

Laser ablation ICP-MS study of trace elements in scheelite in different types of mineralization from the Bohemian Massif

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We describe scheelite in five types of mineral deposit from the Bohemian Massif (greisen, orogenic Au, reduced intrusion-related gold systems (RIRGS), oxidized and reduced skarn) via cathodoluminescence (CL) imaging and laser ablation inductively coupled plasma mass spectrometry (LA-ICP-MS) analyses. We document variability in the scheelite composition within single grains and/or deposit types that reflects the nature and evolution of the ore-forming fluids. We provide scheelite typical trace element characteristics and chondrite-normalized REE patterns for different deposit types, confirming that scheelite can be used as an indicator mineral for their exploration. For example, scheelite from greisen-related deposits is usually rich in Sn, Sr, Y and HREE with various ranges of median values in individual deposits (0.05–8.0 ppm, 232.4–567.1 ppm, 0.06–465.0 ppm and 0.06–632.5 ppm, respectively) and with typical chondrite-normalized REE patterns showing Eu negative or positive anomaly. Under UV light scheelite from Sn-W deposits displays variable colours from whitish via yellowish to bluish. In contrast to greisen-related deposits, scheelite from Au-orogenic association is usually relative rich in Pb (median range in individual deposits: 6.1–11.2 ppm), and relatively poor in Mg, Mo and Nb. Flat and bell-type chondrite-normalized REE patterns show a pronounced Eu positive anomaly. Under UV light it is predominantly homogeneous and bluish. Scheelite from oxidized skarn is enriched in Mo, Nb and LREE (median range in individual deposits: 311.4–7214.1 ppm, 1.3–214.9 ppm, and 1.1–1375.8 ppm, respectively), and scheelite from reduced skarns is characterized by the lowest median ranges of Nb (1.7–107.8 ppm), Y (0.69–41.1 ppm), Mn (0.5–17.0 ppm) and Zn (0.04–0.252 ppm) in individual deposits, when compared to all the groups studied. Scheelite from oxidized and reduced skarn deposits displays several types of chondrite-normalized pattern: (1) LREE-enriched with a positive or negative Eu anomaly, (2) Bell-shaped REE patterns with negative or without Eu anomaly, and (3) Flat REE patterns with a negative Eu anomaly. Scheelite from oxidized and reduced skarn deposits is commonly zoned and mostly bluish under UV light.

Key words: scheelite, gold deposits, skarns, greisen-type deposits, Bohemian Massif.

INTRODUCTION

Laser Ablation Inductively Plasma Coupled Mass Spectrometry (LA-ICP-MS) spot analysis and mapping have been successfully used to document and interpret trends of trace element distribution in sulphides such as pyrite, arsenopyrite, chalcopyrite, sphalerite, molybdenite and others (e.g., Cook et al., 2009, 2013; Large et al., 2009, 2014; Sung et al., 2009; Lin et al., 2011; Thomas et al., 2011; Winderbaum et al., 2012;

Pašava et al., 2013, 2017; George et al., 2015; Caraballo et al., 2023a, b). Cathodoluminescence (CL) imaging has proved to be an effective technique for revealing internal texture, zoning and trace element distribution in scheelite (Gaft et al., 1999; Brugger et al., 2002; Poulin et al., 2016; Ding et al., 2018; Zhang et al., 2018; Sciuba et al., 2020).

Scheelite (CaWO₄) is a common accessory mineral that occurs in different types of ore deposits and, on the basis of study of REE and other trace element distribution, it has been proposed as an important indicator mineral for targeting Au (polymetallic), W and skarn-type deposits (Malec, 1985; Štemprok and Mašková, 1992; Ghaderi et al., 1999; Dostal et al., 2009; Chowdhury et al., 2011; Poulin et al., 2016; Fu et al., 2017; McClenaghan et al., 2017; Manégia et al., 2018; Zhao et al., 2018; Yuan et al., 2019; Sciuba et al., 2020; Miranda et al., 2022, 2024).

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Scheelite is also a major mineral in greisen and skarn-type deposits (Xiong et al., 2006; Ren et al., 2010; Song et al., 2014; Guo et al., 2016; Poulin, 2016; Poulin et al., 2016, 2018) and an accessory mineral in aplite, pegmatite and metamorphosed sedimentary exhalative Fe-Mn (Brugger et al., 1998; Uspensky et al., 1998) and volcanogenic massive sulphide deposits (Poulin, 2016; Poulin et al., 2016, 2018).

Scheelite and powellite (CaMoO_4) form a partial solid solution where Mo^{6+} substitutes for W^{6+} (Tyson et al., 1988). Replacement of W by Mo and Ca by Sr, Pb, Fe, Mn, Ba and REE have been reported (Cottrant, 1981; Raimbault et al., 1993; Ghaderi et al., 1999) and contents of Na, V, Nb, Ta, S, As, Pb, U, Th, Mn, Fe, Au, Ba, B, Co, Cr, K, Ni, Sb, Sc, Bi, Cu, Sn, Zn, Li, Ti and Rb have been measured in various amounts in scheelite in different types of deposits (Anglin, 1996; Eichhorn et al., 1997; Zhigang et al., 1998; Ghaderi et al., 1999; Brugger et al., 2000a, b; Brugger et al., 2002; Xiong et al., 2006; Liu Yan et al., 2007; Dostal et al., 2009; Graupner et al., 2010; Peng et al., 2010; Ren et al., 2010; Song et al., 2014; Hazarika et al., 2016; Poulin, 2016; Poulin et al., 2016; Fu et al., 2017; Sun and Chen, 2017; Sciuba et al., 2020; Miranda et al., 2022, 2024). The REE distribution patterns and total REE concentrations in scheelite play an important role in investigating their precipitation processes. Europium anomalies identified in many deposits can be used to indicate the temperature of scheelite crystallization, the elemental substitution mechanism, and Na activity during scheelite crystallization as well as the redox state (Ghaderi et al., 1999; Brugger et al., 2000a, 2008; Song et al., 2014). Anomalies in Eu values can reflect the degree of magma differentiation, and are also closely related to the pH of the fluid (Guo, Z. et al., 2016; Fu et al., 2017; Ding et al., 2018; Zhang et al., 2018; Yuan et al., 2019; Sciuba et al., 2020). In hydrothermal fluids, Sverjensky (1984) showed that Eu^{2+} predominates at temperatures above 250°C. The high REE concentrations (~10–5000 ppm; Uspensky et al., 1998) in scheelite have been used for Sm-Nd geochronology to date gold mineralization (Anglin et al., 1996; Frei et al., 1998; Uspensky et al., 1998; Kempe et al., 2001; Roberts et al., 2006) whereas the Sr and Nd isotopic compositions have been used to constrain the sources and pathways of auriferous hydrothermal fluids (Bell et al., 1989; Kent et al., 1995; Darbyshire et al., 1996; Frei et al., 1998).

Analysis of bulk mineral chemical composition of different types of scheelite from the Bohemian Massif using instrumental neutron activation analysis (INAA) was done by Malec (1985). On the basis of his collection we provide new *in situ* LA-ICP-MS data on the distribution of REE and other trace elements in scheelite from the Bohemian Massif, from greisen, orogenic gold, reduced intrusion related gold system (RIRGS), and oxidized and reduced skarn associations, and also scheelite from other associations that are not related to any of the previous groups. We also document how different the trace element chemical composition of scheelite can be within a specific type of mineralization, this variation being a function of the origin, composition and physico-chemical conditions of the hydrothermal fluids from which it precipitates. It can also be significantly influenced by the distribution of REE and other trace elements within co-existing mineral phases. We provide further evidence that scheelite chemical composition can be used to fingerprint indicator mineral provenance in exploration and mineral assessment surveys.

CHARACTERISTICS OF SELECTED TYPES OF MINERAL DEPOSITS, SCHEELITE TEXTURE AND MINERAL ASSEMBLAGE

Basic characteristics of the scheelite occurrences studied from the Bohemian Massif, representing greisen/quartz-vein-, orogenic gold-, Au-RIRGS- and oxidized and reduced skarn associations are summarized in Table 1 and their location shown on Figure 1. Major deposits are briefly described in the following text.

SN-W-BEARING GREISENIZED GRANITE AND RELATED QUARTZ VEINS

The Sn-W-(Li) greisen deposits are mostly formed by cassiterite, wolframite and zinnwaldite, and are related to strongly fractionated Variscan S and A type granite affinities within the Saxothuringian Unit along the Czech–German boundary (Erzgebirge). Scheelite is widespread in Sn-W deposits at Čínovec, Horní Slavkov-Krásno-Huber and other localities (Seltman and Štemprok, 1995). Brief characteristics of these sites, including Horní Babákov, Ovesná Lhota and Cětovaz, are given in Table 1.

The Čínovec deposit (Fig. 1) had been exploited since the 14th century. The deposit is genetically linked to a late Variscan highly fractionated A-type granitic intrusion forming the latest evolutionary stage of the volcano-plutonic system of the Teplice caldera. The ore-bearing albite-zinnwaldite granite forms a mostly hidden cupola-like body with an elliptical N–S elongated outcrop (1.4 × 0.3 km). Its upper part, up to a depth of 750 m, is formed by albite-zinnwaldite granites, while the deeper part consists of biotite granites. In the uppermost part of the pluton (the “canopy”) up to a depth of ~300 m, the granite is altered, crosscut with planar quartz-zinnwaldite veins and partly greisenized. Stockscheider (marginal pegmatite) forms a several dm- to 2.5 m-thick layer at the contact between the granite cupola and the surrounding rhyolite (Breiter et al., 2017a).

Tin and tungsten mineralization is spatially associated with a cupola of lithium-albite granite, separated from the enclosing Teplice rhyolite by a continuous pegmatite rim (Stockscheider et al., 2017a). The main minerals forming the Li-Sn-W mineralization are quartz, wolframite, zinnwaldite, topaz, cassiterite, fluorite, feldspar, hydromuscovite and scheelite, while sulphides are represented by arsenopyrite, stannine sphalerite, galena, Bi and Cu minerals (including native bismuth, bismuthinite, chalcopyrite), pyrite and other minerals (Štemprok, 1965; Breiter et al., 2017b). On the Czech side, total geological resources were calculated at ~564 × 10⁶ t of Li-Sn-W ore averaging 0.057% Sn, 0.016% W and 0.2% Li (Newall, 2015). Apart from this, the Čínovec deposit also has significant contents of Nb, Ta and especially Sc (Hreus et al., 2021), and hence differs distinctly from the greisen ores in the western part of the Krušné hory Mts. (Štemprok et al., 1995; Breiter et al., 2017a).

We studied two types of scheelite: (a) greyish-white large grains from massive greisen (SCH-47) and (b) reddish-brown large grains from vein greisen (SCH-48).

The Krásno-Huber deposit, yielded 10–12 000 t of metallic Sn during its 600 year mining history until it was abandoned in 1987. Sn-W-Li ores occur within quartz veins and stocks in greisen. The main minerals include zinnwaldite, topaz, cassiterite, wolframite, chalcopyrite, sphalerite and arsenopyrite, ac-

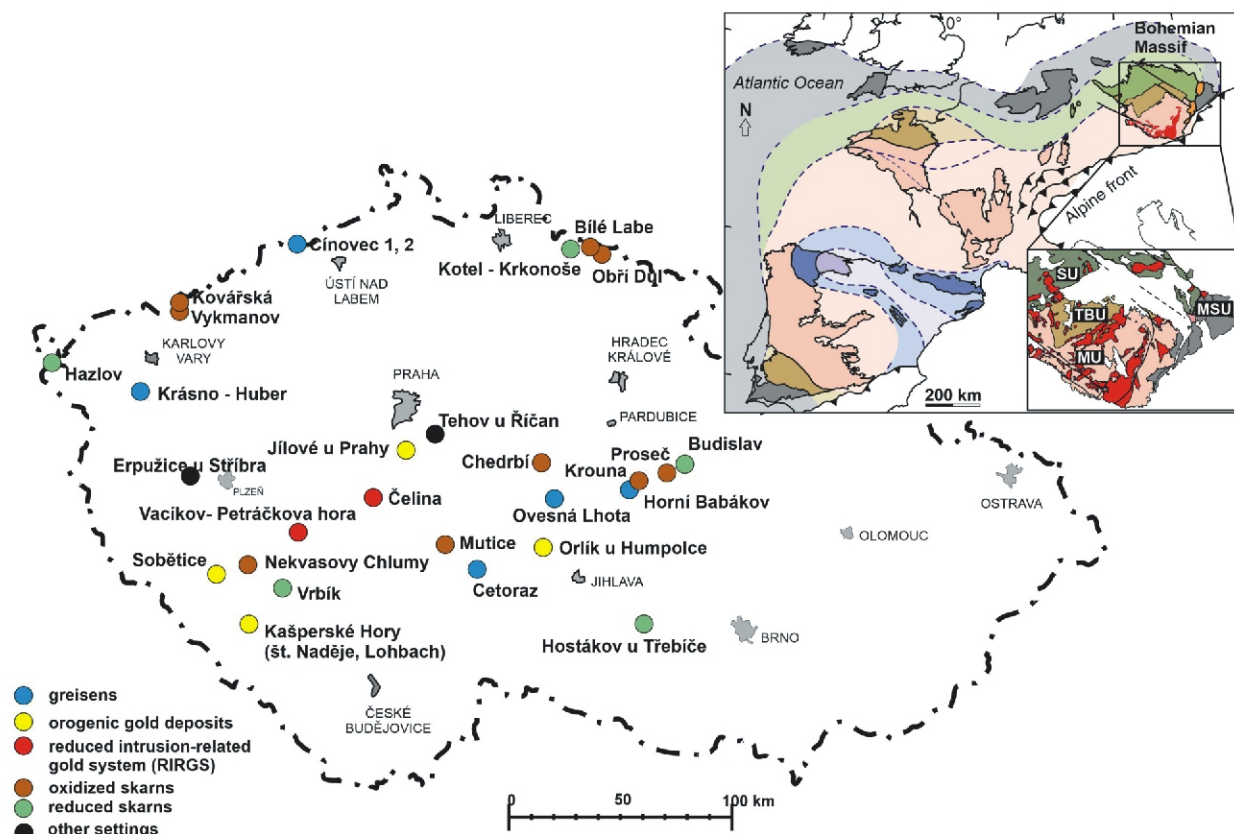


Fig. 1. Distribution of the scheelite localities studied in the Czech Republic

TBU – Teplá-Barrandian Unit, SU – Saxothuringian Unit, MU – Moldanubian Unit, MSU – Moravo-Silesian Unit, adapted from Ackerman et al. (2017)

accompanied by less frequent molybdenite, stannite, bornite, covellite and rare scheelite (Beran and Sejkora, 2006). The ore grade decreased from 1.5–1.7% Sn close to the surface, in the year 1540, to 0.2–0.4% Sn at depth in the 18th century. It is a typical greisen-type deposit, spatially and genetically linked to the younger intrusive complex of S-type acidic intrusions and especially with their Li-F rich members. This was supported by Re-Os dating of molybdenite at 323.3 ± 1.6 (Ackerman et al., 2017), indicating close temporal relationship with granites of the Eibenstock–Nejdek pluton.

We studied two 1-mm grains from greisen (SCH-43), which were very light blue and homogeneous in cathodoluminescence (CL). Scheelite is associated with chalcopyrite and wolframite.

OROGENIC AU DEPOSITS

In the Bohemian Massif, scheelite is commonly found in orogenic gold deposits (e.g., Jílové deposit, Kašperské Hory deposit) and prospects (Soběčice and Orlik u Humpolce – see Table 1 for more information; Morávek, 1995a).

THE JÍLOVÉ AU DEPOSIT

The Jílové deposit (Morávek, 1971; Morávek, 1995b, 1996) is located in the central part of the Bohemian Massif and represents a vein to stockwork type of orogenic gold deposit (Zachariáš et al., 2013). It was the most important mine in the

former Bohemian Kingdom during the medieval times. It is estimated that about 6 to 10 t of gold was mined from the Jílové deposit in the past; remaining resources are estimated to be between 13 and 17 t of Au. On the basis of the $^{40}\text{Ar}/^{39}\text{Ar}$ age of hydrothermal muscovite, the formation of the gold-bearing quartz veins hosted by reverse faults in the Jílové deposit occurred at 339 ± 1.5 Ma. The mineralization is essentially coeval with the late intrusive phases of the Central Bohemian Plutonic Complex and with late-orogenic large-scale tectonic movements at the boundary between two crustal terranes (Teplá-Barrandian and Moldanubian). Morávek (1971) identified six stages of Variscan gold-bearing mineralization in the Jílové deposit. Scheelite-bearing assemblages (stages 2–3) precipitated at $292 \pm 8^\circ\text{C}$ from a low-salinity fluid with calculated values: $^{18}\text{O}_{\text{SMOW}} = +4.2 \pm 0.5\%$ and $^{13}\text{C}_{\text{PDB}} = -11 \pm 1\%$. Some veins in the southern part of the Jílové district contain scheelite in association with dolomite and/or quartz.

The orange scheelite grain, up to 0.5 cm across (SCH-41; generally homogeneous and showing pale yellow luminescence), from a quartz vein several cm thick, comes from the third level of the Pepř mine. The quartz vein is hosted in metavolcanic rocks of the Jílové Belt (Morávek, 1971). Scheelite in the vein is locally cut by a system of tiny carbonate veinlets. According to Morávek (1971) the scheelite is younger than the main quartz matrix and locally crystallization of both quartz and scheelite was contemporaneous. The scheelite is younger than the molybdenite and older than the gold (Morávek, 1971).

Table 1

Summary of the main geological characteristics of scheelite-bearing samples

Deposit	Major metal	Mineral association**	Intrusion	Host rock	Geotectonic Unit	Age
Quartz vein/greisen deposits						
Cínovec/ Zinnwald	Sn-W-Li	Qz+Znw+Sch (We studied two types of scheelite: 1. greyish white large grains from massive greisen (SCH-47) and (2) reddish-brown large grains of scheelite from vein greisen (SCH-48).	Cínovec-Zinnwald granite (cupola) intrusion to paleorhyolite	quartz veins in strongly fractionated and slightly peraluminous rare metal A-type granites	Saxothuringicum	311–321.5 Ma
Krásno-Huber	Sn(W)	Qz+Sch (We studied two 1 mm grains from greisen (SCH-43), showing weak light blue and homogeneous texture in cathodoluminescence (CL). Scheelite is associated with chalcopyrite and wolframite	Carlsbad plutonic intrusion	quartz veins in lithium - topaz granite	Saxothuringicum	320–330 Ma
Cetoraz		Qz+Ms+Kfs+Ghn+Py+Wf+Sch	Pacov Orthogneiss	quartz veins in orthogneiss		494–485 Ma
Ovesná Lhota	Sn-W	Qz+Ms+Tur+Cst+Wf+Sch	small adamellite intrusions related to Moldanubian Pluton	quartz veins in migmatites	Moldanubicum	Variscan
Horní Babákov	W	Qz+Ms+Brl+Flr+Feb+Sch	Nasavrky pluton - biotite granodiorite	quartz veins inside granodiorite body	Bohemicum	Variscan
Oxidized skarns						
Obří Důl	Fe-Cu-As-W	Grt+Cpx+Amp+Sch	Krkonoše-Jizera Massif	dolomite marble in contact zone of I-type porphyritic biotite granite	Saxothuringicum	312 Ma
Kovářská	F-Be-W	Kfs+Cpx+Czo+Amp+Bvn+ Sch	no observable relation to granite intrusion	crystalline rocks of amphibolite facies (paragneiss, mica schist)	Saxothuringicum	Variscan
Vykmanov u Kovářské	W-Pb-Zn	Cpx+Flr+Sph+Sch	no observable relation to granitoid intrusion	crystalline rocks of amphibolite facies (paragneiss, mica schist)	Saxothuringicum	Variscan
Mutice u Mladé Vožice	W	Cpx+Plg+Sch	no observable relation to granitoid intrusion	paragneiss and migmatites	Moldanubicum	pre-Variscan
Nekvasovy-Chlumy	W	Cpx+Plg+Sch	Central Bohemian Massif - Chanovice apophyse (granodiorite)	skarn in cordierite-biotite migmatites	Moldanubicum	Variscan
Proseč	W	Grs+Di+Qtz±Ttn+Sch	Budislav Pluton in Polička crystalline unit	skarn-marble xenoliths in tonalite to granodiorite	Bohemicum	Variscan
Krouna	W	Grs+Di+Qtz±Ttn+Sch	no observable relation to granitoid intrusion	skarn-like calc-silicate rocks of Svratka crystalline unit	Bohemicum	Variscan
Chedrbí	W	Amp+Kfs+Cpx+Grt+Sch	no observable relation to granitoid intrusion	amphibolite	Kutná Hora-Svratka crystalline unit	pre-Variscan
Bílé Labe	Cu-W	Plg+Qz+Sch	Krkonoše-Jizera Massif	impregnations in metasomatized	Saxothuringicum	Variscan

Tabl. 1 cont.

				biotite granite and quartz veins		
Reduced skarns						
Budislav u Litomyšle	W	Cpx+Plg+Wol+Sch	quartz diorite-tonalite-granodiorite Budislav Massif	xenoliths of skarnized marbles and skarn in tonalite	Bohemicum	Variscan
Kotel (Krkonose)	Cu	Grt+Cpx+Act+Ep+Sch	Krkonoše-Jizera Massif	dolomite marble in contact zone of porphyritic biotite granite	Saxothuringicum	Variscan
Hostákov u Třebíče	W	Cpx+Grt+Ves+Flr+Qz+Sch	Třebíč porphyritic amphibole-biotite melasyenite (durbachite) Massif	xenoliths of biotite paragneiss in amphibole-biotite melasyenite (durbachite)	Moldanubicum	Variscan
Hazlov u Chebu	W	Grt+Ves+Wol+Plg+Ep+Sch	Smrčiny granite Massif	skarn xenoliths in granite	Saxothuringicum	Variscan
Vrbík u Horažďovic	W	Grt+Cpx+Scp+Sch	Central Bohemian Massif - amphibole-biotitic granodiorite	skarnized dolomite xenolith in granodiorite	Moldanubicum	Variscan
Reduced Intrusion Related Gold System (RIRGS)						
Vacíkov-Petráčkova hora	Au	Qz+Po+Ccp+Py+Sch+Mol	Central Bohemian Massif	sheeted arrays of quartz veins and veinlets hosted by the Petráčkova hora granodiorite stock		342.1–348.5 Ma
Čelina	Au - W	Cpx-Amp-Bt-Sch (We studied scheelite sample SCH-3) from bedding parallel quartz vein associated with calc-silicate hornfels in amphibolite. It forms small anhedral grains (0.1–0.7 mm) mostly enclosed in amphibole and locally in association with plagioclase.)	Central Bohemian Massif – tonalite	amphibolite		345–358 Ma
Orogenic Gold Deposits						
Kašperské Hory	Au-W	Qz+Sulp+Sch (We studied two types of scheelite: 1. scheelite from thin quartz vein crosscutting finely banded calc-silicate hornfels (rocks) (SCH5) which was sampled in the Naděje gallery at the Kašperské Hory deposit. Scheelite forms up to 0.8 mm subhedral grains in white-grey quartz matrix and is accompanied by arsenopyrite (up to 0.9 mm. 2. – Scheelite (SCH-6) is from the heavy mineral concentrate from the Lohbach brook, which contains abundant ilmenite, garnet, lesser magnetite, rutile, sillimanite and clinozoisite and rare monazite, spinel, pyrite, arsenopyrite, apatite and native gold. Scheelite (up to 3 mm in size) is intergrown with quartz, biotite, sericite, sillimanite (with inclusions of Ti minerals and apatite), garnet, pyrite and limonite. Locally, scheelite contains inclusions of apatite and pyrrhotite. The mineral assemblage of the heavy mineral concentrate indicates several	no observable relation to granitoid intrusion	biotite, sillimanite-biotite, ± cordierite, muskovite, garnet paragneiss	Moldanubicum	344 Ma (Au veins)

Tabl. 1 cont.

		sources: (a) sillimanite-biotite gneiss, (b) quartz veins with gold and (c) skarn.				
Sobětice	Au-Ag	Qz+Au+Sulph+Sch	exocontact of Klatovy granodirite apophyse of the Central Bohemian Massif	coincident quartz lodes in marble at the contact with paragneiss	Moldanubi-cum	Variscan
Jílové u Prahy	Au	Qz+ Cal+Au+Sulph+Sch (We studied a. 0.5 cm orange scheelite grain (SCH-41 – rather homogeneously looking and showing low yellow luminescence) in quartz vein (several cm thick) coming from the third level of the Pepř mine.			Bohemicum	339.0 ±1.5
Orlík u Humpolce	Au	Qz+Plg+Cpx+Amp+Sulph+Sch+Au	no observable relation to granitoid intrusion	quartz veins in paragneiss with calc-silica rocks and quartzites associated with migmatites	Moldanubi-cum	pre-Variscan
Deposit	Deposit size		Temperature and salinity of the fluids		Reference	
Quartz vein/greisen deposits						
Cínovec/ Zinnwald	564 Mt 0.057% Sn, 0.016 % W*, 0.2% Li – total geological resources		384–439°C (topaz); 2.4–9.5wt.% NaCl eq.		Seifert et al. (2011), Seifert and Pavlova (2016), Ackerman et al. (2017), Breiter et al. (2017a, b), Zhang et al. (2017), Starý (2019)	
Krásno-Huber	13.8 Mt 0.224% Sn, 0.032%W* – indicated resources		low-greisenization related to low-salinity (0–7 wt.% NaCl eq.) aqueous fluids with low amount of CO ₂ , CH ₄ , and N ₂ (≤10 mol% in total) at temperatures of ~350–400°C and pressures of 300–530 bar (quartz, topaz, cassiterite)		Suček et al. (1995), Dolníček et al. (2012), Ackerman et al. (2016)	
Cetoraz	83 kt 0.61%W – inferred resources		P-T conditions of 720 ±75 °C at 0.65 ±0.29 GPa calculated from mineral chemistry of muscovite, biotite, quartz, K-feldspar, plagioclase, sillimanite and garnet by THERMOCALC 3.33 software for orthogneiss host rock		Němec and Tenčík (1976), Šponar et al. (1985), Jurák et al. (1986), Losertová et al. (2013), Buřivalová and Losos (2018), Buriánek et al. (2020)	
Ovesná Lhota	occurrence		not available data		Tenčík (1970), Losertová et al (2014)	
Horní Babákov	occurrence		not available data		Hájek (1971)	
Oxidized skarns						
Obří Důl	0.1 Mt 0.43–0.47% W, 0.8 Mt 0.41–0.43 % Cu, 0.19–0.49% Sn – indicated resources		clinopyroxene showed homogenization temperatures (Th) in the range of 387–424°C and only secondary fluid inclusions (Th 228–262°C, salinity up to 8 wt.% NaCl equiv.) were found in garnet (grossular)		Ackerman et al. (2017), Veselovský et al. (2018)	
Kovářská	2.8 Mt, 0.25% W – potencial resources		not available data		Roos and Veselovský (1985)	
Vykmanov u Kovářské	indication		not available data			
Muťice u Mladé Vožice	showing		not available data		Malec (1985)	
Nekvasovy-Chlumy	4 kt 1.67% W – inferred resources		not available data		Hujsl (1988)	
Proseč	occurrence		not available data		Buriánek (2009)	
Krouna	occurrence		500–600°C, 6–15 kbar, based on Fe ²⁺ /Mg ion exchange in garnet/clinopyroxen system		Bubal (2013)	
Chedrbí	occurrence		not available data		Malec (1985)	

Tabl. 1 cont.

Bílá Labe	occurrence	not available data	Tuček (1970)
Reduced skarns			
Budislav u Litomyšle	occurrence	550–520°C	Buriánek (2009)
Kotel (Krkonoše)	occurrence		Žáček (2008)
Hostákov u Třebíče	14.8 kt 0.46% W	not available data	Holub and Páša (2004)
Hazlov u Chebu	700 kt 0.3% W - inferred resources	not available data	Štemprok (1985)
Vrbík u Horažďovic	4.8 kt 3.6 % W - mined out	not available data	Hujsl et al. (1987)
Reduced Intrusion Related Gold System (RIRGS)			
Vacíkov - Petráčková hora	30.7 Mt 1.08 g/t – inferred resources	Polyphase fluid inclusions hosted by igneous quartz of the Petráčková hora granodiorite indicate minimum trapping conditions of about 3 kbar and 550°C. The gold-rich Q1 to Q4 veins may have formed along a quasi-isobaric cooling path at 2.5 to 1.5 kbar and 590 to 400°C. The salinity of primary aqueous inclusions falls into ranges 6 to 23 and 33 to 41 equiv. wt% NaCl; in igneous quartz, populations in salinity were observed between 5 to 16, 35 to 40 and 62 to 70 equiv. wt% NaCl.	Studničný (1994), Zachariáš et al. (2001)
Čelina	7.9 Mt 1.58 g/t Au, 0.285 %W – indicated resources	Corresponding fluids belong to the C-H-O-N system and probably resulted from fluid–rock interactions within the metamorphic series at high P-T conditions (T ≈ 450–550°C and P ≈ 250–400 Mpa). Maximum (550°C calculated from Al, Ca and Na contents in plagioclase and amphibole, minimum based on coexisting biotite-chlorite boundary	Malec (1985), Kovář and Mandík (1993), Boiron et al. (2001), Ackerman et al. (2017)
Orogenic Gold Deposits			
Kašperské Hory	11.7 Mt 4.71 g/t Au; 3.1 Mt 1.32% W – indicated resources	The corresponding fluids of Stages I and II belong to the C-H-O-N system. The P-T conditions correspond to 590–520°C and to 0.44–0.25 GPa for Stage I and 480–390°C and ~0.2 GPa for Stage II; calculated from isotopic equilibrium temperatures for Qz-Ms coexisting pair (stage I) respectively for Qz-Sch coexisting pair (stage II);	Punčochář (1992), Strnad et al. (2012), Ackerman et al. (2019)
Soběčice	old workings	low salinity (1–2.2 wt%NaCl eq.) H ₂ O-CO ₂ fluids in Q1 (ore bearing), estimated T ~350°C, P ~ 400 Mpa; based on H ₂ O-CO ₂ isochores calculated from fluid inclusions in quartz	Černý and Litochleb (1983), Hemalová (2010)
Jílové u Prahy	3.3 Mt 3.0 g/t Au – geological resources	216–238°C 3.5 wt.% eq. NaCl; scheelite originated from low saline H ₂ O type fluid at stable T ~ 298°C, characterized by a more evolved Sr isotopic composition (87Sr/ 86 Sr = 0.7087–0.7105; associated with scheelite).	Zachariáš et al. (2013)
Orlík u Humpolce	old workings	360–380°C, 0.6–0.8 kbar calculated from primary inclusions in ore-bearing quartz and in scheelite	Đurišová et al. (1992), Buřival and Losertová (2015)

* – wolframite is the W ore; ** – mineral abbreviations after Whitney and Evans (2010)

THE KAŠPERSKÉ HORY DEPOSIT

Exploration in the 1980s, and until 1999, resulted in estimation of a total resource of 50 t of Au (Punčochář, 1992). The Kašperské Hory gold deposit is located close to the tectonic boundary between the monotonous and varied units of a regional shear zone of Moldanubicum (Fig. 1). The predominant host rock is variably migmatized biotite or biotite–sillimanite paragneiss with minor layers of quartzite, quartz–feldspar gneiss, calc-silicate rocks, amphibolite and marble, accompanied by two bodies of orthogneiss. Variscan magmatic rocks form small intrusions of granitic lithologies belonging to the Moldanubian Batholith (MB) and the Central Bohemian Plutonic Complex (CBPC) accompanied by granodiorite porphyry and small pegmatite bodies. Four stages of mineralization were identified (Strnad et al., 2012). Stages I–II are represented by three different quartz vein generations formed at ~390–590°C, containing variable amounts of arsenopyrite, scheelite, pyrite, and invisible gold hosted mainly by arsenopyrite. Economic gold mineralization is bound to stage III, characterized by the formation of quartz, galena, native gold and Bi–Te minerals at lower T conditions (180–290°C; Strnad et al., 2012). The gold mineralization forms stratabound bodies bound by layers of calc-silicate rocks (Morávek and Punčochář, 1983) or disseminations related to stage II mineralization. The gold is typically associated with pyrite, pyrrhotite, arsenopyrite, löllingite, molybdenite, Bi–Te minerals, chlorite, scapolite and albite (Pertold and Punčochář, 1996).

Ackerman et al. (2019) concluded, on the basis of arsenopyrite and molybdenite Re/Os (~341 and 333 Ma, respectively) and Ar–Ar geochronological data on silicate minerals (~332 to 325 Ma), that this Au deposit is younger than the gold deposits spatially associated with the CBPC (~349–339 Ma), indicating instead a temporal relationship to the nearby MB.

The scheelite studied comes from two locations:

(1) Scheelite from a thin quartz vein cross-cutting finely banded calc-silicate hornfels (SCH5) was sampled in the Naděje gallery in the Kašperské Hory deposit. Scheelite forms subhedral grains, up to 0.8 mm across, in white-grey quartz matrix and is accompanied by arsenopyrite (up to 0.9 mm). Malec (1985) concluded that this scheelite was remobilized from an earlier scheelite phase disseminated in a stratiform layer in calc-silicate hornfels, during the formation of Au–W mineralization.

(2) Sample SCH-6 is from a heavy mineral concentrate from the Lohbach brook, which contains abundant ilmenite, garnet, lesser magnetite, rutile, sillimanite, clinozoisite, rare monazite, spinel, pyrite, arsenopyrite, apatite and native gold. Scheelite (up to 3 mm in size) is intergrown with quartz, biotite, sericite, sillimanite (with inclusions of Ti minerals and apatite), garnet, pyrite and limonite. Locally, scheelite contains inclusions of apatite and pyrrhotite. The mineral assemblage of the heavy mineral concentrate indicates several sources: (a) sillimanite-biotite gneiss, (b) quartz veins with gold and (c) skarn.

REDUCED INTRUSION-RELATED GOLD SYSTEMS (RIRGS)

Scheelite is also a common mineral in RIRGS of the Bohemian Massif (e.g., Mokrsko and Čelina deposit, Vacíkov-Petráčkova Hora deposit – see Table 1 for more information; Morávek, 1996).

The Mokrsko and Čelina deposits show low-grade (1.5–2.5 g/t Au) mineralization and share features of both orogenic and RIRGS (Zachariáš et al., 2014; Fig. 1). They belong to the Psí Hory gold-bearing district with a long history of gold mining, since Celtic times (the 3rd to 1st centuries B.C.).

The deposits are mainly hosted by tonalite of the calc-alkaline Sázava tonalite suite (~354 Ma, Janoušek et al., 2004) of the CBPC. Gold dominantly occurs in quartz veins (<1 mm to 1 m thick) and stockworks with up to 1% arsenopyrite at Mokrsko, whereas pyrite and pyrrhotite are dominant at Čelina (Morávek 1995a), with relatively abundant scheelite accompanied by native bismuth, native gold, maldonite [Au₂Bi], aurostibite [AuSb₂], joséite-B [Bi₄Te₂S], pilsenite [Bi₄Te₃], tetradymite [Bi₂Te₂S], Bi₂Te, bismuthinite [Bi₂S₃] and other phases (Zachariáš et al., 2014). Molybdenite has yielded Re–Os age values of 342.9 ± 1.4 Ma (Zachariáš and Stein, 2001) and 342.6 ± 2.2 Ma (Ackerman et al., 2017), indicating close relationship of gold deposition with the end of intrusive activity (345–358 Ma, Janoušek et al., 2004, 2010). Chalcopyrite, galena, sphalerite, löllingite, cubanite, tetrahedrite and ullmanite occur as accessory minerals. With total gold resources of 100–140 tonnes, the Psí Hory district belongs among the most important gold districts in Europe (Morávek, 1996).

The scheelite sample (SCH-3) is from a bedding-parallel quartz vein associated with calc-silicate hornfels in amphibolite. It forms small anhedral grains (0.1–0.7 mm) mostly enclosed in amphibole and locally in association with plagioclase.

SKARN

We follow the classification of skarn deposits into oxidized and reduced on the basis of the mineral assemblage and scheelite composition after Miranda et al. (2022). Oxidized skarns are closely associated with magnetite-series intrusions emplaced at shallow crustal levels (<2 kbar; Newberry, 1983; Newberry and Swanson, 1986; Chang et al., 2019) and developed in impure carbonates and/or hematite-rich metasedimentary host rocks (Newberry, 1983; Newberry and Swanson, 1986). In general, reduced skarns are hosted in pure limestone and are associated with ilmenite-series intrusions (Ishihara, 1977), emplaced at greater depth (>2 kbar; Newberry and Swanson, 1986; Chang et al., 2019). Reduced skarns are characterized predominantly by clinopyroxene, amphibole and sulphides with minor garnet, whereas oxidized skarns consist mainly of garnet, clinopyroxene, amphibole and oxides.

OXIDIZED SKARN

We provide new data on scheelite from W occurrences (e.g., Mutice, Krouna, Nekvasovy Chlumy, Chedrbí, Proseč, and Bílé Labe), see more detailed description in Table 1.

The Obří důl polymetallic deposit is located in the Krkonoše mountains of North Bohemia and has been classified as a polymetallic skarn (Fe–Cu–Zn–As) in the exocontact of the Variscan Krkonoše–Jizera pluton (Veselovský et al., 2018; Fig. 1) with Re–Os ages obtained from molybdenite (312 ± 1.4 Ma, Ackerman et al., 2017) and arsenopyrite (307.5 ± 5.6 Ma, Veselovský et al., 2018). The deposit was mined in the past for As and Cu, with resources estimated at ~100,000 t at 0.43–0.47% W and 856,000 t at 0.41–0.43% Cu and 0.19–0.49% Sn. The orebody consists of two lenses, their thicknesses varying from 1 to 12 m. The deposit is formed by pyroxene-garnet skarn hosting polymetallic mineralization represented by irregular clusters, impregnations, and rare veinlets of pyrrhotite, chalcopyrite and arsenopyrite with minor pyrite, sphalerite, molybdenite, cassiterite, stannite, scheelite, malayaite, bismuthinite and other ore minerals (Šrein and Šreinová, 2000; Pašava et al., 2016). Ore is localized in pyroxene-garnet skarn and altered dolomite. Scheelite is present in two associations: (a) with prehnite, pumpellyite and calcite in younger quartz veins (alpine-type) cross-cutting scheelite-bearing con-

tact skarn and (b) in association with ore minerals in contact metamorphosed skarn (Malec, 1985). Most of the scheelite samples came from contact-metamorphosed skarn and were described by Miranda et al. (2024).

REDUCED SKARN

Typical examples of reduced skarn deposits with scheelite in the Bohemian Massif are the Vrbík and Kotel deposits for which scheelite compositions were reported by Miranda et al. (2024). We provide new data for scheelite from Budislav (see Table 1 for more detailed description).

Vrbík deposit at Horažďovice. This small scheelite deposit is characterized by a lense-shaped diopside skarn (~50 × 19 m with a thickness of ~1.8 m) which is in fact a xenolith sitting in the Variscan granitoids of the Central Bohemian Plutonic Complex (CBPC; Hujsl et al., 1987; Fig. 1). Euhedral whitish, greyish or slightly yellowish grains are up to 2 cm in size and, locally, scheelite concentrations reached ~40% (Hujsl et al., 1987).

The Kotel polymetallic deposit (Krkonosé Mts.) is a polymetallic contact skarn, belonging to the Krkonosé-Jizera metamorphic complex, close to the contact with the Variscan Krkonosé-Jizera Granite Pluton. The skarn was briefly exploited in the 19th century (Malásek, 1957). Šrein et al. (1998) reported Cu 0.64 wt.%, Pb 0.032 wt.%, Zn 0.037 wt.% in average ore samples. Greyish-white scheelite grains (up to 4 mm) with white to bluish white CL zones are dispersed in the rock matrix, not intergrown with sulphides. The mineral paragenesis is represented by calcite, quartz, plagioclase, grossular, diopside, vesuvian, epidote, actinolite, chlorite, K-feldspar, titanite, fluorite, fluorapatite and zircon. The ore minerals are pyrite, chalcopyrite, arsenopyrite, sphalerite, minor scheelite and Ag-Cu-Fe sulphide (Šrein et al., 1998; Žáček, 2008).

OTHER SETTINGS

We studied scheelite in quartz veins hosted in Neoproterozoic volcano-sedimentary rocks at Erpužice u Stříbra and Tehov u Říčan.

Erpužice u Stříbra

Scheelite (SCH-14) was taken from an abandoned quarry of Neoproterozoic amphibolite schist (metabasalt), part of the Teplá Barrandian Unit. Scheelite forms rare fine-grained impregnations in a limonitized quartz-veinlet (Malec, 1985).

Tehov u Říčan

Scheelite (SCH-36) is associated with arsenopyrite and cassiterite in quartz-feldspar-tourmaline veinlets and tourmalinized parts of contact metamorphosed schists/hornfels at the contact of the Říčany Variscan granite (Bernard et al., 1981). The scheelite is commonly intergrown with plagioclase but rarely replaces plagioclase and quartz (Malec, 1985).

SAMPLES AND METHODS

Twenty-six polished sections from 26 sites representing five major association types from the Bohemian Massif (greisen-, orogenic gold and RIRGS-, oxidized and reduced skarn deposits, Table 1) were studied. Major and minor element composition ($n = 368$ analysis, including those partly published by Miranda et al., 2024) and cathodoluminescence (CL) images of scheelite were obtained by electron probe micro-analyzer (EPMA) at Université Laval by use of a CAMECA SX-100 instrument equipped with five wavelength-dispersive spectrometers (WDS) and a CL detector (see Appendix 3). The analytical conditions for CL images were using an accelerating voltage set at 15 kV and the beam current at 20 nA with a working distance

of 2 mm. The analyses were performed using a 10- μ m-diameter beam. Analytical conditions for major element analyses (CaO and WO_3) were 20 nA for beam current and 15 kV for accelerating voltage, with a counting time of 20 s at the peak and 10 s in the background. For minor elements (Na, Sr, Mo, Y and Fe), the analytical conditions were 100 nA beam current, 15 kV of accelerating voltage, and counting times of 120 s at the peak and 30 s in the background. Metal tungsten (W), metal molybdenum (Mo), diopside (Ca), celestite (Sr), YPO_4 (Y), magnetite (Fe) and albite (Na) were used as standards. For CL investigation, the accelerating voltage was set at 15 kV, the beam current at 20 nA, with a working distance of 2 mm. The dwell time of each pixel was 0.03 ms. The CL images were recorded in grey-scale in order to document texture and zonation patterns in the scheelite prior to micro-analyses (EPMA and LA-ICP-MS).

The concentration and distribution of the minor and trace elements in the scheelite ($n = 368$ analyses) were determined by LA-ICP-MS at the LabMaTer, Université du Québec à Chicoutimi (UQAC) by use of an Excimer 193 nm RESOLUTION M-50 laser ablation system (Australian Scientific Instrument) equipped with a double volume cell S-155 (Laurin Technic) and coupled with an Agilent 7900 mass spectrometer. The LA-ICP-MS tuning parameters were a laser frequency of 15 Hz, a fluence of 3 J/cm², and scanning speed of 10 μ m/s for the line scans. Spots across the surface of scheelite grains were made with beam sizes of 33 and 55 μ m, depending on grain size. Given the distinct textures in scheelite, the spots were performed in all distinct CL zones to obtain as much variance as possible in the scheelite composition. Data reduction was carried out by use of Lolite v3 running in Igor Pro 6.37 (Paton et al., 2011). ⁴⁴Ca was used for internal standardization for scheelite on the basis of EPMA results (Appendix 1). Synthetic glass reference material NIST-610 was used as an external standard for all elements by use of preferred values from the GeoReM database (Jochum et al., 2005). The basalt glasses GSE-1g, GSD-1g and Gprobe6-A and the NIST-612 (synthetic glass) were used as secondary reference materials for quality control. The reference materials were analyzed at the beginning, throughout and the end of each analytical session, to monitor potential instrumental drift. The results obtained for the monitors allowed for standard deviations on the working values (Appendix 4). The following isotopes were measured: ²³Na, ²⁴Mg, ³⁹K, ⁴⁴Ca, ⁴⁹Ti, ⁵¹V, ⁵⁵Mn, ⁵⁷Fe, ⁶¹Ni, ⁶³Cu, ⁶⁶Zn, ⁷⁵As, ⁸⁸Sr, ⁸⁹Y, ⁹³Nb, ⁹⁵Mo, ¹⁰⁷Ag, ¹¹⁸Sn, ¹³⁷Ba, ¹³⁹La, ¹⁴⁰Ce, ¹⁴¹Pr, ¹⁴⁶Nd, ¹⁴⁷Sm, ¹⁵³Eu, ¹⁵⁷Gd, ¹⁵⁹Tb, ¹⁶³Dy, ¹⁶⁵Ho, ¹⁶⁶Er, ¹⁶⁹Tm, ¹⁷²Yb, ¹⁷⁵Lu, ¹⁸¹Ta, ¹⁸²W, ¹⁸³W, ²⁰⁸Pb, ²⁰⁹Bi, ²³²Th and ²³⁸U. Silicon and S were also monitored to identify possible silicate and sulphide inclusions, respectively. Analyses affected by inclusions were discarded.

Pearson and Spearman correlation coefficients and their significance at levels of confidence $p = 0.001$; 0.01 and 0.05 corresponding to 99.9, 99 and 95% confidence, respectively, were calculated (Holčín and Komenda, 2015) and used for the evaluation of monotonic relationships between the elements analysed in scheelite from different types of deposit/mineralization.

RESULTS

CATHODOLUMINESCENCE AND TRACE ELEMENT VARIATIONS

GREISEN-RELATED SCHEELITE

Scheelite from greisen-type deposits shows a homogeneous texture (Fig. 2A) or two generations (Fig. 2B). We studied two types of scheelite from Cínovec. Scheelite from mas-

sive greisen (Cínovec 1, SCH-47) is heterogeneous, showing brighter CL at grain edges, while the darker central part is more homogenous (Fig. 2C). Reddish-brown grains of scheelite (Cínovec 2, SCH-48) from vein-type greisen are more homogeneous under CL. Under UV light, scheelite from Sn-W deposits displays variable colours from whitish (H. Babákov) via yellowish (Cínovec) to bluish (H. Slavkov-Huber).

Chemical composition

Scheelite from greisens shows variable chemical composition (Appendices 1 and 2 and Table 2) and REE patterns (Fig. 3). Scheelite from Horní Babákov (SCH-37) shows maximum median values of Y (465 ppm), LREE (2295 ppm) and HREE (633 ppm) while scheelite from Krásno-Huber (SCH-43) displays the lowest median concentrations of As (bdl), Y (0.1 ppm), Nb (7.8 ppm), Mo (0.4 ppm), LREE (1 ppm) and HREE (0.1 ppm) of all the scheelite studied from the greisen-type association. The highest median values of Sr (567 ppm), Nb (244 ppm) in the greisen group were found in scheelite from Cětovaz. Scheelite from Ovesná Lhota has median Sr (240 ppm), Nb (30 ppm), Y (357 ppm) and Mo (1.7 ppm) values close to those of Horní Babákov whereas LREE (739 ppm) and HREE (143 ppm) are lower.

Scheelite from massive greisen (Cínovec 1, SCH-47) at Cínovec shows the highest median concentrations of Zn (59 ppm), As (11370 ppm), Ag (9.3 ppm), Sn (8 ppm), Pb (3704 ppm), Bi (940 ppm) and Th (223 ppm) of all greisen-associated scheelite (Appendix 1). Scheelite from vein greisen at Cínovec (Cínovec 2, SCH-48) displays high median values of Mo (7478 ppm) and U (2996 ppm) which are comparable with those of scheelite from the massive greisen at Cínovec.

REE patterns

Scheelite from greisens displays the following REE patterns (Fig. 3): (1) LREE-enriched with a positive Eu anomaly (Cětovaz) (2) HREE-enriched with negative Eu and Ce anomalies (Cínovec 2, SCH-48), (3) Bell-shaped REE patterns with a negative Eu anomaly (Cínovec 1, SCH-47), (4) Flat and bell-shaped REE patterns without a Eu anomaly (Ovesná Lhota), (5) Flat REE patterns with a positive Eu anomaly (Ovesná Lhota and Cětovaz).

Scheelite from massive greisen (Cínovec 1, SCH-47) shows LREE-depleted and bell-type patterns, both with a Eu negative anomaly (Fig. 3A). Scheelite with a LREE-depleted pattern shows enrichment in Sb (67-fold), Bi (6-fold), Y (5-fold), Hf (4-fold), REE (3-fold) and Mo (1.5-fold). Scheelite from vein greisen at Cínovec (Cínovec 2, SCH-48) is dominantly characterized by a HREE-enriched pattern with negative Eu and Ce anomalies (Fig. 3A).

Scheelite from massive greisen at Krásno-Huber shows irregular REE patterns (mostly with a positive Eu anomaly, Fig. 3C), often with the values below detection limits.

Scheelite from Horní Babákov shows bell-shaped REE chondrite-normalized patterns with a Eu positive anomaly (Fig. 3D). In general, scheelite from H. Babákov shows comparable a trace element distribution to scheelite from Krásno-Huber except for enrichment in Y and REE, which is also much higher than in scheelite from other greisen localities.

Scheelite from Cětovaz shows LREE-enriched patterns with a positive or without a Eu anomaly or flat REE patterns showing a positive Eu anomaly (Fig. 3E). It is enriched in Na (like that of H. Babákov), Mn and Sr (the highest median values among all the greisen samples). Median values of Y, LREE and HREE are comparable with those of H. Babákov and significantly higher than those of Cínovec and Krásno-Huber.

Scheelite from Ovesná Lhota shows flat REE patterns with or without a positive Eu anomaly (Fig. 3F).

OROGENIC GOLD SCHEELITE

Scheelite CL from the orogenic gold deposits is predominantly homogeneous (Fig. 2D). Only scheelite from the Jílové deposit shows oscillatory zoning (Fig. 2E). Under UV, scheelite from this setting is dominantly bluish.

Chemical composition

Scheelite from Jílové shows the highest median values of Na (84 ppm), As (15 ppm), Y (466 ppm), LREE (1032 ppm) and HREE (188 ppm) and low Mo median values (23 ppm) when compared to scheelite from other orogenic gold locations in the Bohemian Massif (Appendix 1). Scheelite from Kašperské hory – Naděje and Kašperské hory – Lohbach has peak median Mo values of 113 and 106 ppm, respectively, while scheelite from Orlik is characterized by the highest median Sr value (2118 ppm, Appendix 1).

REE patterns

Scheelite from orogenic gold settings displays bell-shaped REE patterns with a positive Eu anomaly and flat REE patterns with a positive, or small negative, Eu anomaly. REE patterns for the Jílové, Soběčice, Kašperské hory-Naděje, Kašperské hory Lohbach and Orlik u Humpolce deposits are shown on Figure 4A–E.

At Jílové, oscillatory-zoned scheelite has flat and bell-shaped REE patterns in the same scheelite grain (Figs. 2E and 4A). Scheelite with a flat REE pattern and positive Eu anomaly is characterized by higher Ni (20-times), Mg (5-times), Th (3-times) and U (2-times) values, compared to scheelite with a bell-shaped pattern with a positive Eu anomaly, which has significantly higher V (5-times), Na (4-times), REE (3-times), and Y (2-times) concentrations.

At Soběčice, scheelite with a bell-shaped pattern and small negative Eu anomaly is characterized by higher V (4-times), Y (2-times), and REE (2-times) values compared to scheelite with a flat REE pattern and positive Eu anomaly, which shows significantly higher U (30-times) and Sb (4-times) values.

AU RIRGS SCHEELITE

Scheelite from RIRGS is predominantly homogeneous (e.g., Čelina SCH-3; Fig. 2F) under CL and displays creamy white (Vacíkov) to bluish (Čelina) colour under UV light.

Chemical composition

Scheelite from Vacíkov (SCH-1) is characterized by significantly higher median values of Na (113 ppm), Sr (155 ppm), LREE (919 ppm) and similar median concentrations of K, Ti, Mn, As, Nb, Mo and HREE when compared to scheelite from the Čelina deposit (SCH-3) (Appendix 1).

Scheelite from this setting is characterized by the highest median values of Na (54 ppm), K (13.1 ppm), Mg (10 ppm), As (12.3 ppm), V (0.8 ppm) and Cu (1 ppm) of all the groups of mineral deposits studied (Table 2). In comparison with scheelite from orogenic gold deposits, it typically shows significantly higher Nb and Mo median values and lower Sr, LREE and HREE median values (Table 2).

REE patterns

Scheelite from RIRGS displays a bell-shaped REE pattern with a negative Eu anomaly while scheelite from Vacíkov has steep LREE-enriched patterns with a negative Eu anomaly (Fig. 4F, G).

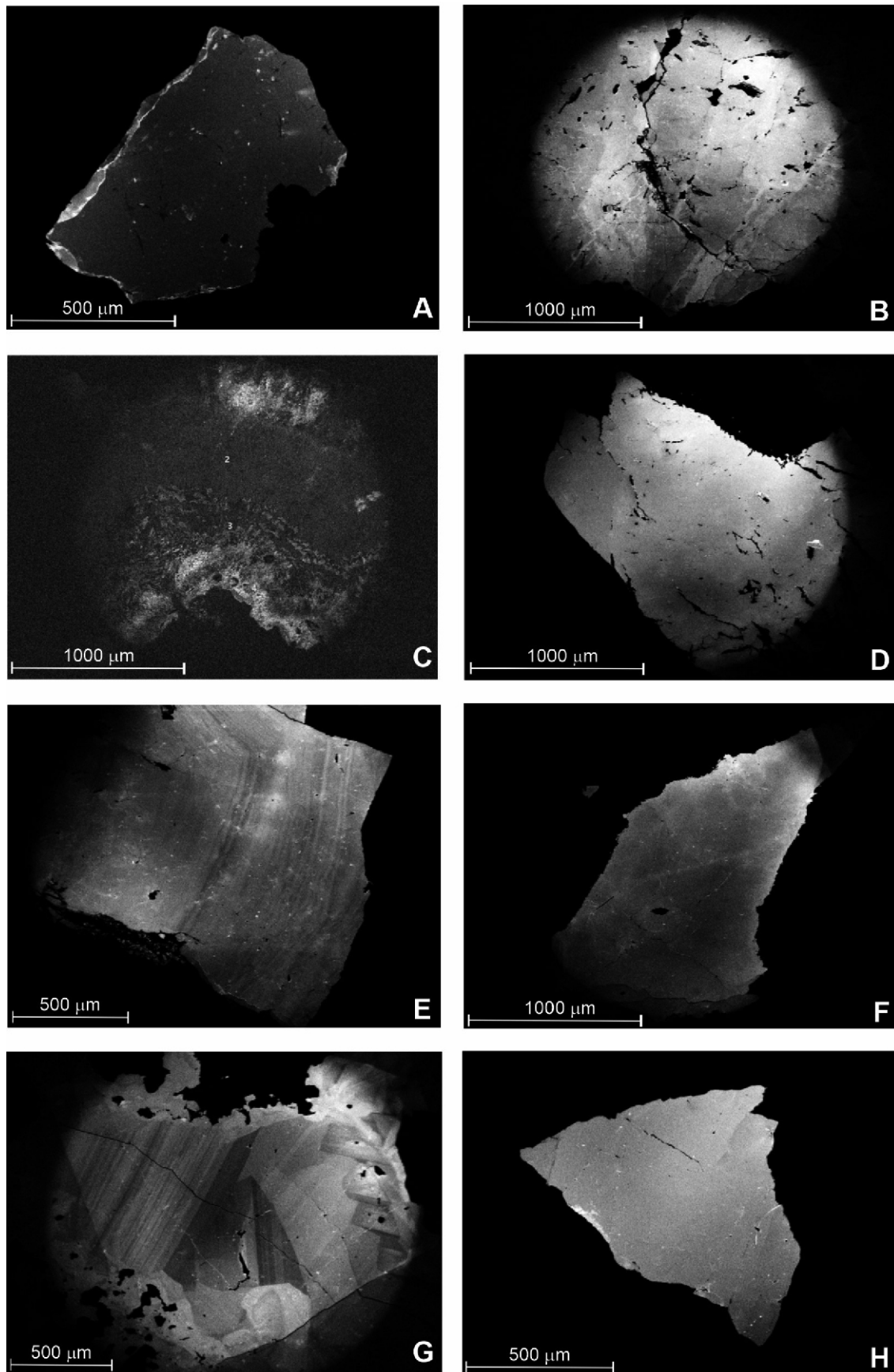


Fig. 2. Cathodoluminescence (CL) images of scheelite grains from different types of mineralization in the Bohemian Massif

A – homogeneous scheelite grain from Krásno-Huber greisen deposit; **B** – oscillatory zoning/two generations of scheelite from a Ovesná Lhota quartz-cassiterite-wolframite vein; **C** – inhomogeneous (brighter) marginal part of grain with more homogeneous (darker) central part of scheelite grain from massive greisen at the Cínovec deposit; **D** – homogeneous scheelite from the Kašperské Hory (Naděje gallery) orogenic Au deposit; **E** – oscillatory zoned scheelite from the Jílové orogenic Au deposit; **F** – homogeneous scheelite from the Čelina – RIRGS Au deposit; **G** – scheelite with oscillatory zoning from oxidized skarn at Mutice; **H** – homogeneous scheelite from reduced skarn at Hostákov

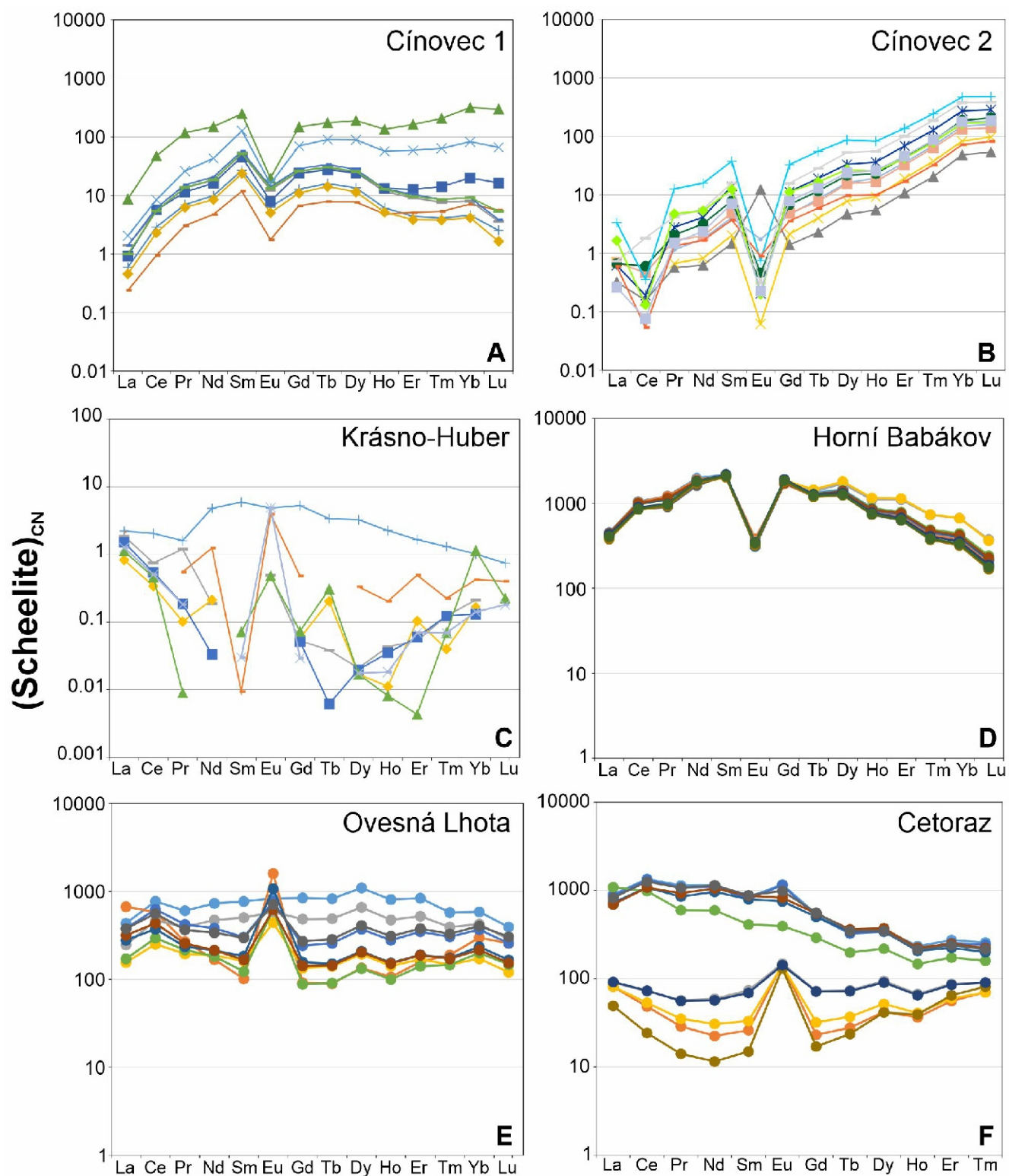


Fig. 3. Rare earth element patterns in scheelite from individual greisen and quartz-cassiterite-wolframite vein localities (Bohemian Massif)

A – Cínovec 1 (massive greisen), **B** – Cínovec 2 (vein-type greisen), **C** – Krásno-Huber deposit, **D** – Horní Babákov deposit (greisen), **E** – Ovesná Lhota (quartz-cassiterite-wolframite vein), **F** – Cetoraz (quartz-wolframite vein)

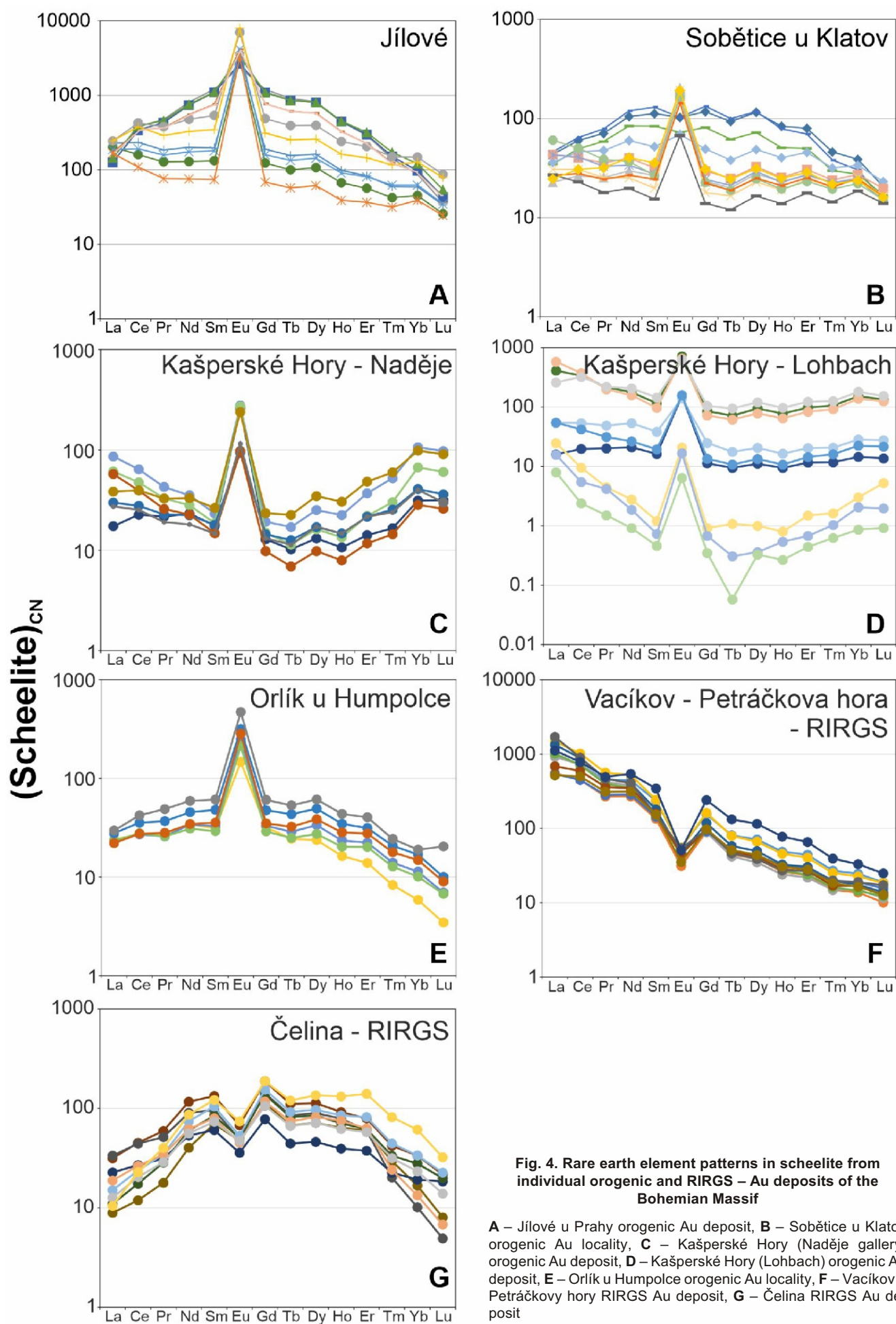


Fig. 4. Rare earth element patterns in scheelite from individual orogenic and RIRGS – Au deposits of the Bohemian Massif

A – Jílové u Prahy orogenic Au deposit, **B** – Soběstice u Klatov orogenic Au locality, **C** – Kašperské Hory (Naděje gallery) orogenic Au deposit, **D** – Kašperské Hory (Lohbach) orogenic Au deposit, **E** – Orlík u Humpolce orogenic Au locality, **F** – Vacíkov u Petrůvkovy hory RIRGS Au deposit, **G** – Čelina RIRGS Au deposit

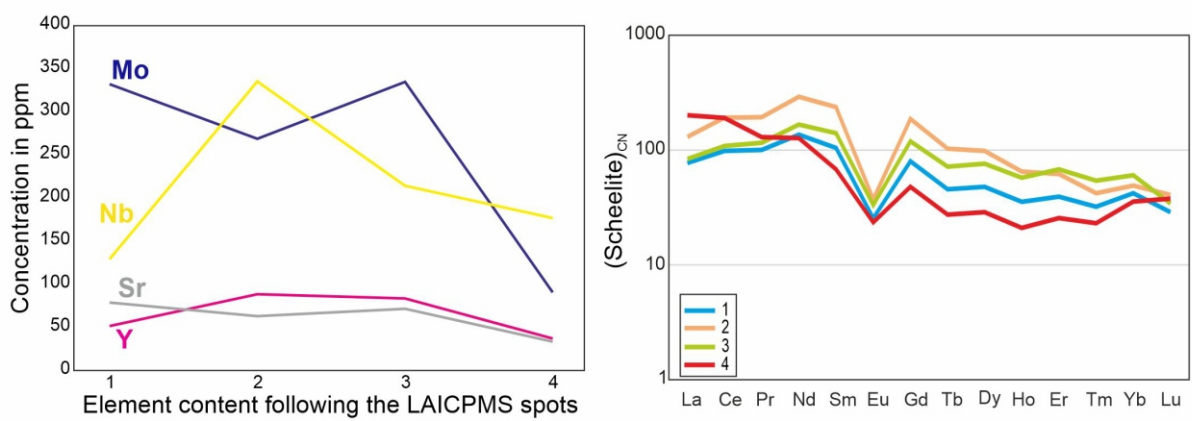
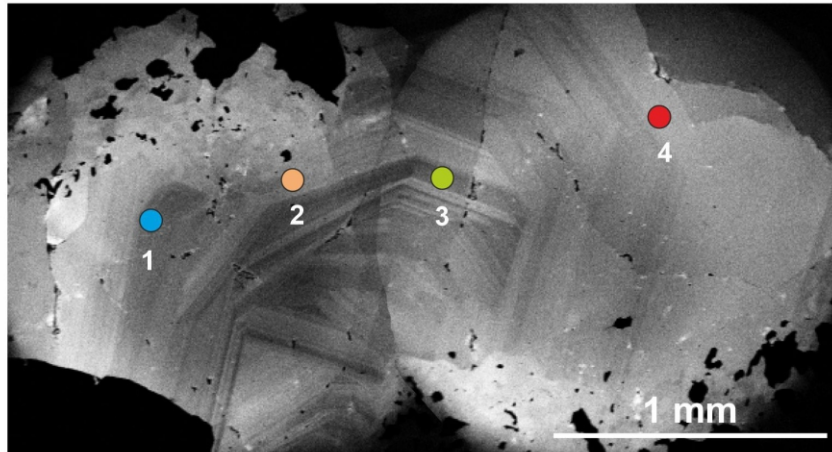
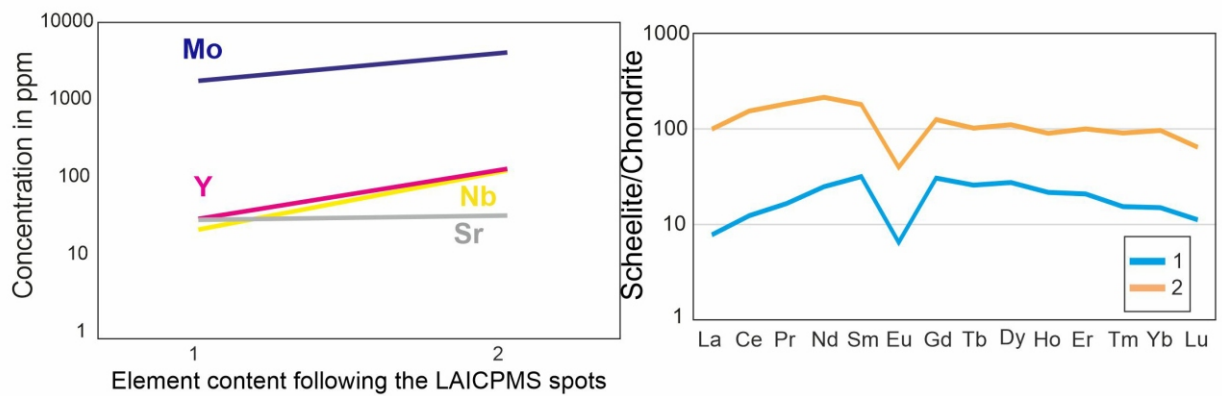
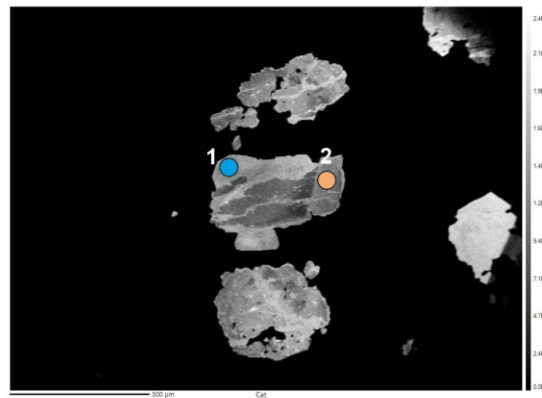
A**B**

Fig. 5. Changes in chemical composition (Mo, Nb, Sr, Y and REE) in scheelite with oscillatory zoning from oxidized skarn at Mutice (A) and Obří důl (B)

anomaly, (2) Bell-shaped REE patterns with a negative or absent Eu anomaly, (3) Flat REE patterns with a negative Eu anomaly (Figs. 6 and 7). REE patterns of scheelite from oxidized skarn deposits (Bílá Labe, Krouna u Poličky, Proseč u Litomyšle, Nekvasovy Chlumy, and Chedrbí) are shown on Figure 7A–E, respectively. REE patterns of scheelite from reduced skarn at Budislav u Litomyšle are shown on Figure 7F.

SCHEELITE FROM THE OTHER SETTINGS

Both scheelite grains studied are homogeneous under CL and show a bluish colour under UV light. They also show very similar distributions of trace and REE contents except for higher Mo and Sr values in scheelite from Tehov u Říčan (Appendix 1). The chondrite-normalized REE patterns are flat with a positive Eu anomaly (Fig. 8). Their chemical signature is similar to that of scheelite from orogenic gold deposits (Fig. 4).

CORRELATION ANALYSIS

Pearson and Spearman correlation coefficients and their significance at levels of confidence $p = 0.001$; 0.01 and 0.05 were calculated for individual groups of scheelite samples analysed. In the following text we list only the highest correlations at the level of confidence 99.9%.

SCHEELITE FROM GRESISEN

Pearson correlation coefficients: **REE** and Na, Mg, Y; **Mo** and U, As, Sn, Hf; **Th** and V, Ag, Bi; **U** and Mo, As, Zn; **Sr** and Nb, Mn; **Sn** and V, As, Mo, Ag; **Bi** and Pb, Hf, V, Th, Sn, Ag, Sb, Ta, As. Molybdenum shows negative correlation with REE, Y and W.

Spearman correlation coefficients: **REE** and Na, Y; **Mo** and U, As, Sn, Hf, Bi, Zn, Cu, Ti; **Th** and Pb, Bi, Ta, Ag, As, Sn, Nb, Cu, Ti, V, Mn; **Sr** and U; **Mn** and LREE, Pb, Nb, Th, Y, Na, Ti.

SCHEELITE FROM OROGENIC GOLD DEPOSITS

Pearson correlation coefficients: **REE** and Y, Hf, Na; **Mo** and Nb; **Th** and Mn, U, Nb; **Sr** and Ta; **As** and Hf, Y, REE, Na, K.

Spearman correlation coefficients: **REE** and Y, Hf, Na, As; **Mo** and Mn, Nb, Th; **Th** and U, Mn, Mo, Ba; **Sr** and As, K, Ta; **Pb** and Y, HREE.

SCHEELITE FROM AU-RIRGS

Pearson correlation coefficients: **LREE** and Na, Sr, Ta, Pb; **HREE** and Y; **Sr** and Pb, LREE; **U** and As, V; **Th** and Sn, Zn.

Spearman correlation coefficients: **LREE** and Ta, Na, Nb, U; **HREE** and Y; **Sr** and Pb, U, Ta.

SCHEELITE FROM OXIDIZED SKARNS

Pearson correlation coefficients: **REE** and Y, Na, Nb, Ta, V, Mn; **Mo** and As.

Spearman correlation coefficients: **REE** and Y, Na, Nb, V, Hf, Ta, Th, U, As, Mn; **Mo** and As, Ti, Na, Ta, Ti, K, Bi, LREE, Pb, Th; **Nb** and LREE, HREE, Y, Hf, Th, Na, V, Ta, U, Sn, As, Mn.

SCHEELITE FROM REDUCED SKARNS

Pearson correlation coefficients: **HREE** and Y; **LREE** and Ta, Nb; **Mo** and Sn; **Sr** and Au, B; Mn and Si.

Spearman correlation coefficients: **HREE** and Pb, W; **LREE** and Y; **Mo** and Y, Ti, Pb; **Y** and LREE, Th, Nb, Mo.

DISCUSSION

DISTRIBUTION OF TRACE ELEMENTS AND TYPES OF REE PATTERNS

Several studies show that the chemical composition of scheelite is strongly influenced by the composition and physicochemical conditions of the hydrothermal fluids from which it precipitates, the composition of the host rock, crystallographic control and the partition coefficients between fluids, scheelite and co-precipitating minerals (Ghaderi et al., 1999; Brugger et al., 2008; Song et al., 2014; Poulin et al., 2018; Sciuba et al., 2020; Miranda et al., 2022; Altenberger et al., 2024; Haupt et al., 2025).

Several authors (Hsu and Galli, 1973; Hsu, 1977; Tyson et al., 1988) showed that Mo^{6+} replaces W^{6+} at moderate to high $f\text{O}_2$ and that Nb^{5+} , Ta^{5+} , As^{5+} and V^{5+} substitute for W^{6+} , whereas divalent cations (Sr^{2+} , Eu^{2+} , Mn^{2+} , Pb^{2+} and Mg^{2+}) are readily incorporated into the Ca-site due their similarities to Ca^{2+} (Ghaderi et al., 1999). Additionally, scheelite may contain significant amounts of REE^{3+} and Y^{3+} that substitute for Ca^{2+} . In hydrothermal fluids, the temperature, pH and oxidation state of the fluids will control the relative abundance of Ce^{4+} and Eu^{2+} (Sverjensky, 1984; Brugger et al., 2008; Pattan et al., 2008).

Miranda et al. (2022) reported that scheelite from magmatic-hydrothermal deposits has higher Mo, Nb, and Mn and lower Sr contents and predominantly negative Eu anomalies when compared to those formed dominantly by metamorphic fluids in orogenic settings. These authors also concluded that Ti, V, Mn, As, Y, Nb, U and Eu anomaly are the most important variables for discrimination of different types of magmatic-hydrothermal deposits.

GREISEN-RELATED SCHEELITE

Miranda et al. (2024) showed that quartz-vein/greisen Sn-W scheelite is characterized by high Th, U, Ta, Y and Pb contents, which differs from oxidized skarns and RIRGS. By comparison with other types of mineral deposit in the Bohemian Massif, scheelite from the greisen association displays high Y, U and Th and the highest median values of Sr, Mn, Bi, Sn, Zn, Hf and HREE. Our data, however show significant differences in the distribution of most of these elements (Mo, Nb, As, Y, U, Th, Mn, Zn, HREE) in scheelite from individual greisens of the Bohemian Massif (Appendix 1). For example, we note low median concentrations of Y and REE at Krásno-Huber and Cínovec, two major greisen-type deposits, and more than three orders of magnitude higher Pb, Bi, As, Th and U in scheelite in greisen from Cínovec (the latter most likely reflecting the presence of different nano-/micro inclusions/impurities). This scheelite is also significantly enriched in Mo, Nb, consistent with the observations of Breiter et al. (2016) and also in Sn. On the other hand, with the exception of scheelite from Cínovec, consistently high Sr at a roughly similar range of As contents and also elevated Sn accompanied by lower Mo values make them distinguishable from most oxidized and reduced skarns and also from RIRGS (Fig. 9A, B). Most outliers with increased Sn values in the group of oxidized skarns are represented by spots measured in scheelite from the Obří důl deposit from which the presence of Sn-bearing phases was described by Šrein and Šreinová (2000). Higher Nb at a much lower range of Sr and at a higher range of Mo values make them distinct from scheelite of the orogenic gold deposits whereas scheelite from oxidized and reduced skarns is typical of higher Sr and lower Mo values (except for the Cínovec deposit with high Mo values; Fig. 9C, D). Interestingly, the primitive chemical composition of rock without Mo, Nb and Ta is typical of Horní Babákov and Ovesná Lhota and yet, the contents of these elements in schee-

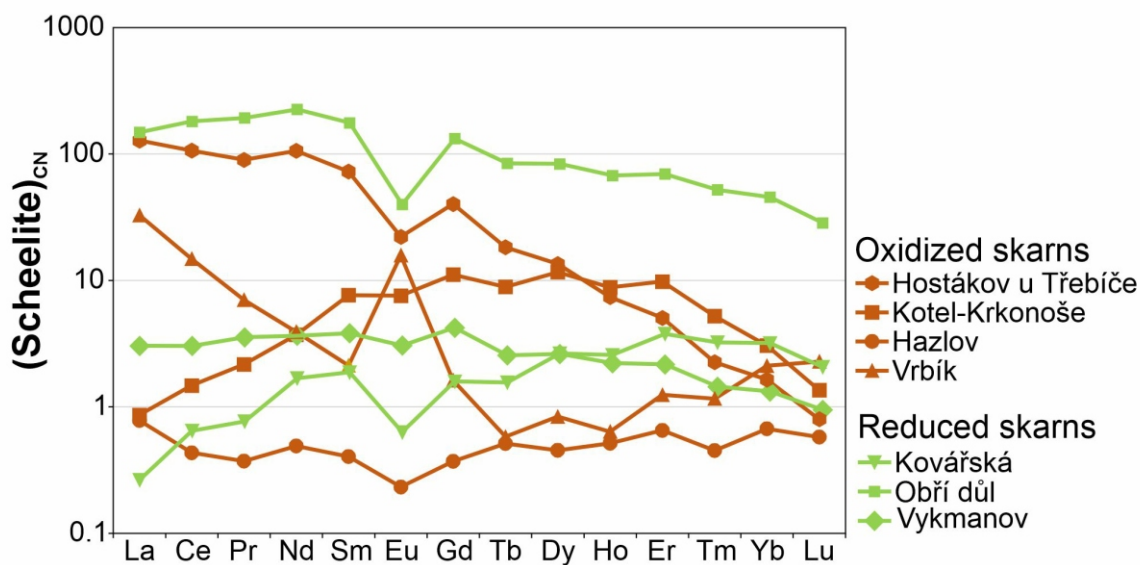


Fig. 6. Rare earth element patterns in scheelite from oxidized and reduced skarns of the Bohemian Massif (on the basis of median values for every locality)

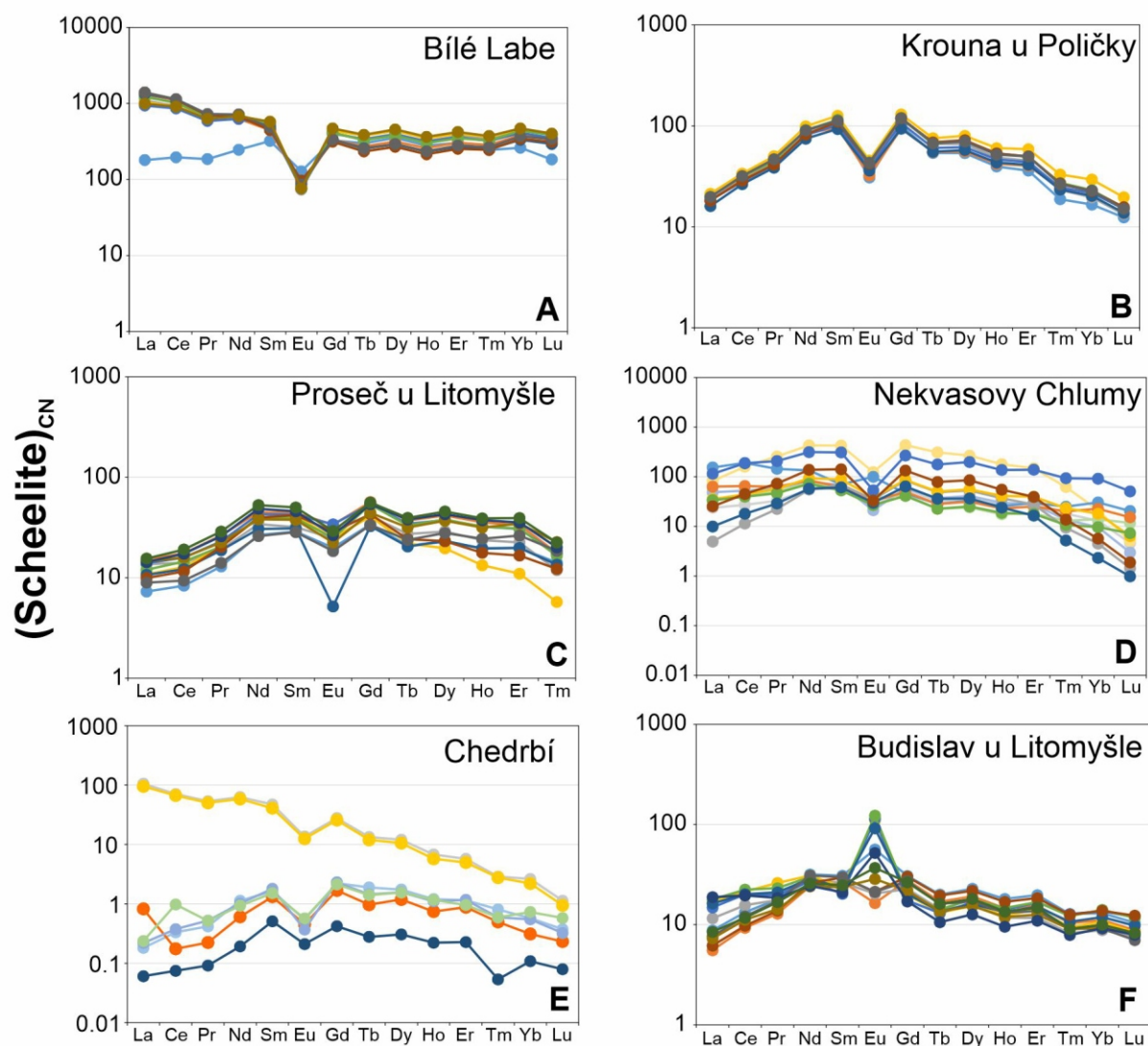


Fig. 7. Rare earth element patterns in scheelite from individual oxidized and reduced skarn deposits/showings of the Bohemian Massif

A – Bílé Labe oxidized skarn, B – Krouna u Poličky oxidized skarn, C – Proseč u Litomyšle oxidized skarn, D – Nekvasovy Chlumy oxidized skarn, E – Chedrbí oxidized skarn, F – Budislav u Litomyšle reduced skarn

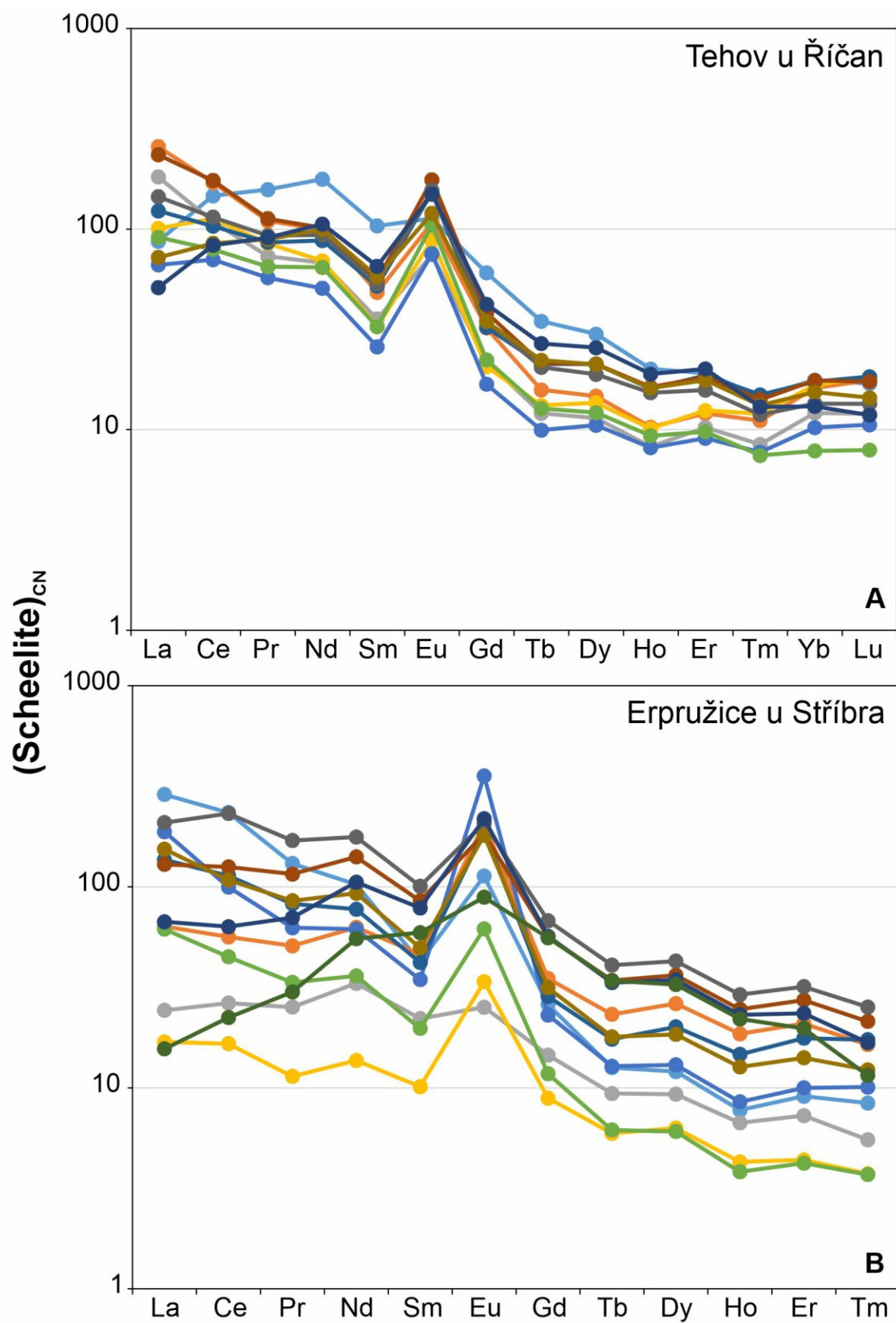


Fig. 8. Rare earth element patterns in scheelite from other individual localities of the Bohemian Massif

A – Tehov u Říčan, B – Erpružice u Stříbra

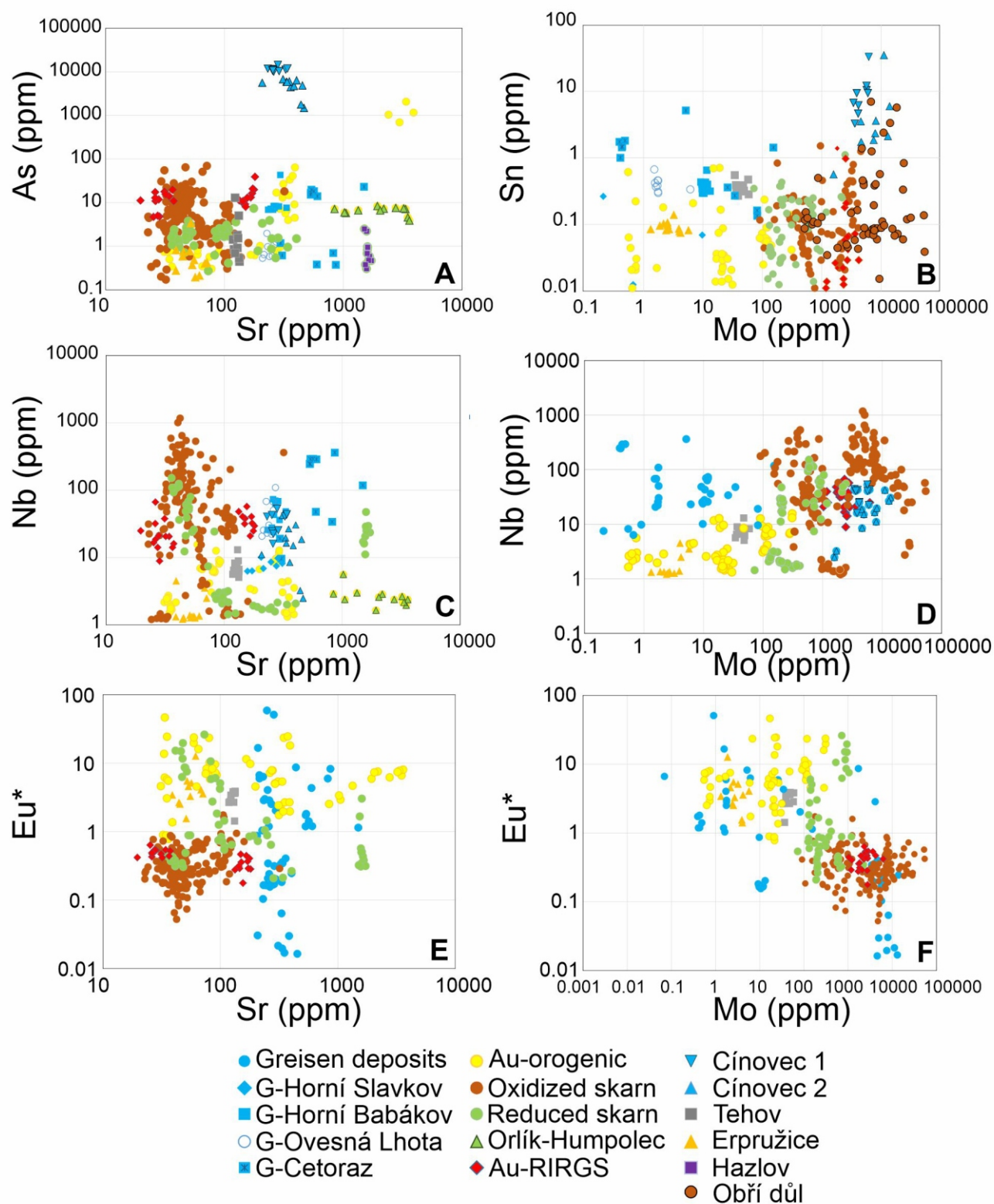


Fig. 9. Binary plots of the minor and trace element composition of the scheelite

A – Sr versus As, B – Mo versus Sn, C – Sr versus Nb, D – Mo versus Nb, E – Sr versus Eu anomaly, F – Mo versus Eu anomaly

lite are similar to those of Mo-rich greisens in the Krušné hory Mts. More information is needed to explain this situation. For example, quartz-cassiterite-wolframite veins at Ovesná Lhota are associated with a small adamellite intrusion which is a part of the Moldanubian Pluton. Morávek (1992) described these veins as only weakly gold-bearing. We can only speculate whether elevated Nb values and a positive Eu anomaly (Fig. 3) could be related to the processes associated with gold enrichment.

Different types of REE pattern and Eu anomaly in scheelite from greisen and related quartz veins of the Bohemian Massif (Fig. 3) reflect different sources of fluids and processes (e.g., Song et al. 2014). The negative Ce anomaly at Cínovec 2 (Fig. 3B) suggests more reduced fluid conditions (Pattan et al., 2005). The shape of the chondrite-normalized REE pattern (typical of a pronounced Eu negative anomaly) from the massive greisen at Cínovec reflects REE distribution in greisen (hostrock) and also other granitic rocks from Cínovec (cf. Breiter et al., 2017a; Fig. 11). This is apparently not the case of scheelite from Krásno-Huber, which is typically of very low REE concentrations and a sharp positive Eu anomaly while the local host rock shows almost two order magnitude higher REE values and a pronounced Eu negative anomaly (Fig. 11). In addition, the shape of the REE pattern in scheelite from Horní Babákov and Cětovaz does not correspond to that of their host rock (Fig. 11) which may indicate that the mineralizing fluids forming that scheelite was not formed by magmatic fluids exsolved from these intrusions. Scheelite from Horní Babákov and Čelina (RIRGS) shows a uniform REE pattern suggesting a homogeneous source of fluids, while scheelite grains from Cětovaz and Ovesná Lhota are characterized by different REE patterns, indicating variable conditions of scheelite formation. Scheelite from the greisen quartz vein at Ovesná Lhota has a similar REE pattern to a group of orogenic Au deposits with a sharp positive Eu anomaly.

A wide range of Eu^* at restricted Sr values characterizes scheelite from the greisen-type association from other groups (Fig. 9E), whereas in the Eu^* vs Mo diagram greisen-type scheelite shows a more scattered picture (Fig. 9F). The correlation between $\text{REE} + \text{Y}$ and Na in scheelite from greisen-type, oxidized skarn and orogenic Au and RIRGS associations (at the deposit-type level; Fig. 10A) indicates a coupled substitution after the following equation: $2\text{Ca}^{2+} = (\text{REE}, \text{Y})^{3+} + \text{Na}^+$ (Nassau, 1963; Nassau and Loiacono, 1963; Cottrant, 1981; Burt, 1989; Rimbault et al., 1993; Brugger et al., 2000a). This type of substitution is considered a major substitution in scheelite (Nassau and Loiacono, 1963; Cottrant, 1981; Burt, 1989; Rimbault et al., 1993; Brugger et al., 2002). However, looking in more detail at the $\text{REE} + \text{Y}$ vs. Na relationship in scheelite from individual deposits from the Bohemian Massif, no positive correlation between these variables at Cínovec and Krásno-Huber (greisen-type), Chedrbí and Vykmanov (oxidized skarn-type) and Soběčice (orogenic Au-type) is visible (Fig. 10B).

OROGENIC AU-RELATED SCHEELITE

On the basis of a study of 25 representative orogenic gold deposits worldwide, Sciuba et al. (2020) reported that scheelite from deposits hosted in low and moderate metamorphic facies rocks commonly have higher REE (median 334 ppm and 632 ppm, respectively) and Y (median 145 ppm and 253 ppm, respectively), Sr, and Na, and low Mn and Nb contents, whereas scheelite from deposits hosted in high metamorphic facies rocks commonly have low REE (median 101 ppm) and Y (median 37 ppm), Sr, and Na, and high Mn and Nb contents. The

chemical composition of scheelite from Jílové (lower metamorphic grade) and of Soběčice, Kašperské Hory and Orlik u Humpolce (high metamorphic grade) is in support of this observation (Appendix 1). The only exception is that scheelite from Orlik u Humpolce shows anomalously high Sr (and low Mn and Nb) content, which likely reflects its close association with the sedimentary (calc-silicate) host rock and scheelite from Soběčice displays low Nb, reflecting its low concentration in the host rock (silicified marble).

The bell-shaped and flat patterns with dominant positive Eu anomaly described from many orogenic Au deposits worldwide (Ghaderi et al., 1999; Brugger et al., 2000b; Roberts et al., 2006; Sciuba et al., 2020) are also typical for the Bohemian orogenic gold scheelite (Fig. 4A–E). REE patterns in scheelite from Jílové, Kašperské Hory-Naděje and Orlik (Fig. 4A, C, E) reflect a homogeneous fluid source, locally with variable redox conditions (e.g., oscillatory zonation of scheelite at Jílové; Fig. 2E; Ďurišová and Sztacho, 1986; Strnad et al., 2012; Zachariáš et al., 2013) or fluctuations in pH from fluid- to rock-buffered conditions, as proposed by Brugger et al. (2008) or even a changing rate of crystallization combined with processes of chemical diffusion (Shore and Fowler, 1996).

Scheelite from Jílové and Kašperské Hory-Naděje, Kašperské Hory-Lohbach and Orlik is typified by a positive correlation between Na and $\text{REE} + \text{Y}$ (Fig. 10B), reflecting most likely coupled substitution. Scheelite from Soběčice shows a negative correlation between Na and $\text{REE} + \text{Y}$ (Fig. 10B) and a slightly positive relationship between Na and Eu^* (Fig. 10C), similar to scheelite from the Nevea Au deposit, Australia (Sciuba et al., 2020), which was ascribed as of both Au skarn (Mueller et al., 2004) and orogenic Au (Goldfarb et al., 2005) type.

Sciuba et al. (2020) reported a negative correlation of Na with Eu^* in scheelite from selected world orogenic Au deposits which is consistent with that of the Jílové, Kašperské Hory-Lohbach and Kašperské Hory-Naděje deposits, while scheelite from Orlik show the opposite trend (Fig. 10C). Sciuba et al. (2020) also noted that the Na content is higher in scheelite from deposits hosted in felsic to intermediate rocks (median 120 ppm and 194 ppm, respectively), and lower in scheelite from deposits hosted in sedimentary rocks (median 53 ppm). Scheelite from our orogenic Au deposits has a much lower median Na value of 21 ppm.

Sciuba et al. (2020) concluded that scheelite from orogenic gold deposits have distinct Sr, Mo, Eu, As, and Sr/Mo, but indistinguishable REE, signatures, compared to scheelite from other deposit types. We show that scheelite from this type of setting in the Bohemian Massif has in general higher REE concentrations when compared to that of reduced skarn (Table 2). Miranda et al. (2024) proposed that orogenic Au deposits and metamorphic scheelite are characterized by low Mo, Nb and Mn, high Sr and Na contents, with commonly positive Eu anomalies. The low Mo content (median value of 21 ppm) and dominant positive Eu anomaly in scheelite from orogenic settings from the Bohemian Massif suggests relatively reduced metamorphic-derived fluids (Song et al., 2014; Poulin et al., 2018; Sciuba et al., 2020). From all the mineral deposit-types studied in the Bohemian Massif, scheelite from orogenic Au deposits shows the lowest median Nb value, which supports this observation. The low (but not lowest) median Mo value is comparable with that of the greisen-type association while the low median Mn value is close to that of RIRGS and reduced skarns. The higher Sr median values were found in the greisen-type association while those of Na were observed in the greisen-type association, RIRGS and oxidized skarns.

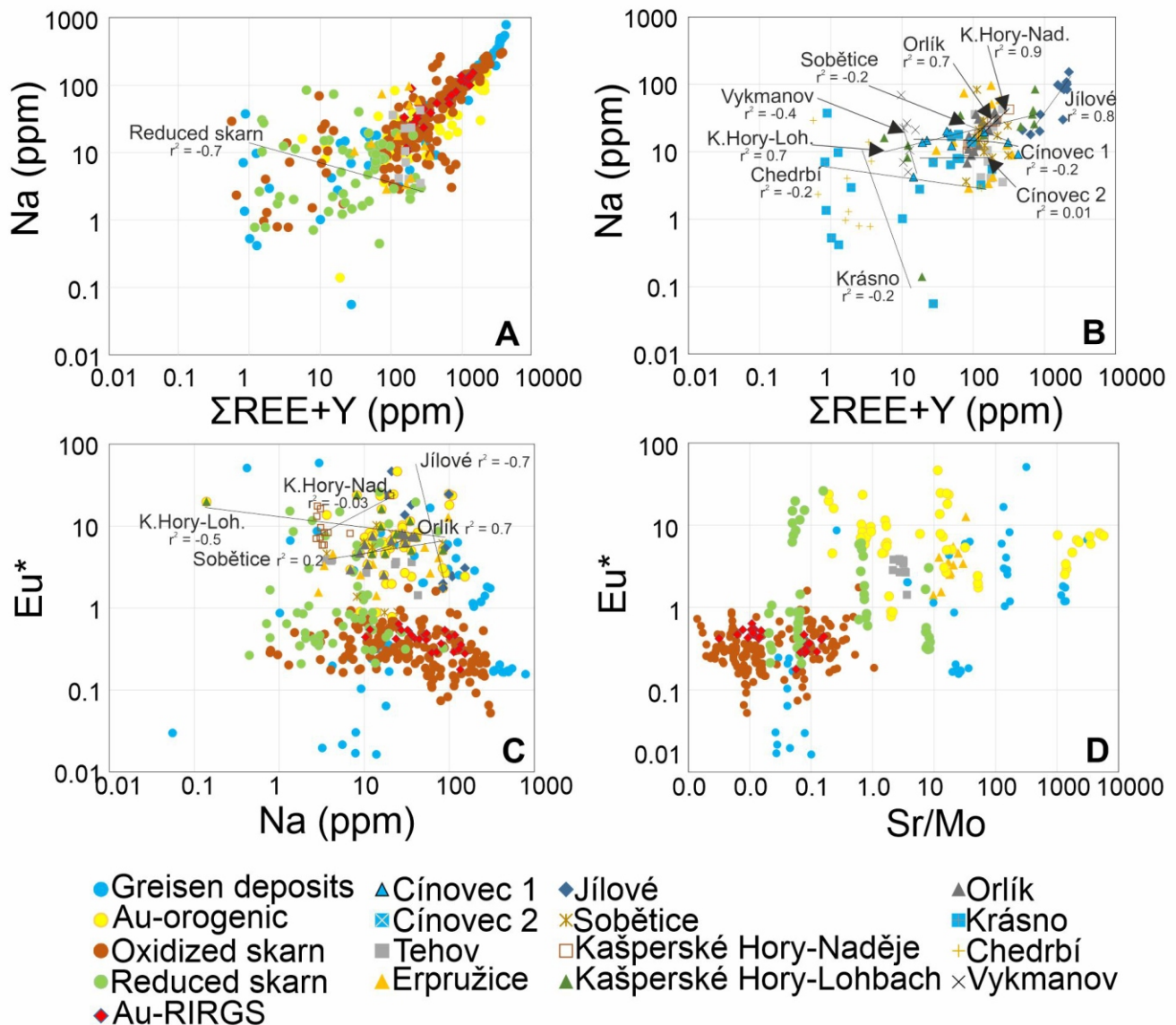


Fig. 10. Binary plots of chemical composition of the scheelite

A – REE + Y versus Na (basic trends), **B** – REE + Y versus Na (selected deposit trends), **C** – Na versus Eu anomaly (selected deposit trends), **D** – Sr/Mo versus Eu anomaly. The values of correlation coefficients (r^2) are shown for selected deposits/mineralizations

Scheelite from orogenic Au deposits can be distinguished from that of oxidized and reduced skarns by higher Sr/Mo values and positive Eu anomalies (Fig. 10D). Scheelite from other settings (quartz veins at Tehov and Erpružice) plots close to that of orogenic Au deposits. Similarly, orogenic Au scheelite shows mostly lower Mo and Nb values when compared to those from oxidized and reduced skarns and RIRGS, and lower Nb values than the group of greisen-related scheelite (Fig. 9D). Scheelite from quartz veins at Tehov and Erpružice again plots close to Au orogenic scheelite. Scheelite from orogenic Au settings can be discriminated from other mineral deposit types in a Sr vs Nb plot (Fig. 9C) and from oxidized skarns by a higher Eu anomaly, lower Mo and larger range of Sr values (see Fig. 9E, F).

RIRGS-RELATED SCHEELITE

Miranda et al. (2024) showed that scheelite from Čelina displays similar REE patterns to that of the felsic intrusions related to the mineralization, suggesting that the exsolved mineralizing fluid REE patterns are controlled by the chemical composition

of the ore-related intrusions and might be reflected by the REE concentrations of scheelite grains. Our new data also show that in comparison with scheelite from orogenic Au settings, scheelite from RIRGS-type mineralizations (Čelina and Vacíkov) is significantly enriched in Mo and Nb, moderately in As and shows similar contents of Mn, Zn, Y and REE. Scheelite from the orogenic Au association contains higher concentrations of Sr, Pb, Th and U. In most of elemental diagrams, scheelite from RIRGS plots close to the field of oxidized skarns (e.g., Fig. 9A–F). From reduced skarns they can be discriminated in a Mo/Sn plot (Fig. 9B).

Scheelite from Vacíkov-Petráčkova hora displays a homogeneous chemical composition and a typical REE pattern with a negative Eu anomaly (Fig. 4F), which is similar to that from Dublin Gulch (intrusion-hosted sheeted quartz veins – not shown, Miranda et al. 2024), consistent with the origin of scheelite from reduced-type fluids (Zachariáš et al., 2001). A predominantly negative Eu anomaly in scheelite from RIRGS results from a different magma composition, salinity and co-genetic REE-bearing minerals (Miranda et al., 2024).

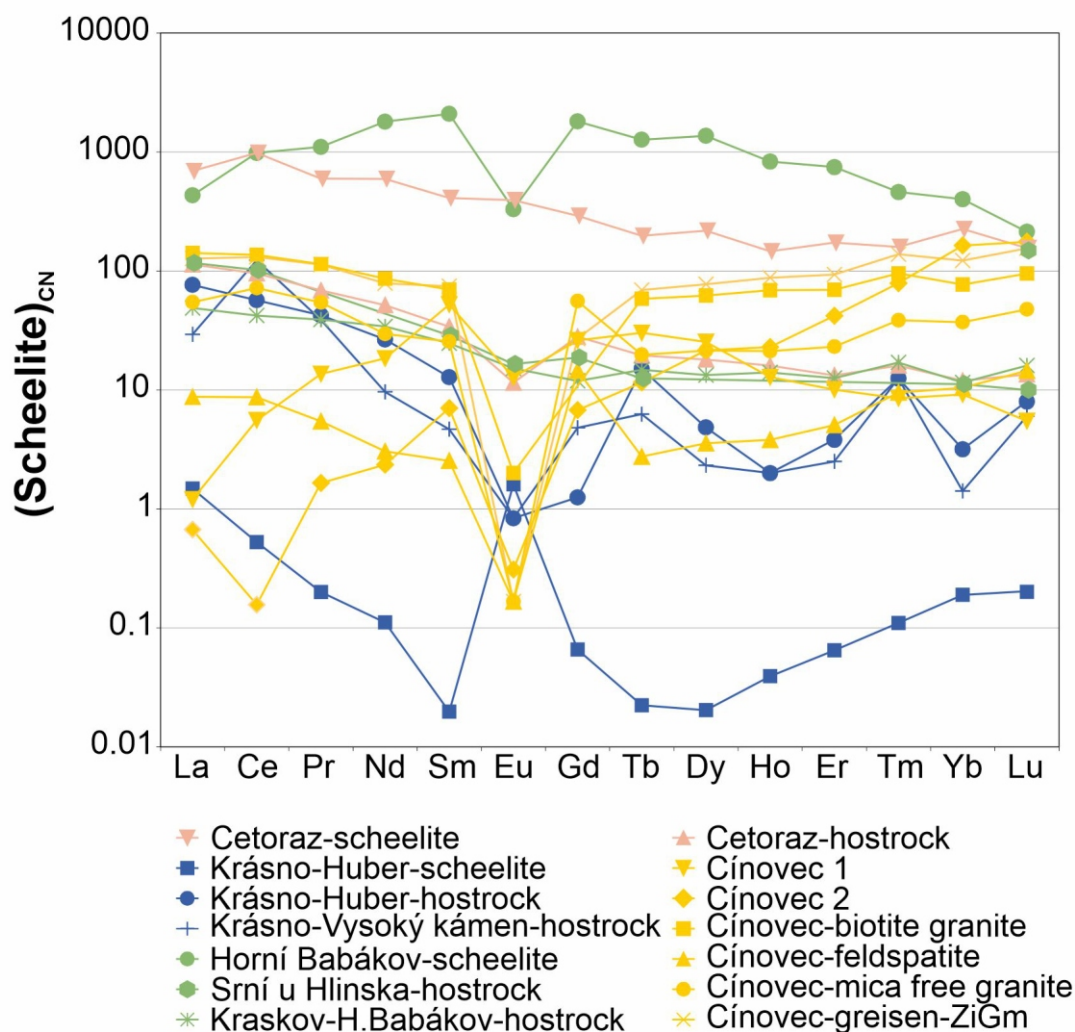


Fig. 11. Rare earth element patterns in selected scheelite (this work) and host rock

Data for Cínovec lithology from Breiter et al. (2017a) and for other hostrocks from the Geochemical Database of the Czech Geological Survey

OXIDIZED AND REDUCED SKARN-RELATED SCHEELITE

On the basis of the study of trace element composition of scheelite from 19 well-documented reduced and oxidized skarn systems, Miranda et al. (2022) concluded that scheelite from reduced and oxidized skarns can be discriminated from those of orogenic and intrusion-related gold deposits due to their lower Sr and higher Mo, Ta, and Nb concentrations. Our study confirmed the highest median Mo, Nb, REE and LREE values in scheelite from oxidized skarn systems of all the mineral deposits/showings studied in the Bohemian Massif. Scheelite from reduced skarns is characterized by the lowest median values of Y and also Mn and As, which are close to those of the orogenic gold setting.

Scheelite from oxidized and reduced skarns is dominated by REE patterns with a negative anomaly, which is consistent with crystallization of plagioclase in their magmatic source rock (Banks et al., 1994; Barker et al., 2004). On the other hand, a combination of increased (high) Sr values with smaller negative Eu anomalies, as documented in scheelite from the Hazlov and Kotel-Krkonoše reduced skarns by Miranda et al. (2024), might indicate release of Sr and Eu from Ca-bearing phases during al-

teration, consistent with observations by Yuan et al. (2019). Large Eu anomalies coupled with low Sr values in scheelite from the reduced skarns at Vrbík and Budislav are somewhat similar to some orogenic Au deposits (Fig. 9E).

The very homogeneous composition of the scheelites is reflected in REE patterns with a negative Eu anomaly from oxidized skarns at Krouna, Proseč and Bílé Labe, and the only slightly inhomogeneous scheelite composition from oxidized skarns at Mutice and Nekvasovy Chlumy is consistent with the findings of Banks et al. (1994), Barker et al. (2004), Miranda et al. (2022) and points to magmatic fluid sources. A change in the shape of REE patterns from LREE-enriched with a negative Eu anomaly (accompanied by significantly higher Y and Nb and lower Mo contents in the scheelite) to flat/bell-shaped with a negative Eu anomaly in oxidized skarn scheelite at Chedrbí documents most likely differences in co-precipitating minerals and in the salinity of the magmatic fluids as reported by Miranda et al. (2022).

On the basis of the comparison of the REE patterns in scheelite grains and host rock at Obří důl (oxidized skarn), Miranda et al. (2024) proposed that the scheelite chemical composition directly reflects, to some degree, the composition of the

mineralizing fluids and host rocks (regardless of the substitution mechanism controlling the REE incorporation in scheelite). They also proposed that the more fractionated REE patterns, such as the steep negative slope at Obří důl, reflect co-precipitation of other mineral phases, types of ligands, and salinity, rather than the original magmatic source.

Primary scheelite from reduced skarns were studied by electron microprobe by Žaludková et al. (2010) who reported higher Mo (up to 2.12 wt.% MoO₃) and low Fe (up to 0.08 wt.% FeO) contents in scheelite from Staré Podhradí and Boží hora u Žulové (Žulovský massif) relative to that from Hostákov and Slavice (Třebíčský massif – up to 0.35 wt.% MoO₃, and up to 0.05 wt.% FeO). They concluded that the chemical composition of the scheelite reflects different chemical compositions of the source fluids and host rock. Miranda et al. (2024) reported a range of 569–2469 ppm Mo (mean 1022 ppm, LA-ICP-MS data) in scheelite from Hostákov. Scheelite from Staré Podhradí showed a high Mo median value (1126 ppm) and low REE (17.5 ppm) and a LREE-enriched pattern with a negative Eu anomaly (this study).

SCHEELITE FROM OTHER SETTINGS

Scheelite from quartz veins at Tehov and Erpružice, characterized by low Mo, Nb, Y, and LREE >HREE, shows many similarities with scheelite from orogenic Au settings (e.g., Fig. 9) and also similar REE patterns (LREE-enriched with positive Eu*).

CONCLUSIONS

The study of chemical composition of scheelite from five major types of deposit within the mineralization of the Bohemian Massif (greisens and related quartz veins, Au-orogenic and Au-RIRGS, oxidized and reduced skarns) can be summarized as follows:

1. Scheelite can display both homogeneous and inhomogeneous chemical compositions within one grain and/or deposit type.
2. Both homogeneous and zoned scheelite seen by cathodoluminescence from greisens and related quartz veins show the highest median values of Sr (281 ppm), Y (88 ppm), Mn (70 ppm), Bi (0.9 ppm), Sn (0.4 ppm), U (0.3 ppm), Th and Zn (0.2 ppm), Hf (0.1 ppm) and HREE (91 ppm) among all the groups studied. Five different types of chondrite-normalized REE pattern in this scheelite reflect different sources and types of fluids and processes. For example, the REE patterns from the massive greisen at Cínovec reflect the REE distribution in the host rock. Scheelites from Cetoraz and Ovesná Lhota are characterized by different REE patterns, indicating changes in fluid chemical composition, whereas scheelite from H. Babákov is typical of a very homogeneous source of fluid.
3. Predominantly homogeneous in CL, scheelite from the Au orogenic setting is typical of the highest median value of Pb (9.2 ppm) and very low median value of Mg

(3.1 ppm) comparable to scheelite from oxidized skarns (3.3 ppm) and a low median Mo value (22 ppm) which is close to that of greisen-type scheelite (11.6 ppm). It also shows the lowest median value of Nb (3 ppm) of all the groups studied. Scheelite from orogenic gold settings displays mostly bell-shaped chondrite-normalized REE patterns and flat REE patterns, both with a positive Eu anomaly.

4. Predominantly homogeneous in CL, scheelite from RIRGS has the highest median values of Si (168 ppm), Mg (10 ppm), As (12 ppm), V (0.8 ppm) and Cu (1 ppm) of all the groups studied. It is, compared to scheelite from orogenic gold deposits, typical of significantly higher Nb and Mo median values and lower Sr, LREE and HREE median values. Scheelite from RIRGS displays mostly bell-shaped normalized REE patterns with a negative Eu anomaly.
5. Predominantly zoned in CL, scheelite from oxidized skarn has the highest median Mo (2494 ppm), REE (274 ppm), LREE (234 ppm) and Nb (77 ppm). Compared to scheelite from reduced skarns, it has significantly higher Y, Mo, LREE and HREE.
6. Predominantly homogeneous in CL, scheelite from reduced skarns is characterized by the lowest median values of Nb (19 ppm), Y (8.2 ppm), Mn (5.9 ppm), and Zn (0.2 ppm) of all the groups studied. A similarly low median As value (1.5 ppm) is close to that of the orogenic gold setting.
7. Scheelite from oxidized and reduced skarns displays five types of chondrite-normalized REE patterns: (1) LREE-enriched with a positive or negative Eu anomaly, (2) Bell-shaped REE patterns with a negative Eu anomaly or without an Eu anomaly, (3) Flat REE patterns with a negative Eu anomaly.
8. Scheelite from other settings (Erpružice u Stříbra and Tehov u Říčan) shows many similarities with that of Au orogenic settings.
9. We suggest that Mo vs. Sn and Nb and Eu* binary plots seem to be suitable, with some exceptions, for determination of scheelite origin from different ore deposit types.

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APPENDIX 1

Basic statistical data on LA-ICP-MS trace element distribution in scheelite in individual deposits within the mineral deposit groups studied

Source file	Category	Na (ppm)	Mg (ppm)	Si (ppm)	K (ppm)	Ti (ppm)	V (ppm)	Mn (ppm)	Ni (ppm)	Cu (ppm)	Zn (ppm)	As (ppm)	Sr (ppm)	Y (ppm)	Nb (ppm)	Mo (ppm)	Ag (ppm)	Sn (ppm)	Sb (ppm)	La (ppm)	Ce (ppm)	Pr (ppm)	Nd (ppm)	Sm (ppm)	Eu (ppm)	Gd (ppm)	Tb (ppm)	Dy (ppm)	Ho (ppm)	Er (ppm)	Tm (ppm)	Yb (ppm)	Lu (ppm)	Hf (ppm)	Ta (ppm)	Pb (ppm)	Bi (ppm)	Th (ppm)	U (ppm)	LREE	HREE
DL_min		0.26	0.008	45.1	0.48	0.067	0.001	0.10	0.03	0.04	0.04	0.03	0.001	0.0007	0.001	0.005	0.003	0.009	0.004	0.0009	0.001	0.0008	0.004	0.003	0.002	0.005	0.0006	0.002	0.0006	0.002	0.0007	0.002	0.0006	0.003	0.0005	0.002	0.001	0.001	0.001		
DL_max		0.78	0.20	175.6	2.10	0.55	0.64	0.31	0.4	0.5	0.7	0.1	0.8	0.09	0.5	0.09	1.3	0.95	3.5	0.20	0.58	0.39	0.13	0.48	0.80	0.49	0.29	0.47	0.19	0.23	0.20	0.58	0.23	0.22	0.22	0.15	0.13	0.24	0.22		
Greisen-/greisenized Q vein - related deposits																																									
H.Babákov	Mean	452.0	278.5	67.3	12.2	2.7	0.3	76.3	1.4	1.2	5.9	10.6	269.6	514.5	44.1	10.9	0.05	0.39	0.02	100.93	586.23	99.73	818.59	311.08	18.73	357.43	46.15	346.05	46.92	124.50	11.98	72.74	5.68	0.16	0.84	5.50	0.04	0.11	0.06	2292.73	654.03
n=12	SD	125.2	24.9	41.1	3.1	0.7	0.3	6.7	3.8	3.4	19.2	9.8	25.1	176.2	18.9	1.2	0.11	0.09	0.02	5.11	43.19	9.24	39.77	7.41	1.17	13.18	2.62	40.95	7.11	26.01	2.92	19.34	1.61	0.02	0.29	1.00	0.06	0.21	0.07	87.01	100.13
SCH-37	Max	792.3	318.3	132.1	17.4	4.6	1.3	93.8	13.9	12.5	69.5	43.1	333.3	950.8	73.3	13.2	0.42	0.65	0.06	107.03	642.18	113.29	899.33	324.57	21.60	381.42	52.13	443.41	63.11	182.23	18.26	113.56	9.23	0.20	1.30	7.05	0.21	0.80	0.28	2470.22	881.64
	Min	331.4	243.0	bdl	7.4	1.8	0.1	69.5	0.2	bdl	bdl	6.9	233.4	373.1	23.8	9.2	0.00	0.29	0.01	89.93	517.08	83.96	739.48	299.41	17.38	336.66	42.91	306.50	40.05	100.91	9.20	54.21	4.10	0.13	0.55	4.32	0.01	0.02	0.01	2136.93	558.62
	Median	431.7	285.2	67.4	11.7	2.6	0.1	74.0	0.2		0.1	7.9	264.1	465.0	39.5	10.8	0.02	0.38	0.01	102.44	601.18	102.37	821.49	309.34	18.57	359.59	45.78	335.96	45.34	119.40	11.38	68.18	5.25	0.17	0.64	5.08	0.02	0.04	0.04	2295.92	632.58
	RSD (%)	0.3	0.1	0.6	0.3	0.2	1.2	0.1	2.7	2.8	3.2	0.9	0.1	0.3	0.4	0.1	213.73%	23.69%	91.21%	5.06%	7.37%	9.26%	4.86%	2.38%	6.25%	3.69%	5.68%	11.83%	15.15%	20.89%	24.40%	26.59%	28.40%	10.62%	35.02%	18.24%	129.23%	189.13%	114.61%	3.79%	15.31%
Ovesná Lhota	Mean	140.0	4.6	60.5	10.7	1.5	0.02	38.0	8.2	2.4	9.7	1.0	241.1	531.6	46.2	2.2	0.05	0.41	0.09	79.55	296.43	30.39	146.29	42.79	45.31	54.12	9.90	92.93	15.27	52.03	6.66	54.71	5.68	0.04	2.99	11.53	1.77	0.23	0.09	694.87	237.18
n=9	SD	54.6	0.3	53.4	2.5	0.3	0.04	4.4	14.7	6.5	19.0	0.5	20.5	257.2	28.5	1.4	0.06	0.11	0.20	34.92	96.29	11.46	80.19	30.43	18.43	46.23	8.24	73.48	11.92	34.89	3.38	21.64	2.09	0.04	1.82	1.61	3.88	0.23	0.07	265.26	154.01
SCH-38	Max	237.0	5.3	154.3	15.6	1.9	0.1	46.4	44.5	20.9	57.0	2.0	272.3	1081.4	108.8	6.1	0.19	0.67	0.65	158.46	471.91	55.88	333.18	113.70	90.07	167.50	29.83	268.69	43.98	134.00	14.19	99.52	9.65	0.14	7.14	14.14	12.61	0.61	0.22	1291.77	599.85
	Min	59.8	4.1	bdl	7.8	1.0	0.001	31.5	0.2	bdl	bdl	0.5	209.3	294.7	20.7	1.5	0.01	0.30	0.00	36.75	153.60	18.01	76.73	15.01	24.78	17.49	3.21	32.55	5.43	22.48	3.60	29.16	2.96	0.01	1.31	8.67	0.01	0.03	0.01	368.84	104.85
	Median	125.9	4.5	44.5	10.3	1.5	0.01	37.9	0.3	0.1	0.2	1.0	239.6	357.0	30.1	1.7	0.03	0.38	0.02	74.92	286.62	24.78	97.86	27.08	40.10	31.25	5.42	51.15	8.37	30.29	4.66	50.46	6.30	0.03	1.90	11.63	0.06	0.09	0.07	739.02	143.40
	RSD (%)	39.0	7.6	88.1	23.2	19.6	187.2	11.6	180.0	265.7	195.3	45.8	8.5	48.4	61.7	64.5	124.19	26.00	213.39	43.90	32.49	37.70	54.81	71.11	40.67	85.42	83.26	79.07	78.05	67.06	50.77	39.55	36.83	90.18	61.01	13.97	219.39	98.65	79.90	38.17	64.93
Cetoraz	Mean	153.2	6.3	75.8	10.8	3.9	0.3	279.8	6.8	6.4	0.4	10.0	651.1	301.7	177.1	35.0	1.4	1.4	0.01	117.98	408.06	49.32	256.52	64.66	30.57	57.05	7.04	51.43	7.37	25.63	3.80	37.49	3.88	0.03	2.54	41.74	2.64	1.63	2.15	984.18	136.65
n=11	SD	111.4	4.5	50.7	2.4	4.8	0.6	106.4	13.9	11.1	0.9	8.8	313.9	213.8	126.1	48.0	4.3	1.3	0.01	93.52	347.89	43.65	228.54	56.10	23.40	46.35	5.13	34.14	4.43	13.71	1.76	14.83	1.34	0.02	1.87	28.04	3.31	1.17	3.42	828.07	74.74
SCH-39	Max	304.4	15.3	180.7	15.6	16.7	2.1	481.5	37.5	33.4	2.9	23.1	1494.3	583.8	361.4	153.5	15.0	5.1	0.03	257.29	824.27	105.78	522.64	130.24	65.61	110.92	13.04	91.32	12.70	43.65	6.34	61.44	6.19	0.07	7.37	123.02	12.23	4.82	12.65	1961.85	231.56
	Min	15.7	3.2	bdl	6.7	1.0	0.0	156.4	0.2	bdl	bdl	0.4	296.8	56.3	9.6	0.4	bdl	0.1	0.00	11.63	14.82	1.30	5.25	2.21	7.37	3.38	0.85	10.16	1.98	8.92	1.72	20.02	2.31	bdl	0.50	21.32	0.06	0.40	0.14	45.96	48.29
	Median	189.7	3.7	63.9	10.4	1.3	0.1	283.4	0.2			14.0	567.1	294.8	243.9	5.1		1.4	0.01	163.60	601.87	55.32	271.33	60.74	22.16	57.82	7.13	53.65	7.99	27.66	3.93	38.27	3.83	0.04	2.66	31.00	2.06	1.56	1.09	1326.53	142.47
	RSD (%)	72.7	72.4	66.8	22.0	121.5	189.3	38.0	205.0	174.4	206.9	88.2	48.2	70.9	71.2	137.3	296.0	98.8	84.21	79.26	85.26	88.50	89.09	86.76	76.55	81.24	72.88	66.39	60.16	53.47	46.33	39.57	34.44	62.38	73.87	67.17	125.18	71.78	159.22	84.14	54.69
Krásno-Huber	Mean	7.6	11.8	63.8	11.3	0.8	0.2	3.7	1.6	1.1	5.4	2.4	234.3	0.4	7.9	1.5	0.2	0.2	0.27	0.37	0.43	0.05	0.37	0.11	0.12	0.15	0.02	0.12	0.02	0.05	0.01	0.07	0.01	0.001	0.39	0.20	0.44	0.03	0.11	1.61	0.29
n=8	SD	11.8	3.8	95.1	3.5	0.5	0.3	1.1	4.1	2.3	11.0	6.3	51.8	0.7	1.4	3.1	0.4	0.2	0.39	0.11	0.35	0.05	0.72	0.29	0.12	0.34	0.04	0.26	0.04	0.08	0.01	0.07	0.01	0.002	0.01	0.12	1.06	0.02	0.09	1.81	0.48
SCH-43	Max	37.5	16.7	290.5	18.1	1.3	0.8	6.0	12.5	7.1	33.4	19.2	311.2	2.2	10.1	9.7	1.3	0.6	1.11	0.53	1.25	0.15	2.21	0.88	0.28	1.06	0.12	0.80	0.12	0.27	0.03	0.19	0.02	0.01	0.41	0.36	3.24	0.08	0.33	6.35	1.54
	Min	0.4	8.4	bdl	7.2	bdl	bdl	2.7	bdl	bdl	bdl	bdl	157.8	bdl	6.3	bdl	bdl	0.0	0.00	0.20	0.00	0.00	0.00	0.00	bdl	0.01	bdl	bdl	bdl	bdl	bdl	0.02	bdl	bdl	0.38	0.01	bdl	0.01	0.04	0.52	0.04
	Median	2.1	9.8	10.4	10.4	0.9	0.1	3.5	0.1	0.1			232.4	0.1	7.8	0.4		0.1	0.01	0.35	0.32	0.02	0.05		0.09	0.01		0.01		0.01		0.03	0.01		0.38	0.22	0.01	0.02	0.08	1.03	0.06
	RSD (%)	155.2	31.8	149.1	30.9	63.1	127.5	28.1	252.4	215.2	204.0	263.8	22.1	186.2	17.3	205.9	263.8	127.4	146.34	29.74	81.26	108.00	192.67	257.37	93.80	224.42	222.99	223.92	218.62	169.62	156.13	94.13	101.42	153.30	2.35	61.83	242.98	79.81	80.32	112.71	167.78
Cínovec 1-Voj Jáma 3.p.	Mean	14.6	23.7	46.8	13.8	7.2	5.0	85.3	0.3	14.3	55.8	11485.6	283.4	24.7	35.2	4639.1	9.0	9.5	0.71	0.60	8.76	2.94	17.87	11.45	0.66	8.59	1.83	12.31	1.76	5.77	1.10	11.91	1.66	1.35	3.57	5947.81	3187.69	219.77	2534.25	50.87	36.35
n=9	SD	4.9	5.8	48.8	4.2	3.7	2.5	24.5	0.2	8.7	15.5	1171.3	36.7	34.8	11.4	1035.9	3.7	8.5	0.57	0.73	11.97	3.70	20.82	10.13	0.34	8.31	1.76	13.30	2.18	7.91	1.61	18.33	2.71	1.57	1.95	5101.23	4085.02	114.03	283.89	54.35	47.08
SCH-47	Max	20.3	34.8	141.8	18.8	14.3	10.3	110.5	0.6	30.6	73.9	14442.1	339.4	98.1	53.0	6171.6	15.6	32.8	1.64	2.06	35.72	10.90	68.25	36.82	1.17	29.50	6.28	46.57	7.39	26.01	5.12	54.07	7.34	4.52	7.85	14859.10	11662.10	444.52	2998.23	177.69	152.77
	Min	4.2	13.8	bdl	7.5	3.3	1.2	34.9	0.0	0.0	26.1	10049.7	231.3	2.3	16.0	3405.5	1.1	bdl	0.08	0.06	0.58	0.28	2.17	1.75	0.10	1.33	0.28	1.89	0.26	0.62	0.09	0.71	0.04	0.11	1.18	1220.42	183.48	33.36	2051.64	6.28	4.74
	Median	14.3	23.6	36.1	15.2	5.9	4.7	94.9	0.3	12.5	59.4	11370.2	270.9																												

SCH-41	Min	18.1	2.7	0.0	8.5	bdl	bdl	2.2	0.2	bdl	bdl	4.0	308.6	100.9	1.3	5.9	0.0	bdl	0.00	27.22	67.00	7.09	34.61	10.95	110.09	13.61	2.07	15.12	2.13	5.87	0.78	6.64	0.60	0.01	0.36	7.46	bdl	0.02	bdl	358.58	33.22		
	Median	84.3	5.8	0.0	13.6	0.7	0.1	3.3	0.2	0.1		13.8	345.4	462.2	2.3	22.0	0.0		0.00	39.75	209.26	34.61	234.91	95.98	187.51	125.10	18.14	119.61	15.45	34.54	2.96	15.30	0.98	0.07	0.41	9.12	0.01	0.13	0.19	1079.82	213.21		
	RSD (%)	59.8	117.4	174.5	50.5	174.0	146.5	20.7	212.8	181.5	260.8	88.8	8.5	43.0	45.0	46.5	107.1		185.22	25.45	32.32	43.38	56.58	67.04	44.92	69.05	67.27	64.06	60.50	52.87	42.21	37.23	43.09	52.17	9.14	24.28	157.86	60.17	129.34	35.00	58.13		
Soběčice	Mean	21.2	4.3	23.6	5.0	1.0	0.1	6.4	2.1	1.7	0.3	1.6	34.6	55.8	2.1	20.0	0.1		3.11	8.66	24.98	3.77	23.59	7.22	7.70	9.37	1.35	11.54	1.92	5.97	0.64	4.55	0.42	0.01	0.34	15.26	0.11	0.04	0.04	85.28	26.39		
n=12	SD	20.1	2.3	36.6	4.6	0.2	0.1	0.9	6.1	5.4	0.6	1.9	1.9	21.1	0.4	2.8	0.2	bdl	7.32	2.52	7.92	1.72	14.53	5.45	2.77	7.76	1.05	8.37	1.24	3.02	0.21	0.87	0.06	bdl		0.01	10.28	0.21	0.03	0.05	36.49	14.59	
SCH-42	Max	83.4	10.1	116.8	15.3	1.2	0.5	8.6	22.2	19.5	2.4	7.3	39.1	101.3	2.7	26.0	0.5	bdl	25.44	14.32	39.34	7.28	54.49	19.25	11.40	26.45	3.60	28.86	4.56	12.72	1.13	6.57	0.57	0.02	0.35	47.12	0.75	0.10	0.14	163.49	57.23		
	Min	3.6	2.6	bdl	bdl		0.7	bdl		bdl		0.3	31.1	27.5	1.5	15.8	bdl	bdl	0.02	5.39	14.04	1.67	8.99	2.28	3.79	2.77	0.43	4.06	0.75	2.83	0.36	3.16	0.34	bdl		0.32	10.08	0.01	0.01	0.01	39.93	11.93	
	Median	15.6	3.0		4.7	1.0		6.2	0.2		0.1	0.9	34.7	46.5	2.1	20.8			0.11	8.37	24.81	3.18	17.51	4.39	8.65	5.36	0.82	7.44	1.31	4.46	0.56	4.30	0.40	0.01	0.33	11.21	0.02	0.04	0.02	74.15	19.01		
	RSD (%)	94.8	54.0	155.0	92.7	17.0	138.1	13.9	288.6	319.4	249.8	115.0	5.6	37.9	17.8	14.2	163.5		235.52	29.06	31.69	45.63	61.57	75.54	35.95	82.85	78.05	72.50	64.37	50.65	32.97	19.25	15.17		2.49	67.36	195.83	59.95	104.93	42.79	55.28		
Kašperské hory-Naděje	Mean	19.8	4.9	45.1	8.9	1.8	0.1	14.7	0.2	1.3	5.7	1.1	83.2	47.9	7.0	110.3		0.1	0.05	12.03	26.38	3.01	13.39	3.11	10.39	3.27	0.51	5.16	1.00	4.64	0.91	11.36	1.44		0.43	7.79	0.05	0.20	0.14	71.58	25.02		
n=12	SD	2.1	199.7	268.2	7.6	0.6	0.3	0.8	22.9	2.0	0.5	1.1	3.4	5.9	10.4	39.1		0.1	0.01	0.34	1.27	0.29	3.00	1.41	0.28	2.40	0.25	2.03	0.34	1.02	0.09	0.55	0.05		0.01	0.61	0.05	0.02	0.06	8.34	4.21		
SCH-5	Maximum	42.4	20.4	204.3	21.4	7.1	0.7	18.9	0.4	11.1	66.7	1.8	93.0	158.2	12.6	123.4	0.2	0.2	0.21	28.66	57.41	6.09	25.01	5.02	15.55	5.34	1.10	12.94	2.75	14.43	2.88	30.83	3.47	0.01	0.52	12.29	0.22	0.39	0.23	138.35	68.41		
	Minimum	10.1	2.7	bdl	bdl		0.8	bdl		bdl		0.3	74.9	19.1	3.1	90.4	bdl	0.0	0.01	4.11	13.84	1.77	6.71	1.28	5.17	1.30	0.20	2.11	0.41	1.87	0.36	4.81	0.64	bdl		0.36	5.57	0.01	0.04	0.07	41.75	10.76	
	Median	16.2	3.1	12.5	7.8	1.3		15.2	0.2	0.1		1.1	82.4	33.1	6.5	112.9		0.1	0.02	10.90	24.19	2.71	11.70	2.70	10.61	2.92	0.43	4.20	0.80	3.45	0.61	7.62	1.01		0.43	7.05	0.02	0.19	0.14	66.76	17.16		
	RSD (%)	10.9	4103.5	595.3	85.1	30.8	346.9	5.4	10335.2	148.1	8.1	106.9	4.1	12.3	148.1	35.4		71.9	20.44	2.79	4.79	9.80	22.42	45.45	2.67	73.24	48.57	39.25	33.94	21.95	9.85	4.85	3.30		2.03	7.79	98.17	11.07	41.92	11.65	16.83		
Kašperské hory-Lohbach	Mean	26.3	24.8	279.2	12.2	0.7	0.2	29.9	7.8	4.4	8.4	2.7	182.6	68.8	8.9	138.1		0.2	0.09	36.48	80.89	8.36	37.71	8.56	17.94	8.28	1.31	11.34	2.04	7.57	1.22	12.01	1.53	0.004	0.54	13.37	0.26	2.14	6.34	198.22	37.01		
n=12	SD	20.9	68.8	696.8	5.9	0.6	0.4	21.8	13.3	10.1	19.3	4.6	94.5	70.8	3.7	118.3		0.2	0.06	42.19	87.50	8.53	38.22	8.67	16.73	8.40	1.36	11.88	2.12	7.82	1.25	12.30	1.53	0.01	0.05	10.41	0.16	1.81	6.81	203.38	38.17		
SCH-6	Maximum	84.4	253.0	2557.6	29.2	2.2	1.4	67.0	40.3	34.8	69.5	16.9	292.7	191.5	15.6	309.7	bdl	0.7	0.22	137.61	229.35	20.42	100.50	23.88	40.31	22.34	3.60	30.58	5.34	19.75	3.13	30.80	3.78	0.02	0.62	35.00	0.67	4.77	25.30	523.61	95.82		
	Minimum	0.1	1.7	bdl	6.1	bdl	bdl	10.2	bdl	bdl	bdl	0.3	59.4	0.8	3.3	15.0	bdl	bdl	0.01	1.89	1.47	0.14	0.42	0.07	0.36	0.07	bdl		0.08	0.01	0.07	0.02	0.15	0.02	bdl		0.49	3.70	0.07	0.07	0.08	4.42	0.35
	Median	20.6	4.1	36.8	11.4	0.8		15.1	0.3	0.1	0.8	0.8	189.6	27.0	8.5	105.8			0.08	12.93	29.36	3.72	18.30	4.28	8.81	3.82	0.51	4.18	0.74	2.79	0.45	4.32	0.61		0.51	7.19	0.22	1.27	6.07	80.82	13.60		
	RSD (%)	79.7	277.6	249.6	48.7	85.5	239.1	72.9	171.6	230.5	229.5	171.9	51.7	102.9	41.1	85.6		163.5	66.11	115.65	108.17	101.99	101.35	101.35	93.24	101.39	104.35	104.77	104.16	103.33	102.28	102.44	99.96	128.21	9.34	77.87	62.77	84.27	107.27	102.60	103.15		
Orlík u Humpolce	Mean	20.9	4.7	47.1	17.5	1.2	0.1	4.3	3.4	4.5	5.3	6.5	2236.9	36.9	2.7	0.9	0.1	0.1	0.01	5.75	18.35	2.86	18.13	5.82	13.68	7.95	1.19	8.74	1.36	3.54	0.32	1.67	0.16	0.005	0.54	8.68	0.09	0.11	1.88	72.54	16.99		
n=12	SD	10.9	6.4	67.5	6.4	1.1	0.1	0.9	6.8	6.5	16.8	1.2	985.3	18.2	0.9	0.5	0.1	0.2	0.03	0.59	2.98	0.67	4.35	1.79	7.41	2.50	0.44	3.53	0.55	1.51	0.15	0.87	0.12	0.004	0.02	7.30	0.19	0.16	3.42	18.78	7.00		
SCH-7	Maximum	38.9	25.4	237.7	33.4	4.4	0.3	6.0	25.0	19.5	61.2	8.4	3600.1	72.0	5.6	2.1	0.4	0.6	0.10	6.99	25.77	4.52	27.24	9.92	28.63	13.93	2.15	15.72	2.37	6.44	0.60	3.21	0.50	0.01	0.57	21.41	0.70	0.51	10.29	111.67	30.02		
	Minimum	6.9	bdl	bdl	7.8	bdl	bdl	3.4	bdl	bdl	bdl	3.9	843.7	14.7	1.6	0.5	bdl	bdl	bdl	4.95	15.14	2.33	14.08	4.31	4.73	5.57	0.75	4.85	0.75	1.66	0.13	0.61	0.06	0.001	0.51	1.41	bdl	bdl	bdl	54.55	8.87		
	Median	18.9	2.5	10.4	16.7	1.1	0.1	4.0	0.2	0.7		6.9	2117.7	35.1	2.5	0.7				5.58	16.87	2.52	16.01	5.18	12.91	7.00	1.06	7.71	1.19	3.38	0.32	1.67	0.15	0.004	0.54	6.08	0.01	0.02	0.01	65.34	15.32		
	RSD (%)	52.2	136.6	143.2	36.7	89.4	71.6	19.9	199.3	144.8	317.9	19.0	44.1	49.4	35.2	52.7	139.1	159.5	205.31	10.26	16.26	23.32	24.00	30.79	54.15	31.40	36.87	40.40	40.30	42.71	46.62	51.77	72.80	71.80	3.22	84.09	223.22	140.71	181.77	25.89	41.23		
Au-RIRGS-related	Mean	36.7	35.2	137.6	11.7	1.8	0.6	12.1	6.7	8.0	4.4	12.1	29.1	109.4	23.9	3021.4	0.3	0.1	0.46	3.94	14.81	3.06	29.52	12.44	2.79	24.67	2.72	19.75	3.85	10.40	0.84	4.19	0.38	0.01	0.48	1.85	0.05	0.01	0.02	91.21	42.11		
Čelina	SD	19.7	50.0	255.3	6.8	1.6	0.3	5.6	14.4	12.2	11.1	4.8	4.8	45.7	14.3	1062.5	0.6	0.3	1.16	1.87	5.95	1.03	9.61	3.37	0.62	7.04	0.79	5.93	1.28	4.05	0.38	2.12	0.18	0.01	0.04	0.52	0.06	0.01	0.02	26.51	14.09		
n=13	Max	89.0	168.2	973.0	32.0	6.0	1.2	27.2	51.4	37.5	38.9	19.8	37.4	240.3	66.4	6199.4	2.1	1.0	4.45	7.94	27.69	5.50	53.38	19.68	4.17	37.39	4.34	33.32	7.20	22.42	2.02	10.40	0.79	0.02	0.60	2.53	0.19	0.03	0.05	154.92	80.48		
SCH-3	Min	10.3	0.4	bdl	4.6	bdl		0.2	7.9	bdl	bdl	bdl	4.7	19.6	52.8	8.9	2275.4	bdl	bdl	0.00	2.11	7.30	1.66	18.33	8.63	1.89	15.50	1.60	11.33	2.15	6.00	0.50	1.72	0.12	bdl		0.42	0.52	bdl	bdl	bdl	62.75	25.36
	Median	33.4	10.7	29.2	10.3	1.4	0.5	9.6		0.1		11.8	28.0	100.6	20.5	2588.2			0.03	3.00	14.03	2.70	25.97	11.80	2.68	23.32	2.67	20.16	3.86	9.62	0.73	3.38	0.36	0.01	0.48	1.92	0.01	0.01		81.80	40.02		
	RSD (%)	53.7	142.1	185.5	58.2	93.0	57.1	46.3	215.8	153.5	250.5	39.8	16.5	41.8	59.7	35.2	231.2	189.1	252.99	47.50	40.14	33.72	32.57	27.13	22.14	28.53	28.89	30.04	33.11	38.93	45.88	50.58	47.96	58.15	8.69	28.30	127.71	67.48	114.20	29.07			

n=3	SD	4.3	3.9	275.2	6.1	0.3	1.9	49.3	0.2	39.3	0.7	0.1	3.6	0.8	0.1	30.6			0.01	0.35	6.44	0.11	0.45	0.08	0.01	0.32	0.02	0.12	0.04	0.12	0.02	0.10	0.01	0.02	0.03	3.63	0.01	0.32	0.02	7.23	0.40
SCH-16	Maximum	10.8	16.7	583.8	45.9	1.9	4.2	107.0	0.6	83.4	1.7	0.7	111.9	4.5	1.7	1052.2	bdl	bdl	0.02	0.78	13.90	0.28	1.29	0.32	0.05	0.96	0.06	0.67	0.17	0.75	0.09	0.71	0.06	0.05	0.42	9.27	0.03	0.73	0.04	16.93	2.50
	Minimum	1.5	7.2	bdl	32.2	1.1	0.0	2.1	0.2	bdl	bdl	0.3	103.4	2.7	1.5	977.2	bdl	bdl	bdl	0.01	0.10	0.02	0.18	0.13	0.03	0.26	0.03	0.41	0.08	0.44	0.05	0.48	0.04	bdl	0.36	0.97	0.002	0.04	bdl	0.74	1.53
	Median	1.7	11.5		44.5	1.7	0.5	2.8	0.3	0.1	0.8	0.5	105.5	4.2	1.6	1014.7			0.01	0.06	0.40	0.07	0.78	0.28	0.04	0.32	0.06	0.66	0.14	0.61	0.08	0.55	0.05		0.40	2.34	0.01	0.05		2.61	2.17
	RSD (%)	92.5	32.9	141.4	15.0	21.7	118.6	132.1	50.4	141.2	80.1	29.0	3.4	20.6	4.6	3.0			86.02	122.47	134.04	88.75	60.42	33.44	24.15	61.82	32.38	21.14	29.26	20.61	23.31	16.81	19.07	141.42	6.89	86.64	91.42	116.60	141.42	106.99	19.54
Vykmanov	Mean	23.2	75.6	478.0	26.1	2.5	1.2	13.9	0.2	0.2	12.1	0.6	114.8	3.1	1.6	1573.9	1.1	0.3	0.05	0.83	2.02	0.36	2.18	0.73	0.22	0.89	0.09	0.75	0.13	0.35	0.03	0.19	0.02		0.51	8.64	1.19	0.62	0.23	7.23	1.57
n=7	SD	20.4	38.3	383.3	25.3	1.9	2.0	12.9	0.1	0.1	23.6	0.3	25.8	0.3	0.3	483.6	1.4	0.5	0.04	0.66	0.31	0.14	0.82	0.29	0.07	0.22	0.04	0.16	0.02	0.04	0.00	0.07	0.01	0.003	0.02	4.23	1.17	0.50	0.09	1.39	0.24
SCH-44	Maximum	69.5	127.9	875.7	82.0	7.0	6.1	34.6	0.4	0.4	69.5	1.2	159.2	3.6	2.2	2446.4	4.3	1.5	0.11	2.36	2.59	0.65	3.29	1.24	0.36	1.32	0.15	1.01	0.16	0.42	0.04	0.27	0.03	0.01	0.54	17.38	3.48	1.77	0.37	9.81	2.02
	Minimum	5.0	21.4	bdl	bdl	1.0	0.1	4.0	0.1	bdl	bdl	0.0	74.6	2.7	1.4	949.4	0.1	bdl	bdl	0.18	1.64	0.17	1.25	0.38	0.15	0.64	bdl	0.58	0.10	0.29	0.03	0.07	0.02	bdl	0.47	3.28	0.07	0.20	0.09	5.50	1.30
	Median	21.1	92.4	569.9	15.3	1.6	0.3	5.3	0.3	0.2	1.3	0.7	107.3	3.0	1.5	1495.6	0.4	0.1	0.05	0.73	1.88	0.33	1.68	0.57	0.17	0.85	0.09	0.65	0.12	0.35	0.04	0.23	0.02	0.003	0.52	7.92	0.56	0.52	0.27	7.47	1.50
	RSD (%)	88.1	50.7	80.2	97.2	77.5	167.8	93.0	45.5	75.6	194.6	54.1	22.5	8.8	17.8	30.7	131.3	186.5	75.38	78.87	15.14	40.04	37.83	39.79	32.58	24.64	47.01	20.94	18.04	10.50	12.25	33.37	27.23		4.30	48.93	98.00	81.03	39.03	19.18	15.12
Bilé Labe	Mean	230.6	9.7	61.0	12.9	2.4	1.2	88.2	17.0	12.2		5.3	43.5	522.4	141.5	2724.3		0.4	0.07	256.19	564.37	56.55	286.28	72.27	4.84	74.99	11.13	87.83	15.63	52.92	7.45	66.54	8.30	0.04	9.98	25.62	6.55	0.27	0.11	1315.49	249.81
n=10	SD	41.9	6.7	49.6	3.2	3.4	0.9	5.8	37.1	26.7		1.5	2.4	73.3	30.0	908.1		0.2	0.12	80.10	159.12	13.69	59.66	10.48	0.88	10.22	1.83	15.01	2.63	8.70	1.17	10.72	1.51	0.01	2.86	4.14	8.95	0.23	0.10	316.29	39.00
SCH-28	Maximum	271.1	29.2	161.2	17.0	12.5	3.8	98.6	125.1	86.2	0.1	8.4	50.2	633.8	186.5	3391.6	0.2	0.8	0.42	330.40	697.78	67.28	327.21	84.93	7.20	93.27	13.94	111.90	19.84	67.28	9.20	79.65	9.86	0.05	12.41	33.64	24.60	0.76	0.33	1565.79	311.65
	Minimum	118.0	4.6	bdl	8.1	0.8	0.1	77.8	0.1	bdl	bdl	2.7	41.3	409.1	83.4	154.3	bdl	bdl	bdl	42.53	119.82	17.18	112.31	47.54	4.14	62.41	8.42	66.03	11.70	40.45	6.02	44.33	4.50	0.03	2.32	16.00	0.01	0.05	0.00	411.63	198.71
	Median	244.1	8.2	54.2	11.8	1.5	1.0	89.6	0.2	0.1		4.9	42.7	528.9	143.6	2914.8		0.4	0.01	274.59	594.92	59.42	308.86	73.46	4.45	73.25	11.33	89.79	15.81	52.26	7.47	67.35	8.51	0.04	11.00	25.45	0.65	0.18	0.07	1375.82	249.60
	RSD (%)	18.2	68.6	81.3	24.5	139.5	73.4	6.6	218.9	219.0		28.7	5.6	14.0	21.2	33.3		53.4	173.56	31.26	28.19	24.21	20.84	14.50	18.17	13.63	16.48	17.09	16.85	16.44	15.66	16.11	18.18	20.13	28.64	16.17	136.62	83.23	91.90	24.04	15.61
Chedrbí	Mean	6.0	3.7	74.4	10.3	3.9	0.2	2.9	3.6	4.0		0.7	33.7	2.3	19.4	1808.0		0.1	0.01	4.36	7.78	0.90	5.26	1.31	0.15	1.19	0.11	0.70	0.09	0.25	0.02	0.12	0.01	0.001	1.37	2.60	0.02	0.04	0.06	20.94	1.31
n=11	SD	8.2	1.4	89.1	2.3	5.1	0.1	2.2	7.2	7.3	0.1	0.6	7.8	3.5	38.5	564.1		0.1	0.02	9.14	16.16	1.85	10.55	2.46	0.28	1.96	0.16	0.98	0.12	0.29	0.02	0.14	0.01	0.001	2.10	0.74	0.01	0.06	0.08	42.38	1.73
SCH-9	Maximum	29.2	6.6	318.3	14.3	18.1	0.5	7.8	19.5	23.6	0.2	2.1	50.6	10.6	108.3	2233.7	bdl	0.2	0.06	24.99	43.23	4.96	28.63	6.96	0.76	5.53	0.48	2.95	0.37	0.91	0.07	0.44	0.03	0.003	5.87	3.79	0.04	0.19	0.23	115.08	5.26
	Minimum	0.8	2.5	bdl	7.5	1.1	0.0	1.2	0.2	bdl	bdl	0.2	24.2	0.2	1.2	618.6	bdl	bdl	0.00	0.01	0.03	0.01	0.06	0.04	0.01	0.06	0.01	0.08	0.01	0.02	bdl	0.02	0.00	bdl	0.37	1.38	0.01	0.01	bdl	0.24	0.14
	Median	2.9	2.8	58.4	10.6	1.3	0.1	2.1	0.2			0.6	32.2	0.8	1.3	2037.7			0.01	0.04	0.11	0.03	0.32	0.19	0.02	0.33	0.03	0.29	0.04	0.14	0.01	0.06	0.01	0.001	0.38	2.88	0.01	0.01	0.02	1.15	0.58
	RSD (%)	135.7	39.2	119.7	22.5	130.8	90.3	76.7	198.9	181.6		76.8	23.1	151.4	198.8	31.2		83.8	125.42	209.91	207.67	205.73	200.75	187.25	185.89	164.35	148.64	139.96	129.08	119.35	109.53	111.24	77.52	121.86	153.43	28.30	68.88	142.76	132.47	202.38	132.16
Krouna	Mean	13.3	128.9	195.5	12.5	1.5	0.3	5.3	18.2	0.8	0.3	7.6	104.9	48.3	33.1	483.3		0.1	0.02	4.50	18.50	4.05	39.39	15.92	2.21	22.06	2.28	16.06	2.66	7.36	0.62	3.74	0.37	0.01	0.39	8.99	0.06	0.03	0.09	106.62	33.09
n=9	SD	2.1	199.7	268.2	7.6	0.6	0.3	0.8	22.9	2.0	0.5	1.1	3.4	5.9	10.4	39.1		0.1	0.01	0.34	1.27	0.29	3.00	1.41	0.28	2.40	0.25	2.03	0.34	1.02	0.09	0.55	0.05	0.004	0.01	0.61	0.05	0.02	0.06	8.34	4.21
SCH-45	Maximum	18.1	550.4	720.0	30.6	2.8	1.3	6.8	59.8	6.4	1.5	8.9	109.3	59.8	52.5	589.4	bdl	0.2	0.04	5.05	20.63	4.64	45.18	18.61	2.56	25.80	2.72	19.59	3.28	9.40	0.81	4.99	0.48	0.02	0.41	9.94	0.15	0.09	0.21	122.45	41.27
	Minimum	10.7	2.7	bdl	4.2	0.8	0.1	3.9	0.2	bdl	bdl	5.4	98.4	39.3	20.4	446.2	bdl	bdl	0.01	3.79	16.19	3.57	34.04	13.75	1.73	18.65	1.94	13.21	2.16	5.75	0.46	2.82	0.30	0.002	0.38	7.96	0.01	0.01	0.02	92.03	26.65
	Median	12.9	14.6	51.4	9.5	1.4	0.2	5.4	0.3		0.2	8.4	104.7	46.3	30.4	476.8		0.1	0.02	4.59	18.60	4.02	39.48	16.17	2.36	23.24	2.39	16.26	2.64	7.21	0.61	3.77	0.37	0.01	0.39	9.12	0.03	0.02	0.08	105.25	32.73
	RSD (%)	16.2	154.9	137.2	60.8	37.7	107.2	15.0	125.8	265.2	136.9	14.9	3.3	12.2	31.5	8.1		67.7	43.12	7.45	6.84	7.28	7.63	8.87	12.53	10.86	10.97	12.61	12.73	13.85	14.42	14.74	12.76	48.48	2.25	6.74	93.01	76.68	68.00	7.82	12.73
Mutice	Mean	9.7	22.4	109.5	6.1	1.2	0.2	23.1	0.8	5.6	0.9	1.8	65.6	85.4	256.0	290.8		0.2	0.04	33.31	104.64	14.29	93.35	23.57	2.06	25.11	2.71	18.98	3.06	9.71	1.12	8.76	0.86	0.01	0.40	0.69	0.01	0.20	1.13	296.34	45.20
n=12	SD	4.5	28.5	102.9	5.1	0.4	0.3	8.9	1.3	6.6	1.7	0.4	16.6	44.3	139.6	99.5		0.1	0.06	13.88	40.57	4.06	29.86	9.23	0.51	10.60	1.21	8.59	1.39	4.27	0.44	2.62	0.20	0.01	0.12	0.33	0.01	0.12	1.30	94.56	17.78
SCH-10	Maximum	18.6	86.7	287.7	16.7	2.4	1.1	50.0	4.4	18.1	6.3	2.5	90.2	180.7	531.0	419.8	bdl	0.3	0.18	58.10	187.09	20.72	135.80	36.18	2.81	38.36	4.56	33.36	5.53	17.72	1.99	13.58	1.35	0.03	0.77	1.54	0.03	0.44	4.84	460.48	77.71
	Minimum	3.0	2.8	bdl	bdl	0.8	0.0	14.6	0.1	bdl	bdl	1.1	33.4	30.8	85.2	90.6	bdl	0.1	0.01	16.03	60.05	9.10	45.45	8.58	1.28	7.85	0.83	6.06	1.01	3.49	0.48	4.99	0.62	0.003	0.32	0.35	0.01	0.06	0.21	180.89	17.58
	Median	8.5	8.5	1																																					

	RSD (%)	41.3	42.8	101.4	20.0	22.4	27.8	11.1	103.6	252.6	291.7	35.0	5.6	26.9	57.0	25.4	166.7		238.75	20.99	22.58	21.43	20.49	17.60	30.29	19.50	22.96	26.92	30.06	31.31	28.85	34.21	45.69	75.04	15.12	51.62	204.84	48.69	84.18	19.62	26.34
Reduced skarn	Mean	18.9	1.8	37.9	12.3	1.5	0.1	16.3	0.4	3.7	1.6	1.1	42.9	67.1	10.2	250.4		0.1	0.01	2.06	8.51	2.03	19.62	8.85	3.44	14.53	1.95	14.85	2.68	8.06	0.82	5.28	0.62	0.01	0.54	4.04	0.10	0.02	0.03	59.03	34.25
Budislav	SD	12.4	1.8	31.8	5.2	1.2	0.2	2.2	0.7	7.4	5.0	0.5	15.3	81.5	10.4	85.6		0.0	0.02	1.08	3.45	0.74	7.19	5.20	1.04	9.77	1.72	14.81	2.75	8.82	1.00	7.10	0.80	0.01	0.13	0.82	0.26	0.01	0.05	25.53	36.97
n=12	Maximum	55.6	5.8	82.0	23.6	5.3	0.8	19.5	2.6	20.9	18.1	2.2	68.5	334.6	41.8	348.9	bdl	0.2	0.04	5.31	16.12	3.81	37.25	24.48	5.13	45.59	7.52	63.25	11.68	36.78	4.06	28.30	3.18	0.03	0.95	5.56	0.97	0.04	0.18	132.65	154.77
SCH-13	Minimum	1.0	bdl	bdl	bdl	0.4	bdl	12.9	bdl	0.0	bdl	0.6	28.0	27.8	3.2	106.1	bdl	bdl	0.00	1.23	5.27	1.31	13.11	5.18	1.82	7.99	1.00	6.96	1.23	3.45	0.28	1.40	0.12	bdl	0.47	3.04	0.004	0.003	bdl	38.84	15.04
	Median	16.3	2.5	34.1	12.7	1.1	0.1	17.0	0.2	0.1		0.9	32.5	41.2	6.3	280.5			0.01	1.64	6.99	1.65	17.07	7.03	3.26	11.90	1.44	10.65	1.93	5.53	0.50	2.84	0.37	0.01	0.50	3.79	0.01	0.01	0.01	49.78	23.26
	RSD (%)	65.7	96.5	83.9	42.2	84.5	151.6	13.3	166.3	198.7	302.0	48.3	35.6	121.5	102.2	34.2		67.3	116.65	52.53	40.50	36.56	36.67	58.81	30.31	67.21	88.09	99.75	102.66	109.47	122.25	134.45	130.36	77.86	23.33	20.38	271.81	81.34	183.88	43.25	107.93
Hazlov	Mean	15.9	28.1	109.1	10.6	1.6	0.3	1.5	4.2	4.1	0.2	0.9	1598.8	0.9	24.6	204.5		0.2	0.07	0.37	0.32	0.06	0.34	0.08	0.02	0.11	0.02	0.13	0.03	0.11	0.01	0.11	0.03		0.63	1.83	0.03	0.02	0.22	1.31	0.44
n=11	SD	24.0	68.5	102.5	4.2	0.5	0.5	1.8	10.7	7.7	0.4	0.7	56.0	0.6	9.0	7.9		0.2	0.11	0.35	0.15	0.05	0.32	0.05	0.03	0.08	0.01	0.05	0.01	0.04	0.01	0.05	0.04		0.12	0.08	0.04	0.01	0.52	0.88	0.17
SCH-31	Max	84.8	244.6	244.6	19.5	3.2	1.4	5.1	37.5	26.4	1.0	2.5	1720.8	2.4	47.3	219.6	bdl	0.4	0.33	1.30	0.74	0.18	1.16	0.16	0.10	0.31	0.04	0.22	0.06	0.19	0.03	0.24	0.14	bdl	0.83	1.95	0.14	0.03	1.81	3.42	0.81
	Min	0.8	4.5	bdl	5.1	1.3	0.0	0.3	0.2	0.0	bdl	0.3	1506.8	0.4	11.1	193.2	bdl	bdl	bdl	0.12	0.18	0.02	0.09	0.04	bdl	0.05	0.01	0.07	0.01	0.06	0.01	0.06	0.01	bdl	0.44	1.68	0.01	0.004	bdl	0.51	0.24
	Median	2.6	5.9	141.8	9.6	1.4	0.0	0.5	0.2	0.1		0.7	1594.3	0.7	23.9	204.2		0.1	0.01	0.18	0.27	0.03	0.22	0.06	0.01	0.07	0.02	0.11	0.03	0.10	0.01	0.11	0.01		0.61	1.84	0.01	0.02		0.86	0.41
	RSD (%)	151.4	243.9	94.0	39.4	32.7	161.0	118.2	257.1	189.3	162.1	74.6	3.5	65.6	36.6	3.9		97.6	162.54	95.86	46.75	83.77	92.32	57.84	113.20	69.78	52.89	39.65	46.21	38.59	49.66	45.00	140.52		19.15	4.46	125.26	50.18	240.64	67.44	37.61
Hostákov	Mean	5.6	9.1	19.8	16.3	1.2	0.1	9.0	3.2	4.1	0.2	1.5	41.2	9.2	97.2	1022.0		0.2	0.14	28.34	61.41	8.23	49.77	9.87	1.11	7.28	0.61	3.15	0.40	0.81	0.05	0.31	0.02	0.002	3.25	1.45	0.13	0.24	0.17	166.02	5.34
n=11	SD	6.7	7.2	21.3	10.8	0.8	0.3	1.8	10.3	13.0	0.6	0.4	4.7	1.7	32.0	721.7		0.3	0.34	8.49	16.90	2.13	13.92	3.66	0.47	2.40	0.18	0.81	0.07	0.15	0.02	0.09	0.01	0.002	1.50	0.56	0.21	0.19	0.17	45.48	1.26
SCH-11	Max	19.5	26.4	66.7	41.7	3.1	1.1	11.3	37.5	47.3	2.1	2.4	50.5	12.5	151.4	2468.6	0.1	1.1	1.25	39.34	90.35	12.66	78.95	16.00	1.95	10.40	0.84	4.46	0.54	1.12	0.08	0.56	0.04	0.01	4.96	2.78	0.71	0.73	0.58	249.65	7.62
	Min	bdl	bdl	bdl	5.4	bdl	bdl	6.2	bdl	bdl	bdl	1.1	35.9	6.4	38.9	568.5	bdl	bdl	bdl	13.71	31.83	4.66	29.00	4.07	0.36	3.20	0.32	1.84	0.26	0.59	0.00	0.22	0.01	bdl	0.95	0.82	bdl	0.05	bdl	86.82	3.32
	Median	2.9	7.5	12.5	11.2	1.2		9.5				1.5	40.2	9.2	107.9	598.4		0.2	0.01	30.26	65.12	8.34	48.44	10.74	1.24	7.99	0.66	3.32	0.40	0.81	0.06	0.28	0.02		3.73	1.37	0.04	0.19	0.11	171.19	5.60
	RSD (%)	118.1	79.1	107.6	66.2	65.8	297.1	20.2	321.7	314.4	283.2	24.7	11.4	18.9	32.9	70.6		113.7	250.55	29.95	27.53	25.82	27.97	37.08	41.84	32.94	28.91	25.65	18.90	18.56	49.80	30.06	41.64	91.36	45.99	38.82	153.61	78.41	98.91	27.39	23.61
Kotel	Mean	5.5	5.5	82.4	11.4	2.9	bdl	1.0	0.7	4.1	77.7	2.2	237.2	14.6	1.8	261.2	0.4	0.1	0.49	0.26	1.19	0.29	2.80	1.57	0.38	2.97	0.42	3.63	0.62	1.73	0.14	0.61	0.04	0.003	0.37	16.33	0.03	0.02	0.10	9.47	7.19
n=12	SD	5.0	3.0	65.3	3.2	3.2	0.1	0.4	1.6	8.4	167.0	1.9	80.7	10.1	0.2	92.5	1.1	0.1	0.99	0.17	0.92	0.23	2.58	1.34	0.19	2.37	0.31	2.52	0.43	1.06	0.08	0.38	0.03	0.002	0.01	35.47	0.04	0.01	0.14	7.54	4.53
SCH-17	Max	13.5	11.4	186.3	17.5	12.5	0.3	1.9	6.0	30.6	514.3	7.4	404.5	32.2	2.1	412.8	3.9	0.3	3.34	0.64	3.20	0.81	9.13	4.67	0.73	8.45	1.08	8.88	1.47	3.68	0.26	1.36	0.12	0.01	0.40	127.88	0.17	0.05	0.40	27.14	16.24
	Min	0.4	2.7	bdl	6.3	0.6	bdl	0.4	0.1	bdl	bdl	0.4	108.0	1.2	1.4	72.1	bdl	0.0	0.00	0.01	0.05	0.02	0.27	0.17	0.04	0.38	0.05	0.42	0.07	0.20	0.02	0.10	0.01	bdl	0.35	1.19	0.01	0.01	0.001	0.94	0.92
	Median	2.1	3.9	68.8	11.3	1.6	bdl	1.0	0.2	0.3	0.1	1.5	244.1	11.3	1.8	245.5		0.1	0.02	0.20	0.90	0.20	1.70	1.13	0.43	2.21	0.32	2.86	0.48	1.57	0.13	0.52	0.03		0.37	2.14	0.01	0.02	0.02	6.78	6.35
	RSD (%)	90.1	55.2	79.3	28.3	108.4		38.0	221.3	208.3	214.9	89.1	34.0	69.0	12.1	35.4	273.2	75.9	201.97	66.94	77.52	77.47	91.94	85.49	51.03	79.56	72.96	69.44	69.28	61.49	56.54	62.54	80.82	55.28	3.31	217.15	138.12	52.53	150.65	79.59	62.98
Vrbík	Mean	16.2	8.7	52.0	9.2	1.0	0.1	6.4	7.2	15.8	0.1	1.7	53.9	2.0	39.0	858.3	0.1	0.1	0.10	8.48	10.43	0.84	2.86	0.44	1.33	0.40	0.03	0.23	0.04	0.20	0.04	0.43	0.07	0.001	0.48	4.80	0.11	0.79	0.93	24.78	1.04
n=12	SD	20.3	10.3	58.0	6.2	0.6	0.2	2.2	15.7	23.4	0.3	1.0	11.4	0.9	22.6	165.6	0.1	0.1	0.23	3.93	5.35	0.59	2.97	0.47	1.11	0.31	0.03	0.16	0.03	0.09	0.02	0.23	0.04	0.004	0.03	1.04	0.15	0.79	0.73	12.54	0.48
SCH-40	Max	73.7	36.1	169.6	22.2	2.1	0.8	11.3	54.2	69.5	1.0	3.8	82.7	4.1	79.2	1187.1	0.3	0.3	0.70	16.22	18.88	2.40	11.79	1.90	4.45	1.38	0.12	0.59	0.12	0.37	0.08	0.95	0.15	0.02	0.53	6.99	0.56	2.86	2.63	44.94	2.12
	Min	1.3	bdl	bdl	bdl	bdl	bdl	4.3	0.1	0.0	0.0	0.3	41.6	0.9	7.4	526.8	bdl	bdl	bdl	1.94	2.78	0.30	0.53	0.08	0.47	0.18	bdl	bdl	bdl	0.08	0.02	0.18	0.03	bdl	0.43	3.63	0.00	0.04	0.00	9.13	0.45
	Median	8.8	2.9	25.7	9.5	1.1		5.5	0.3	0.6		1.7	50.0	1.9	31.1	879.9				7.76	9.06	0.65	1.79	0.31	0.89	0.32	0.02	0.21	0.03	0.20	0.03	0.36	0.06		0.48	4.56	0.05	0.64	0.87	21.72	1.05
	RSD (%)	125.1	118.9	111.4	66.8	55.4	211.2	33.6	217.6	148.0	232.3	59.9	21.2	45.0	57.8	19.3	150.7	148.3	225.65	46.30	51.25	70.25	104.04	105.87	83.87	77.79	112.28	67.46	76.29	45.54	57.28	54.33	58.63	280.34	5.63	21.66	142.57	99.28	78.33	50.60	46.36
Other settings	Mean	29.9	101.6	28.6	9.7	1.9	0.1	18.0	5.1	4.2	1.6	0.7	53.2	13.4	1.8	3.1		0.1	0.04	26.67	58.27	6.70	36.44	7.27	8.72	6.88	0.74	5.27	0.80	2.53	0.31	2.85	0.40	0.003	0.33	3.72	0.04	0.13	0.02	150.96	12.91
Erpružice	SD	31.0	20.3	39.4	6.2	1.2	0.1	2.9	10.2	9.6	3.8	0.8	9.9	6.3	1.0	1.2			0.05	19.51	43.49	4.22	20.50	3.90	5.10	3.78	0.42	2.99	0.45	1.39	0.16	1.47	0.21	0.002	0.01	0.98	0.03	0.10	0.02	89.78	6.83
n=12	Maximum	97.3	139.0	115.4	20.9	4.6	0																																		

APPENDIX 2

Electron microprobe analyses of scheelite from different localities representing different types of mineralization in the Bohemian Massif

Sample	Point	Locality	Type of mineralization/deposit	Oxide			Weight%				Det.Lim ppm(G)			
				CaO	WO3	Total	Na	Fe	Mo	Sr	Na	Fe	Mo	Sr
SCH1 1	SCH1-1	Vacikov. Petráčkova hora	RIRGS	19.2	81.2	100.3	0.02	0.00	0.10	0.00	15	35	37	96
SCH1 2	SCH1-2	Vacikov. Petráčkova hora	RIRGS	19.0	81.4	100.4	0.03	0.00	0.11	0.00	14	35	38	95
SCH1 3	SCH1-3	Vacikov. Petráčkova hora	RIRGS	19.1	81.0	100.1	0.02	0.00	0.12	0.00	14	35	37	96
SCH1 4	SCH1-4	Vacikov. Petráčkova hora	RIRGS	19.3	81.8	101.0	0.02	0.00	0.11	0.00	14	35	37	96
SCH1 5	SCH1-5	Vacikov. Petráčkova hora	RIRGS	19.2	81.5	100.7	0.01	0.01	0.08	0.00	14	35	36	96
SCH1 6	SCH1-6	Vacikov. Petráčkova hora	RIRGS	19.1	81.9	101.1	0.00	0.00	0.13	0.00	15	35	37	97
SCH1 7	SCH1-7	Vacikov. Petráčkova hora	RIRGS	19.2	81.7	100.9	0.01	0.00	0.11	0.00	15	35	37	96
SCH3 1	SCH3 1	Čelina. JP-10. 60 m	RIRGS	18.8	79.5	98.4	0.01	0.00	0.15	0.00	14	34	38	95
SCH3 2	SCH3 2	Čelina. JP-10. 60 m	RIRGS	19.3	80.0	99.3	0.01	0.00	0.24	0.00	15	35	40	94
SCH3 3	SCH3 3	Čelina. JP-10. 60 m	RIRGS	19.3	80.1	99.4	0.01	0.00	0.39	0.00	15	34	43	93
SCH3 4	SCH3 4	Čelina. JP-10. 60 m	RIRGS	19.3	80.1	99.4	0.01	0.00	0.19	0.00	14	35	39	94
SCH3 5	SCH3 5	Čelina. JP-10. 60 m	RIRGS	19.1	80.2	99.3	0.01	0.00	0.16	0.00	14	35	38	96
SCH3 6	SCH3 6	Čelina. JP-10. 60 m	RIRGS	19.1	80.2	99.2	0.01	0.00	0.17	0.00	15	35	39	96
SCH3 7	SCH3 7	Čelina. JP-10. 60 m	RIRGS	19.3	80.0	99.3	0.00	0.00	0.18	0.00	15	34	38	95
SCH3 8	SCH3 8	Čelina. JP-10. 60 m	RIRGS	19.2	80.5	99.7	0.00	0.00	0.15	0.00	15	35	38	96
SCH3 9	SCH3 9	Čelina. JP-10. 60 m	RIRGS	19.3	80.7	100.0	0.01	0.01	0.17	0.00	14	34	38	96
SCH3 10	SCH3 10	Čelina. JP-10. 60 m	RIRGS	19.5	79.4	99.0	0.01	0.00	0.36	0.00	14	35	43	96
SCH3 11	SCH3 11	Čelina. JP-10. 60 m	RIRGS	19.2	80.7	99.9	0.00	0.01	0.15	0.00	15	35	38	95
SCH7 1	SCH7 1	Humpolec. Orlík	Orogenic Au	19.2	80.6	99.7	0.01	0.00	0.00	0.15	15	34	35	100
SCH7 2	SCH7 2	Humpolec. Orlík	Orogenic Au	19.0	80.5	99.5	0.00	0.00	0.00	0.15	15	35	35	99
SCH7 3	SCH7 3	Humpolec. Orlík	Orogenic Au	19.0	80.5	99.4	0.02	0.00	0.00	0.19	15	35	35	98
SCH7 4	SCH7 4	Humpolec. Orlík	Orogenic Au	19.0	80.4	99.4	0.01	0.00	0.00	0.17	15	34	35	102
SCH7 5	SCH7 5	Humpolec. Orlík	Orogenic Au	19.1	80.6	99.6	0.02	0.00	0.00	0.20	14	35	36	100
SCH7 6	SCH7 6	Humpolec. Orlík	Orogenic Au	19.0	80.0	99.0	0.04	0.00	0.00	0.16	15	34	35	102
SCH7 7	SCH7 7	Humpolec. Orlík	Orogenic Au	19.0	80.0	99.0	0.01	0.00	0.00	0.17	15	34	36	100
SCH7 7	SCH7 7	Humpolec. Orlík	Orogenic Au	19.0	80.8	99.8	0.01	0.00	0.00	0.15	14	35	35	101
SCH7 8	SCH7 8	Humpolec. Orlík	Orogenic Au	18.8	79.9	98.7	0.00	0.00	0.00	0.16	15	34	35	100
SCH7 9	SCH7 9	Humpolec. Orlík	Orogenic Au	18.9	79.0	97.9	0.01	0.00	0.00	0.17	15	35	35	102
SCH7 10	SCH7 10	Humpolec. Orlík	Orogenic Au	18.9	80.0	98.8	0.01	0.00	0.00	0.16	14	35	36	102
SCH9 1	SCH9 1	Chedrbí u Čáslavi	Oxidized skarn	19.3	79.7	99.0	0.01	0.00	0.10	0.00	14	35	38	97
SCH9 2	SCH9 2	Chedrbí u Čáslavi	Oxidized skarn	19.3	80.3	99.6	0.00	0.00	0.10	0.00	14	34	38	97
SCH9 3	SCH9 3	Chedrbí u Čáslavi	Oxidized skarn	19.5	80.9	100.3	0.00	0.00	0.05	0.00	14	35	36	96
SCH9 4	SCH9 4	Chedrbí u Čáslavi	Oxidized skarn	19.3	80.9	100.2	0.00	0.00	0.04	0.00	15	35	37	94
SCH9 5	SCH9 5	Chedrbí u Čáslavi	Oxidized skarn	19.4	80.7	100.2	0.00	0.00	0.15	0.00	15	35	37	96
SCH9 6	SCH9 6	Chedrbí u Čáslavi	Oxidized skarn	19.3	79.8	99.1	0.00	0.00	0.11	0.00	15	35	38	97
SCH9 7	SCH9 7	Chedrbí u Čáslavi	Oxidized skarn	19.5	80.4	99.8	0.01	0.00	0.13	0.00	14	35	38	96
SCH9 8	SCH9 8	Chedrbí u Čáslavi	Oxidized skarn	19.5	80.6	100.1	0.00	0.00	0.12	0.00	15	34	38	96
SCH39 1	SCH39 1	Cetoraz u Pacova	Greisen Sn-W and related quartz vein	19.4	81.3	100.8	0.01	0.00	0.00	0.01	14	35	35	99
SCH39 2	SCH39 2	Cetoraz u Pacova	Greisen Sn-W and related quartz vein	19.1	79.8	98.9	0.02	0.00	0.01	0.00	14	35	35	98
SCH39 3	SCH39 3	Cetoraz u Pacova	Greisen Sn-W and related quartz vein	19.4	80.8	100.2	0.01	0.00	0.00	0.00	14	35	35	97
SCH39 4	SCH39 4	Cetoraz u Pacova	Greisen Sn-W and related quartz vein	19.3	81.2	100.5	0.01	0.00	0.00	0.00	15	35	36	98
SCH39 5	SCH39 5	Cetoraz u Pacova	Greisen Sn-W and related quartz vein	19.4	80.4	99.8	0.03	0.00	0.00	0.01	15	35	35	97
SCH39 6	SCH39 6	Cetoraz u Pacova	Greisen Sn-W and related quartz vein	19.2	80.6	99.8	0.04	0.00	0.00	0.00	15	35	36	100
SCH43 1	SCH43 1	Krásno - Huber	Greisen Sn-W and related quartz vein	19.5	81.0	100.5	0.00	0.00	0.00	0.00	15	35	36	97
SCH43 2	SCH43 2	Krásno - Huber	Greisen Sn-W and related quartz vein	19.5	81.1	100.7	0.01	0.00	0.00	0.00	14	35	35	96
SCH43 3	SCH43 3	Krásno - Huber	Greisen Sn-W and related quartz vein	19.6	81.4	100.9	0.00	0.00	0.00	0.00	14	35	36	97
SCH43 4	SCH43 4	Krásno - Huber	Greisen Sn-W and related quartz vein	19.6	80.9	100.4	0.00	0.00	0.00	0.00	15	34	35	98
SCH41 1	SCH41 1	Jílové. 3. Pěpř. ž. Tobola. 1967	Orogenic Au	19.2	81.2	100.4	0.00	0.00	0.00	0.00	15	35	35	94
SCH41 2	SCH41 2	Jílové. 3. Pěpř. ž. Tobola. 1968	Orogenic Au	19.1	81.1	100.2	0.01	0.00	0.00	0.00	15	35	35	96
SCH41 3	SCH41 3	Jílové. 3. Pěpř. ž. Tobola. 1969	Orogenic Au	19.2	80.3	99.5	0.02	0.00	0.01	0.00	15	35	35	95
SCH41 4	SCH41 4	Jílové. 3. Pěpř. ž. Tobola. 1970	Orogenic Au	19.1	81.0	100.1	0.01	0.00	0.00	0.00	15	35	36	95
SCH41 5	SCH41 5	Jílové. 3. Pěpř. ž. Tobola. 1971	Orogenic Au	19.4	80.7	100.1	0.01	0.00	0.00	0.00	15	35	36	94
SCH41 6	SCH41 6	Jílové. 3. Pěpř. ž. Tobola. 1972	Orogenic Au	19.1	80.7	99.8	0.01	0.00	0.01	0.00	15	35	35	93
SCH41 7	SCH41 7	Jílové. 3. Pěpř. ž. Tobola. 1973	Orogenic Au	19.2	81.3	100.5	0.01	0.00	0.00	0.00	15	35	36	94
SCH41 8	SCH41 8	Jílové. 3. Pěpř. ž. Tobola. 1974	Orogenic Au	19.3	80.8	100.0	0.01	0.01	0.00	0.00	15	35	36	95
SCH38 1	SCH38 1	Ovesná Lhota u Světlé nad Sázavou	Greisen Sn-W and related quartz vein	19.0	80.6	99.5	0.03	0.00	0.00	0.00	15	35	35	96
SCH38 2	SCH38 2	Ovesná Lhota u Světlé nad Sázavou	Greisen Sn-W and related quartz vein	19.2	80.7	99.9	0.01	0.00	0.00	0.00	15	35	35	96
SCH38 3	SCH38 3	Ovesná Lhota u Světlé nad Sázavou	Greisen Sn-W and related quartz vein	18.9	78.6	97.5	0.04	0.00	0.00	0.00	15	35	35	93
SCH38 4	SCH38 4	Ovesná Lhota u Světlé nad Sázavou	Greisen Sn-W and related quartz vein	19.1	80.2	99.3	0.03	0.00	0.00	0.00	15	35	36	94
SCH10-1	SCH10-1	Mutice u Mladé Vožice	Oxidized skarn	19.1	80.6	99.8	0.00	0.01	0.02	0.00	15	35	37	91
SCH10-2	SCH10-2	Mutice u Mladé Vožice	Oxidized skarn	19.1	79.5	98.6	0.00	0.00	0.04	0.00	15	35	37	90
SCH10-3	SCH10-3	Mutice u Mladé Vožice	Oxidized skarn	19.2	78.9	98.1	0.00	0.00	0.03	0.00	14	35	37	90
SCH10-4	SCH10-4	Mutice u Mladé Vožice	Oxidized skarn	19.2	79.9	99.2	0.00	0.00	0.04	0.00	15	35	37	93
SCH10-5	SCH10-5	Mutice u Mladé Vožice	Oxidized skarn	19.3	79.6	98.9	0.01	0.00	0.01	0.00	15	35	37	90
SCH10-6	SCH10-6	Mutice u Mladé Vožice	Oxidized skarn	19.2	79.3	98.5	0.00	0.00	0.03	0.00	15	35	37	92
SCH10-7	SCH10-7	Mutice u Mladé Vožice	Oxidized skarn	19.2	79.4	98.7	0.00	0.01	0.04	0.00	15	35	37	91
SCH10-8	SCH10-8	Mutice u Mladé Vožice	Oxidized skarn	19.3	79.5	98.7	0.00	0.00	0.02	0.00	15	35	37	88
SCH10-9	SCH10-9	Mutice u Mladé Vožice	Oxidized skarn	19.2	79.0	98.2	0.00	0.00	0.03	0.00	15	35	36	90
SCH10-10	SCH10-10	Mutice u Mladé Vožice	Oxidized skarn	19.3	80.1	99.3	0.00	0.00	0.02	0.00	15	35	36	90
SCH13-1	SCH13-1	Budislav u Litomyšle	Reduced skarn	19.2	80.0	99.2	0.00	0.00	0.00	0.00	15	35	36	91
SCH13-2	SCH13-2	Budislav u Litomyšle	Reduced skarn	19.2	80.6	99.8	0.00	0.00	0.01	0.00	15	35	36	90
SCH13-3	SCH13-3	Budislav u Litomyšle	Reduced skarn	19.3	79.8	99.0	0.00	0.00	0.01	0.00	15	35	36	89
SCH13-4	SCH13-4	Budislav u Litomyšle	Reduced skarn	19.3	79.8	99.0	0.00	0.00	0.00	0.00	15	35	36	91
SCH13-5	SCH13-5	Budislav u Litomyšle	Reduced skarn	19.2	80.0	99.3	0.01	0.00	0.02	0.00	15	35	36	92
SCH13-6	SCH13-6	Budislav u Litomyšle	Reduced skarn	19.1	80.5	99.6	0.00	0.00	0.00	0.00	15	35	36	90
SCH13-7	SCH13-7	Budislav u Litomyšle	Reduced skarn	19.2	80.4	99.5	0.00	0.00	0.00	0.00	15	35	36	91
SCH29 1	SCH29 1	Nekvasovy-Chlumy	Oxidized skarn	19.3	81.5	100.8	0.01	0.00	0.01	0.00	15	35	36	97
SCH29 2	SCH29 2	Nekvasovy-Chlumy	Oxidized skarn	19.4	81.5	100.8	0.00	0.00	0.03	0.00	15	35	37	97
SCH29 3	SCH29 3	Nekvasovy-Chlumy	Oxidized skarn	19.3	81.5	100.8	0.00	0.00	0.05	0.00	15	35	38	97
SCH29 4	SCH29 4	Nekvasovy-Chlumy	Oxidized skarn	19.3	81.7	101.0	0.01	0.00	0.03	0.00	14	35	36	98
SCH29 5	SCH29 5	Nekvasovy-Chlumy	Oxidized skarn	19.2	80.5	99.7	0.01	0.00	0.03	0.00	15	35	37	97
SCH29 6	SCH29 6	Nekvasovy-Chlumy	Oxidized skarn	19.4	81.1	100.6	0.00	0.01	0.05	0.00	15	35	36	95
SCH29 7	SCH29 7	Nekvasovy-Chlumy	Oxidized skarn	19.3	81.1	100.5	0.00	0.00	0.07	0.00	15	35	38	96
SCH29 8	SCH29 8	Nekvasovy-Chlumy	Oxidized skarn	19.3	80.8	100.2	0.00	0.00	0.04	0.00	14	35	37	97
SCH29 9	SCH29 9	Nekvasovy-Chlumy	Oxidized skarn	19.2	80.9	100.1	0.00	0.00	0.07	0.00	15	35	37	96
SCH29 10	SCH29 10	Nekvasovy-Chlumy	Oxidized skarn	19.2	80.9	100.1	0.01	0.00						

SCH46 7	SCH46 7	Proseč u Litomyšle	Oxidized skarn	19.6	81.7	101.2	0.01	0.00	0.02	0.00	14	35	36	95
SCH46 8	SCH46 8	Proseč u Litomyšle	Oxidized skarn	19.6	81.3	100.9	0.01	0.00	0.02	0.00	15	35	36	96
SCH45 1	SCH45 1	Krouna u Poličky	Oxidized skarn	19.4	80.7	100.1	0.00	0.00	0.04	0.00	15	35	36	95
SCH45 2	SCH45 2	Krouna u Poličky	Oxidized skarn	19.4	80.2	99.6	0.03	0.00	0.03	0.00	14	35	37	97
SCH45 3	SCH45 3	Krouna u Poličky	Oxidized skarn	19.3	80.0	99.3	0.02	0.00	0.05	0.00	15	35	36	95
SCH45 4	SCH45 4	Krouna u Poličky	Oxidized skarn	19.3	80.5	99.8	0.00	0.00	0.03	0.00	15	35	37	98
SCH45 5	SCH45 5	Krouna u Poličky	Oxidized skarn	19.3	80.6	99.9	0.00	0.00	0.03	0.00	15	35	37	97
SCH45 6	SCH45 6	Krouna u Poličky	Oxidized skarn	19.3	79.7	99.1	0.00	0.00	0.06	0.00	15	35	36	97
SCH36 1	SCH36 1	Tehov u Řičan	Quartz vein	19.3	79.4	98.6	0.01	0.00	0.00	0.00	15	35	36	95
SCH36 2	SCH36 2	Tehov u Řičan	Quartz vein	19.2	79.4	98.7	0.01	0.01	0.00	0.00	14	34	35	95
SCH36 3	SCH36 3	Tehov u Řičan	Quartz vein	19.3	79.7	98.9	0.00	0.00	0.00	0.00	15	35	36	96
SCH36 4	SCH36 4	Tehov u Řičan	Quartz vein	19.2	79.1	98.3	0.01	0.00	0.00	0.00	14	35	36	97
SCH36 5	SCH36 5	Tehov u Řičan	Quartz vein	18.6	74.5	93.0	0.01	0.00	0.00	0.00	14	34	36	94
SCH36 6	SCH36 6	Tehov u Řičan	Quartz vein	19.3	79.2	98.5	0.00	0.00	0.01	0.00	15	35	35	97
SCH36 7	SCH36 7	Tehov u Řičan	Quartz vein	19.1	78.9	98.0	0.00	0.00	0.00	0.00	15	35	36	95
SCH36 8	SCH36 8	Tehov u Řičan	Quartz vein	19.1	78.9	98.0	0.01	0.00	0.00	0.00	14	35	36	96
SCH48 1	SCH48 1	Čínovec 2	Greisen Sn-W and related quartz vein	19.6	75.7	95.4	0.00	0.00	0.52	0.00	15	34	45	94
SCH48 3	SCH48 3	Čínovec 2	Greisen Sn-W and related quartz vein	19.8	75.5	95.3	0.00	0.01	0.52	0.00	14	34	45	95
SCH48 4	SCH48 4	Čínovec 2	Greisen Sn-W and related quartz vein	19.4	77.2	96.7	0.00	0.00	0.14	0.00	14	34	38	96
SCH48 5	SCH48 5	Čínovec 2	Greisen Sn-W and related quartz vein	17.7	76.4	94.1	0.01	0.02	0.12	0.00	14	35	39	94
SCH48 6	SCH48 6	Čínovec 2	Greisen Sn-W and related quartz vein	19.6	75.9	95.5	0.00	0.00	0.54	0.00	14	34	45	94
SCH48 7	SCH48 7	Čínovec 2	Greisen Sn-W and related quartz vein	20.2	76.7	96.9	0.00	0.00	0.09	0.00	14	35	38	93
SCH47 2	SCH47 2	Čínovec 1	Greisen Sn-W and related quartz vein	18.5	74.9	93.3	0.02	0.03	0.25	0.00	14	34	41	93
SCH47 3	SCH47 3	Čínovec 1	Greisen Sn-W and related quartz vein	18.6	74.1	92.7	0.00	0.02	0.41	0.00	14	34	42	91
SCH42 1	SCH42 1	Sobětice u Klatov	Orogenic Au	19.0	79.3	98.3	0.01	0.01	0.00	0.00	15	35	35	94
SCH42 2	SCH42 2	Sobětice u Klatov	Orogenic Au	19.2	80.2	99.4	0.01	0.00	0.00	0.00	14	35	35	98
SCH42 3	SCH42 3	Sobětice u Klatov	Orogenic Au	19.1	80.0	99.1	0.01	0.00	0.00	0.00	14	35	36	97
SCH42 4	SCH42 4	Sobětice u Klatov	Orogenic Au	19.2	80.0	99.2	0.02	0.00	0.00	0.00	15	35	36	97
SCH42 5	SCH42 5	Sobětice u Klatov	Orogenic Au	19.4	79.8	99.2	0.00	0.00	0.00	0.00	15	35	36	96
SCH33 1	SCH33 1	Supíkovice u Jeseníku	Orogenic Au	19.3	80.0	99.3	0.00	0.00	0.03	0.00	14	35	36	94
SCH33 2	SCH33 2	Supíkovice u Jeseníku	Orogenic Au	19.2	80.5	99.7	0.00	0.00	0.02	0.00	14	35	36	96
SCH33 3	SCH33 3	Supíkovice u Jeseníku	Orogenic Au	19.2	78.6	97.7	0.00	0.00	0.03	0.00	14	35	36	96
SCH33 4	SCH33 4	Supíkovice u Jeseníku	Orogenic Au	19.4	79.9	99.3	0.00	0.00	0.02	0.00	14	35	36	95
SCH33 5	SCH33 5	Supíkovice u Jeseníku	Orogenic Au	19.2	80.5	99.6	0.00	0.00	0.03	0.00	14	35	36	95

Data for Obří důl, Kovářská and Vykmánov (oxidized skarns) and Kotel, Vrbík, Hazlov and Hostákov (reduced skarns) are from Miranda et al. (2024)

APPENDIX 3

Detailed laser setting specifications

Laser make, model and type	Resolution M-50 Excimer
Laser wavelength	193 nm
Laser pulse duration	5 ns
Laser energy density	3 J cm ⁻²
Laser repetition rate	15 Hz
Laser spot diameter	33 to 55 µm spot
Ablation cell	S-155 (Laurin Technic)
He carrier gas flow	350 ml/min
N2 additional gas	2 ml/min
Gas blank acquisition time	30 s
ICP-MS model and type	7900 Agilent quadrupole
Ar plasma gas flow	14 L/min
Ar auxiliary gas flow	0.0 L/min
Ar makeup gas flow	0.0 L/min
Injector	3.0 mm i.d., quartz
Torch depth/z-position	0.4 mm
RF power	1550 W
Cone	Ni-Pt
m/z monitores (amu)	11, 23, 24, 39, 44, 49, 51, 55, 57, 59, 61, 63, 66, 75, 85, 88, 89, 93, 95, 107, 118, 137, 139, 140, 141, 146, 147, 153, 157, 159, 163, 165, 166, 169, 172, 175, 181, 182, 183, 208, 209, 232, 238.
Dwell time (ti)	5.00 ms
Quad. settling time (τ)	1.90 ms
Total sweep time (Tn)	400 ms

APPENDIX 4

Reference materials

Element	Li7	B11	Na23	Mg24	Si29	K39	Ti49	V51	Cr53	Mn55	Co59	Ni60	Cu63	Zn66	As75	Sr88	Y89	Nb93	Mo95	Ag107	Sn118	Ba137	La139	Ce140	Pr141	Nd146	Sm147	Eu153	Gd157	Tb159	Dy163	Ho165	Er166	Tm169	Yb172	Lu175	Ta181	Pb208	Bi209	Th232	U238
GSD-1 g Working value	43	50	26708.5	21709	248656.8	25300	7431.7	44	42	220	40	58	42	54	27	69.4	42	42	39	23	29	67	39.1	41.4	45	44.7	47.8	41	50.7	47	51.2	49	40.1	49	50.9	51.5	40	50	35	41	41
Stdev	6	20	1483.8	241.2	3739.2	300	359.6	2	3	20	2	4	2	2	8	0.7	2	3	3	3	6	1	0.4	0.4	1	0.5	0.5	2	0.5	2	0.5	2	0.4	2	0.5	0.5	4	2	4	2	2
Average GSD-1 g	48.7	71	27949	22485	274633	26780	7123	43	41.9	228	38.3	55.9	41.8	55.8	37	69	39	37	39	22.1	33.0	68	39	39	41	44	47	40	48	45	50	47	37	45	50	48	39	49	31	39	40
Stdev (n = 25)	2.1	6	1120	2254	5900	887	104	1	1.9	11	1.0	1.9	1.2	2.5	3	5	3	3	3	0.8	1.1	3	4	4	4	1	1	3	1	4	1	4	1	4	1	4	3	3	3	5	4
RSD* (%)	4%	9%	4%	10%	2%	3%	1%	3%	5%	5%	3%	3%	3%	4%	9%	8%	8%	9%	8%	4%	3%	4%	9%	10%	10%	2%	2%	7%	2%	9%	2%	9%	3%	9%	2%	8%	9%	7%	10%	13%	9%
GSE-1 g Working value	430	330	28934	21106	250994	21800	450	440	400	590	380	440	380	460	260	447	410	420	390	200	280	427	392	414	460	453	488	410	514	480	524	501	595	500	520	518	390	378	320	380	420
Stdev	60	120	1484	181	7011	200	42	20	80	20	20	30	40	10	90	5	30	40	30	20	50	5	4	4	10	5	5	20	6	20	6	8	6	20	5	6	40	12	30	20	30
Average GSE-1g	467.7	416	29774	21296	270300	23984	416	422	384.3	642	387.3	434.2	396.2	487.6	340	443	398	384	393	197.6	331.6	423	405	411	450	452	491	413	511	501	521	515	571	504	521	526	425	384	316	378	419
Stdev (n = 25)	12.3	16	1060	1684	6133	993	8	10	8.3	16	8.5	11.0	10.1	9.2	41	19	11	11	15	6.2	9.7	12	9	9	12	14	13	14	15	15	15	14	21	14	16	16	14	17	16	11	13
RSD (%)	3%	4%	4%	8%	2%	4%	2%	2%	2%	2%	2%	3%	3%	2%	12%	4%	3%	3%	4%	3%	3%	3%	2%	2%	3%	3%	3%	3%	3%	3%	3%	4%	3%	3%	3%	3%	4%	5%	3%	3%	
Gprobe6-A Working value	6	n.d	15506	51318	224259	1577	7012	238	300.2	1255	46.9	145.7	89.5	70.8	n.d	166.6	19.33	4.16	n.d	n.d	1.33	173	5.2	12	1.69	8.4	2.36	0.98	2.87	0.51	3.3	0.72	2.1	0.31	2.06	0.32	0.28	3.28	n.d	0.33	0.29
Stdev	1.39	n.d	1335	6935	9418	166	1319	27.1	21.5	93	3.51	18.9	20.31	16.6	n.d	26.1	1.78	0.41	n.d	n.d	0.71	25.62	0.56	0.7	0.12	0.98	0.31	0.13	0.46	0.06	0.46	0.09	0.34	0.05	0.33	0.04	0.03	0.78	n.d	0.04	0.06
Average Gprobe6-A	6.1		16431	53166	256466.7	1815	6436	255	330.6	1324	52.9	168.8	97.7	96.4		163	17	3		0.6	1.3	167	5	10.9	1.50	8.08	2.29	0.99	2.94	0.48	3.18	0.67	1.94	0.27	2.00	0.29	0.23	3.72		0.31	0.31
Stdev (n = 25)	1.6		1175	5970	5800.0	221	55	5	7.9	43	1.5	5.3	3.3	3.4		12	2	0.3		0.1	0.2	3	0.5	1	0.15	0.38	0.11	0.11	0.13	0.05	0.11	0.08	0.09	0.03	0.08	0.03	0.04	0.28		0.07	0.04
RSD (%)	26%		7%	11%	2%	12%	1%	2%	2%	3%	3%	3%	3%	3%		7%	9%	8%		12%	14%	2%	10%	10%	10%	5%	5%	11%	4%	10%	4%	12%	5%	12%	4%	10%	18%	8%		22%	13%
NIST-612 Working value	42	35	103858	77	n.d	66.3	44	39	35	38	35.5	38.8	37.7	38	37	78.4	38	40	38	22	38	39.7	35.8	38.7	37.2	35.9	38.1	35	36.7	36	36	38	38	38	39.2	36.9	40	38.57	30	37.79	37.38
Stdev	6	6	n.d	60	n.d	0.8	10	8	3.3	2	1.2	0.2	0.9	8	14	0.2	4	6	4	0.3	4	0.8	0.8	0.8	1.8	0.8	0.8	2	0.8	6	0.8	2	1.8	2	0.8	0.8	4	0.2	12	0.08	0.08
Average NIST-612	43.4	41	101657	62		68	39.8	37.6	35.7	40.4	33.9	37.5	35.7	34.9	32.7	75.0	35.5	32.7	37.1	19.7	34.9	38.2	34.9	35.2	33.7	35.9	38.1	35.2	38.4	35.2	35.5	35.4	35.6	33.3	38.7	33.9	34.9	37.1	29.7	35.1	34.5
Stdev (n = 25)	2.0	8	1471	2		8	1.3	1.0	1.8	2.2	1.0	1.5	1.2	1.8	1.2	6.3	3.3	3.1	2.8	0.7	1.8	1.1	3.7	3.8	3.8	0.8	0.9	2.7	0.8	3.7	0.8	3.7	1.8	3.6	0.8	3.4	3.9	1.9	3.1	3.5	3.5
RSD (%)	5%	19%	1%	4%		12%	3%	3%	5%	5%	3%	4%	3%	5%	4%	8%	9%	10%	8%	3%	5%	3%	11%	11%	11%	2%	2%	8%	2%	10%	2%	10%	5%	11%	2%	10%	11%	5%	10%	10%	10%

* – relative standard deviation