological correlations with the stratotype goniatite-bearing section. Its uppermost part — the Kyjovice Formation in the region of Toszek (Poland) — has been studied palynologically for the first time. Miospore assemblages of the upper part of the Western European *nitidus–carnosus* (NC) miospore Zone of the early Serpukhovian have been distinguished. This agrees well with a general stratigraphic range of the Kyjovice Formation within the Goγ–E1a goniatite zones.

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## INTRODUCTION

The Early Carboniferous age of the Moravian-Silesian flysch succession in the territory of Poland, based principally on lithostratigraphical correlations with the stratotype section of the Moravia region (Unrug, 1964, 1971, 1974, 1977; Unrug and Dembowski, 1971; Kumpera, 1971a, b, c, 1977; Dvořak, 1973, 1994) has been confirmed locally, by records of rare goniatites, molluscs, orthocones and foraminifers (Żelichowski, 1962a, b; Geroch, 1971). Palynology was first used in the 1960's to recognise the stratigraphy of the uppermost part of this succession the Kyjovice Formation in the region of Toszek — but did not bring any results (Żelichowski, 1962a). However, recently resumed miospore studies have yielded new data. In selected samples of the Horny Benešov Formation (Głubczyce region), scant traces of organic matter have been recorded. The best results have been obtained from the Kyjovice Formation at Toszek (Fig. 1). Its late Viséan age within the Goy - Goniatites granosus Zone determined by Żelichowski (1962a), was based principally on lithological correlations with the Goniatite "Series" of Bobrovnik in the Nizky Jesenik Mts. (Knopp, 1929), containing an analogous fauna to that found at Toszek (Fig. 1). Preliminary palynological investigations supported Żelichowski's opinion (Trzepierczyńska, 2002), the results of further studies being presented in this paper.

# STRATIGRAPHY OF THE CULM IN THE MORAVIAN–SILESIAN ZONE

The stratotype section of the Early Carboniferous flysch deposits (Culm) in the Moravian–Silesian zone is exposed in the Nizky Jesenik Mts., the Odra Mts. and the Drahany Upland of the Moravia region, Czech Republic (Patteisky, 1929; Hokr, 1955; Havlena, 1964; Koverdynský, 1964; Kumpera, *op. cit.*; Dvořak, 1973). Natural exposures of the Culm in Poland occur in the Eastern Sudetes, east of Głuchołazy, and at the western margin of the Upper Silesia Coal Basin, in the vicinity of Toszek (Fig. 1).

Most exposures are located in the region of Ghubczyce, and the deposits from that area correspond to the part of the flysch succession from the Horny Benešov to the Hradec Formations (Żelichowski, 1962*b*; Unrug, 1977). The graywacke sandstones and shales exposed in the region of Toszek, a longitudinal, narrow area of Culm outcrops in the northern part of the zone, correspond to the uppermost part of the flysch succession (the Kyjovice Formation). They represent the northernmost and probably the youngest part of the flysch deposits cropping out in the Moravian–Silesian zone (Żelichowski, 1962*a*; Paszkowski, 1995).

Although the Culm deposits in Poland have been much studied they are still poorly understood due to scarcity of natural exposures and to unsatisfactory biostratigraphic data.



Fig. 1. Generalised geological map of the region studied (after Fusan *et al.*, 1967; Pożaryski and Dembowski, 1983; Dvořak, 1994; modified by the author)

HJMts — High Jesenik Mts.; NJMts — Nizky Jesenik Mts.; OMts — Odra Mts.; RHZ — Rheno-Hercynian Zone; USCB — Upper Silesia Coal Basin; KLFZ — Kraków–Lubliniec Fold Zone; MM — Małopolska Massif; FSB — Fore-Sudetic Block; DU — Drahany Upland

Flysch sedimentation in the Moravian–Silesian zone started in the Famennian and finished in the early Serpukhovian (Grygar and Trzepierczyński, 1995) (Fig. 2). The Andelska Hora Formation, an epimetamorphic rock succession of sandy shales with conglomerates and crinoidal limestones of Famennian–middle Viséan age, was distinguished in the lower part of the flysch succession (Kumpera and Martinec, 1995) (Fig. 2). In Poland, natural exposures of this formation occur south-east of Głuchołazy, in the vicinity of Jarnołtówek (Fig. 1).

The younger, Horny Benešov Formation is biostratigraphically documented only in Moravia. It is predominantly formed by unfossiliferous graywackes. Only plant fossils referred to *Asterocalamites scrobiculatus* (Schlotheim) indicate an Early Carboniferous age (Patteisky, 1929 *fide* Unrug, 1977), but more precise dating is based on micro- and macrofauna found at the base and top respectively (Fig. 2). In the siliceous-radiolarite shales of the Ponikiev Beds (Drahany Upland), Tournaisian and early Viséan conodonts were found (Zikmundova, 1967), whereas in the overlying and most fossiliferous Moravice Formation, goniatites of the late Viséan — Goa and Go $\beta_4$  zones were recorded (Kumpera, 1971*a*, 1977) (Fig. 2). The occurrence of *Goniatites striatus striatus* (?) (Sow.) and *Nomismoceras germanicum* Schmidt determine the late Viséan age of the Moravice Formation between the Goa and Go $\beta_4$  zones (Żelichowski, 1962*b*), whereas in Moravia, the age of the formation is between the Goa<sub>2–3</sub> and Go $\beta$  zones respectively (Kumpera, 1971*a*) (Fig. 2). In Poland, natural exposures of the Horny Benešov and Moravice Formations occur south-west of Głubczyce.

Lithologically different, but well biostratigraphically dated, the Hradec (Unrug, 1977) and Hradec-Kyjovice (Dvořak, 1973, 1994) Formations represent the uppermost part of the flysch succession in the Moravian–Silesian zone. The Hradec Formation *s.s.* (south-east and north of Głubczyce), comprising thick-bedded sandstones with intercalations of conglomerates and thin-bedded mudstones, corresponds to the late Viséan between the Go $\beta_5$  and Go $\beta_7$  zones (Żelichowski, 1962*b*), whereas in Moravia, the age of the Hradec-Kyjovice Formation is between the Go $\beta_{spi}$ –E1a zones (Kumpera, 1971*c*; Dvořak, 1994) (Fig. 2).

In Poland, natural exposures of graywacke sandstones and dark shales with abundant plant detritus, characteristic of the uppermost part of the flysch succession — the Kyjovice Formation — occur in the western margin of the Upper Silesia Coal Basin, in the region of Toszek and the Stradunia River Valley (north of Głubczyce) (Fig. 1). The stratigraphical position of these deposits was first estimated by Roemer (1870), who found *Asterocalamites scrobiculatus* (Schlotheim) and *Lepidodendron tetragonum* (?) Goeppert, which indicated an Early Carboniferous age, but this has not been confirmed by contemporary investigations. Only individual specimens of *Asterocalamites* sp. and *Lepidodendron* sp. have been described from the outcrop 3 (Toszek Castle Hill) here studied, but these have no stratigraphic significance (Żelichowski, 1962*a*).

More precise dating has been based on macrofauna. Michael (1913) recorded *Posidonia becheri* Bronn, a taxon of late Viséan age. Knopp (1929), however, suggested that the studied deposits lithologically correspond to the "Series" of Bobrovnik (the Nizky Jesenik Mts.), documented by goniatites. *Sudeticeras wilczeki-hoeferi* Patteisky, recorded in these beds, indicate a late Viséan age, the Go<sub>Y2</sub> Zone (Knopp, 1933 *fide* Żelichowski, *op. cit.*). Ruzička (1956) showed that specimens found by Michael (1913) are similar to *Posidonia corrugata*, characteristic of the late Viséan–early Serpukhovian. *P. corrugata* was also found by Żelichowski (1962*a*) in outcrop 3 together with *Posidoniella minor* Brown and *Orthoceras* sp.

### MATERIAL AND METHODS

Palynological samples have been retrived from the Horny Benešov (HBFm) and Kyjovice Formations (KFm). Fifty samples of graywacke sandstones and shales with plant detritus from the exposures at the Braciszów quarry (south-west of Głubczyce, HBFm) and Toszek Castle Hill (western margin of the USCB, KFm) were taken (Fig. 1).

The standard method of sample processing for Carboniferous palynomorphs was used.

Crushed 10 g samples were processed using 96% nitric acid and 40% hydrofluoric acid. One thousand and five hundred microscope slides were analysed, and useful palynological data were obtained. Though the organic matter was metamorphosed, miospores were preserved (Figs. 3 and 4). Their number in a single slide was estimated between several and several dozen specimens. Several hundred specimens have been identified and assigned to 26 genera and 46 species.



Fig. 2. Stratigraphy of the Culm in the Moravian-Silesian zone

### PALYNOSTRATIGRAPHY

In selected samples of the Horny Benešov Formation, miospores occur rarely. Individual specimens of the Carboniferous genera: *Densosporites* sp., *Granulatisporites* sp. ? (=*Waltzispora* sp. ?) have been determined, but they have little stratigraphic significance.

The best results were obtained from the Kyjovice Formation. Twenty three rock samples were collected from the exposure at Stary Młyn 2, located at the eastern side of Toszek Castle Hill (Fig. 1). This is a two-metres thick succession of laminated dark shales with abundant plant detritus and with a layer of fine-grained, olive-gray graywacke sandstone at the base (see locality No 3, Żelichowski, 1962*a*; locality No 5, Paszkowski, 1995; Biernat, 1964). Within the interval between 0.35 m and 1.30 m above the sandstone, seven miospore-bearing levels were recognised (Fig. 5). Abundant and taxonomically diverse miospore associations of stratigraphically important taxa were obtained (Table 1). The assemblages are dominated by representatives of the late Viséan genera *Tripartites* and *Schulzospora*.



Fig. 3. Miospore assemblage of the early Serpukhovian from the Toszek Castle Hill section

1 — Waltzispora planiangulata Sullivan, 1964, sampled level A; 2 — Waltzispora cf. sagittata Playford, 1962, sampled level C; 3 — Acanthotriletes falcatus (Knox) Potonić et Kremp, 1955, sampled level A; 4 — Acanthotriletes castanea Butterworth et Willams, 1958, sampled level G; 5 — Microreticulatisporites concavus ? Butterworth et Williams, 1958, sampled level A; 6 — Triquitrites trivalvis (Waltz) Potonić et Kremp, 1956, sampled level F; 8 — Triquitrites marginatus ? Hoffmeister, Staplin et Malloy, 1955, sampled level F; 9, 10 — Triquitrites tripertitus (Horst) Sullivan et Neves, 1964: 9 — sampled level B, 10 — sampled level D; 11 — Triquitrites distinctus Williams, 1973: 13 — sampled level B, 14 — sampled level D; 12 — Triquitrites cristatus Dybová et Jachowicz, 1957, sampled level A; 17 — Tripartites complanatus Staplin, 1960, sampled level A; 18, 19 — Tripartites vetustus Schemel, 1950, sampled level F; 20 — Tripartites cristatiformis Jachowicz, 1962, sampled level A; 21 — Tripartites cristations (Horst) Sullivan de Neves, 1964; 9 = Sampled level A; 22, 23 — Tripartites trilinguis (Horst) Smith and Butterworth: 22 — sampled level A; 18, 19 — Tripartites vetustus Schemel, 1950, sampled level F; 20 — Tripartites cristatiformis Jachowicz, 1962, sampled level A; 21 = Tripartites cristatics respondenties trilinguis (Horst) Smith and Butterworth: 22 — sampled level B, 23 — sampled level A; 24, 25 — Tripartites astricus Jachowicz, 1962, sampled level A; 26 — Diatomozonotriletes trilinearis Playford, 1963, sampled level A; 27 — Diatomozonotriletes ubertus (Ischenko) Jachowicz, 1962, sampled level C; 32 — Reticulatisporites carnosus ? (Knox) Neves, 1964; sampled level F; 29 — Savitrisporites nux (Butterworth et Williams) Smith et Butterworth, 1967, sampled level C; 32, 33 — Rotaspora cf. fracta (Schemel) Smith et Butterworth, 1967; 32 — sampled level A; 30 — sampled level A; 30 — Sampled level A; 31 — Rotaspora knoxi Butterworth et Williams, 1958, sampled level C; 32, 33 — Rotaspora cf. fracta (Schemel) Smit



Fig. 4. Miospore assemblage of the early Serpukhovian from the Toszek Castle Hill section

1, 4 — *Murospora* cf. *parthenopia* Neves et Ioannides, 1974: 1 — sampled level G, 4 — sampled level D; 2 — *Murospora complicata* Ravn, 1991, sampled level F; 3 — *Murospora horrens* (Ischenko) Beju, 1970, sampled level A; 5 — *Potoniespores delicatus sensu* Playford, 1963, sampled level A; 6 — *Potoniespores* cf. *delicatus* Playford, 1963, sampled level D; 7, 8 — *Crassispora kosankei* (Potonié et Kremp) Bharadwaj, 1957: 7 — sampled level A, 8 — sampled level D; 9 — *Densosporites anulatus* (Loose) Smith et Butterworth, 1967, sampled level A; 10 — *Densosporites* sp., sampled level A; 11 — *Pseudoannulatisporites polonicus* ? Karczewska, 1967, sampled level A; 12 — *Lycospora* sp, sampled level B; 13, 14 — *Cingulizonates* sp., sampled level A; 15 – *Kraeuselisporites* sp.; sampled level A; 16 — *Discernisporites micromanifestus* (Hacquebard) Sabry et Neves, 1971, sampled level F; 17 — *Auroraspora* sp. ?, sampled level B; 18–20 — *Schulzospora primitiva* Dybová-Jachowiczowa, 1966, sampled level B; 23 — *Florinites* sp. ?, sampled level B; 24 — *Cymatiosphaera* sp., sampled level B; Magn. × 500



Fig. 5. Palynostratigraphical position of Culm deposits at the Toszek Castle Hill section

Specimens of *Rotaspora*, *Murospora*, *Diatomozonotriletes* and *Lycospora* are also common.

The assemblages include *Lycospora pusilla* and also some other species that appear in the late Viséan, such as: *Tripartites vetustus*, *Schulzospora primitiva*, *Sch. campyloptera*, *Savitrisporites nux*, *Rotaspora knoxi*, *R.* cf. *fracta*, *Murospora horrens*, *Diatomozonotriletes ubertus*, *Diatomozonotriletes trilinearis*, *Pseudoannulatisporites polonicus*?, *Acanthotriletes* 

*falcatus* and *Acanthotriletes castanea* (Figs. 3 and 4). Some relic species characteristic of the Asbian–Brigantian *tessellatus–clavata* (TC) and *nigra–marginatus* (NM) miospore zones were also recorded: *Triquitrites tripertitus*, *Triquitrites comptus*?, *Triquitrites marginatus*?, *Tripartites distinctus*, *Murospora* cf. *parthenopia*, *Potoniespores delicatus* and *Potoniespores interitorsus* (Figs. 3 and 4) (Neves *et al.*, 1973; Neves and Ioannides, 1974).

Within the interval between 0.35 m and 1.10 m above the sandstone (Fig. 5 and Table 1), individual specimens of *Crassispora kosankei*, *Reticulatisporites carnosus*?, *Florinites* sp.? and *Bellispores nitidus*? have been recorded. These are the principal and/or index taxa of the *nitidus–carnosus* (NC) miospore Zone (Clayton *et al.*, 1977).

The abundance of *Schulzospora* and *Tripartites* and the presence of the index species *Tripartites vetustus* and *Rotaspora* cf. *fracta* indicate an age not older than the latest Viséan. But the occurrence of the index species *Reticulatisporites carnosus*? and *Bellispores nitidus*? accompanied by *Crassispora kosankei* and *Florinites* sp.? suggests the upper part of the *nitidus–carnosus* (NC) miospore Zone.

#### CONCLUSIONS

The shales of the Kyjovice Formation exposed at Toszek Castle Hill belong to the *nitidus–carnosus* (NC) Zone of the Western European miospore zonation scheme (Clayton *et al.*, 1977) and indicates an early Serpukhovian — E1a age of the deposits. This agrees well with the dating based on the presence of *Posidonia corrugata* (Etheridge) indicative of the late Viséan Goα–early Serpukhovian E1a age. The results reported here show the deposits, though, to be younger than previously thought. This agrees well with the general stratigraphic range of the Kyjovice Formation in the Nizky Jesenik Mts. in Moravia (Goγ<sub>1</sub>–E1a; after Dvořak, 1994), as well as of its facies equivalent — the Malinovice Beds, the uppermost lithostratigraphic Culm unit distinguished in Silesia (Paszkowski, 1995).

The success of these studies of the Kyjovice Formation at the Toszek Castle Hill section demonstrates that palynology is an efficient method of dating the uppermost part of the Culm exposed in the entire Moravian–Silesian zone.

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#### Table 1

Miospore taxa	Toszek Castle Hill section early Serpukhovian NC ( <i>nitidus–carnosus</i> ) spore zone sampled level						
	Α	В	C	D	E	F	G
Tripartites vetustus Schemel, 1950	+	+	+	+	+	+	+
Tripartites cristatus Dybová et Jachowicz, 1957	+	+					
Tripartites complanatus Staplin, 1960	+				+		
Tripartites cristatiformis Jachowicz, 1962	+		+			+	
Tripartites cf. insignitus Jachowicz, 1962	+						
Tripartites trilinguis (Horst) Smith et Butterworth, 1967	+	+		+			
Tripartites astricus Jachowicz, 1962	+		+	+		+	
Tripartites distinctus Williams, 1973	+			+		+	+
Rotaspora cf. fracta (Schemel) Smith et Butterworth, 1967	+	+					
Rotaspora knoxi Butterworth et Williams, 1958			+				
Schulzospora primitiva Dybová-Jachowiczowa, 1966	+	+	+	+	+	+	+
Schulzospora vetusta Dybová-Jachowiczowa, 1966	+			+			+
Schulzospora campyloptera (Waltz) Hoffmeister, Staplin et Malloy, 1955	+	+					+
Pseudoannulatisporites polonicus? Karczewska, 1967	+						
Bugensipollenites ovatus? Dybová-Jachowiczowa, 1966		+					
Waltzispora planiangulata Sullivan, 1964	+	+	+				
Waltzispora cf. sagittata Playford, 1962			+				
Murospora cf. parthenopia Neves et Ioannides, 1974				+			+
Murospora complicata Ravn, 1991						+	
Murospora horrens (Ischenko) Beju, 1970	+						
Potoniespores delicatus Playford, 1963	+			+		+	
Reticulatisporites carnosus? (Knox) Neves, 1964						+	
Cingulizonates spp.	+	+	+	+	+	+	+
Triquitrites comptus? Williams, 1973							+
Triquitrites cf. tribullatus (Ibrahim) Schopf, Wilson et Bentall, 1944				+			+
Triquitrites tripertitus (Horst) Sullivan et Neves, 1964	+	+		+			
Triquitrites trivalvis (Waltz) Potonié et Kremp, 1956	+				+		
Triquitrites marginatus? Hoffmeister, Staplin et Malloy, 1955						+	
Discernisporites micromanifestus (Hacquebard) Sabry et Neves, 1971						+	
Savitrisporites nux (Butterworth et Williams) Smith et Butterworth, 1967	+						
Bellispores nitidus? (Horst) Sullivan, 1964	+						
Densosporites annulatus (Loose) Smith et Butterworth, 1967	+						
Diatomozonotriletes trilinearis Playford, 1963	+						
Diatomozonotriletes ubertus (Ischenko) Jachowicz, 1962			+				
<i>Lycospora</i> spp.	+	+	+	+	+	+	+
Crassispora kosankei (Potonié et Kremp) Bharadwaj, 1957	+			+			
Florinites sp.?		+					

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